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Carbon and oxygen isotopic paleoceanography of the Indian and South Atlantic Oceans — Paleoclimate and paleo-ocean circulation —

By

Koji SETO

With 7 Tables, 69 Text-figures and 1 Appendix

(Received, May 31, 1995)

Abstract: Ocean circulation is intimately associated with continental arrangement and global climate. The purposes of this study are to reconstruct the water mass structure and the deep ocean circulation in the Indian and South Atlantic Oceans during the Cenozoic.

Oxygen and carbon isotopes were studied in Cenozoic sediments at six sites (Sites 752, 754, 756, 757, 758, and 762) in ODP (Ocean Drilling Program) Legs 121 and 122 in the northeastern Indian Ocean. These isotopic records are related to global events occurring in middle Miocene, the Eocene / Oligocene boundary, middle to late Eocene, and the Paleocene / Eocene boundary. To compile those records along with a number of published isotopic data from the Indian and South Atlantic Oceans, adjustments to isotopic ratios have been calculated for different foraminiferal species, and benthic and planktonic foraminiferal isotopic data converted into δ values of dissolved inorganic carbon (DIC) of marine water. The general trends of oxygen and carbon isotopic values show an increase to the south.

Averaged values in one million year intervals of oxygen and carbon isotopes were calculated for each ODP and DSDP (Deep Sea Drilling Project) site, and the time and spatial distributions of the oxygen and carbon isotopic values were examined from the estimated paleodepth. In the Paleocene ocean, the vertical distribution of isotopic ratios is uniform. However, notable negative shift in oxygen isotopic the remarkable in the Miocene are recognized at about 1500m paleodepth in the northeastern Indian Ocean. The source of the water masses are assumed to be as follows: AABW (Antarctic Bottom Water) or proto-AABW formed in the Southern Ocean (Atlantic sector) throughout the Cenozoic. In the Paleocene, another water mass may have formed at low latitudes including the Tethyan Sea, and this water may could have been warm and highly saline, judging from oxygen isotopic ratios. This water mass corresponds to WSDW (Warm Saline Deep Water), which have encountered Proto-AABW at mid latitudes during the early Paleogene. This water mass rapidly reduced in size with the closing of the Tethyan Sea at the Paleocene / Eocene boundary, but still continued to 50 Ma in the Indian Ocean and to 40 Ma in the South Atlantic Ocean. AAIW (Antarctic Intermediate Water) developed from the Oligocene (30 Ma) in the Indian Ocean. Proto-NADW (Proto-North Atlantic Deep Water) distinctly developed from the late Pliocene (3 Ma).

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I. Introduction

It is generally believed that ocean circulation is intimately associated with global climate and continental arrangement. In order to reconstruct of paleoclimate and paleoceanography, oxygen and carbon isotopic analyses have been measured in deep sea sediments, as those isotopes are can be used to estimate paleotemperature and tracer of ocean circulation. Urey (1947) first pointed out that the oxygen isotopic composition of fossils can be used to determination paleotemperature, and the first paleotemperature determinations were published from belemnite shells of the Peedee Formation (Urey et al., 1951). Epstein et al. (1951) empirically determined the relationship between the oxygen isotopic composition of mollusk-shells and growth temperature as a temperature scale. This scale was adjusted by Craig (1965) and Horibe and Oba (1969). Woodruff and Savin (1989) showed that the distribution pattern of fossil foraminiferal $\delta^{13}\text{C}$ values during the Holocene is similar to the $\delta^{13}\text{C}$ pattern of dissolved inorganic carbon (DIC) in the

modern ocean which is related to deep water formation (Kroopnick, 1985). Based on this evidence, they proposed the existence of Tethyan Indian Saline Water (TISW), which flowed from the Tethys into the northern Indian Ocean.

Cenozoic oxygen and carbon isotopic records at many DSDP (Deep Sea Drilling Project) and ODP (Ocean Drilling Program) sites have been published (Shackleton et al., 1984; Oberhansli et al., 1984; Poore and Matthews, 1984; Vincent et al., 1985; Oberhansli, 1986; Miller et al., 1989; Stott et al., 1990; Kennett and Stott, 1990; Stott and Kennett, 1990; Barrera and Huber, 1990; 1991; Katz and Miller, 1991; Woodruff et al., 1990; Boersma and Mikkelsen, 1990; Woodruff and Chamber, 1991; Vincent et al., 1991; Zachos et al., 1992a; 1992b; Rea et al., 1991).

A continuous record of post late Maastrichtian sediments were also recovered from various water depths in the Broken and Ninetyeast Ridge (Leg 121) and the Exmouth Plateau (Leg 122) in the northeastern Indian Ocean. In this area, the isotopic studied using foraminiferal tests have been made by Vincent et al. (1985), Oberhansli (1986), Rea et al. (1991), Seto et al. (1991), and Nomura et al. (1992).

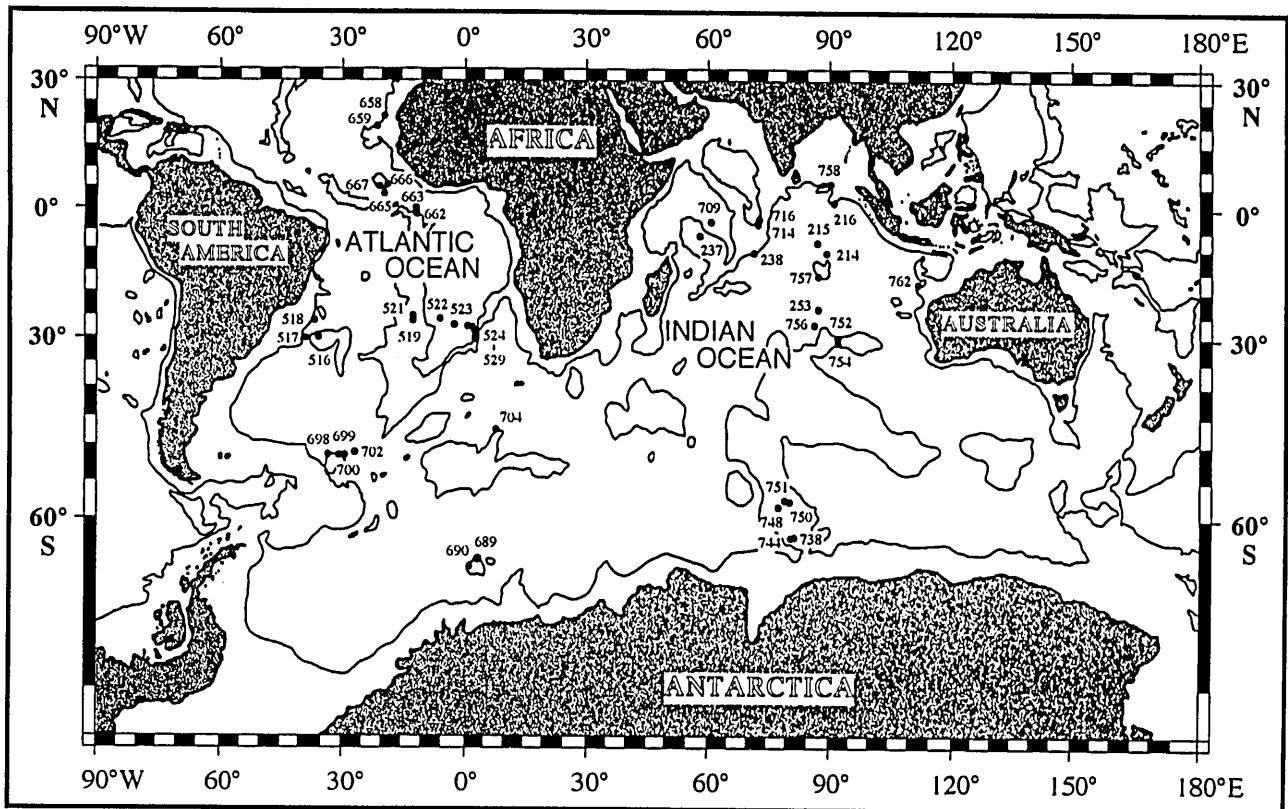


Fig. 1. Location of DSDP and ODP sites examined in this study.

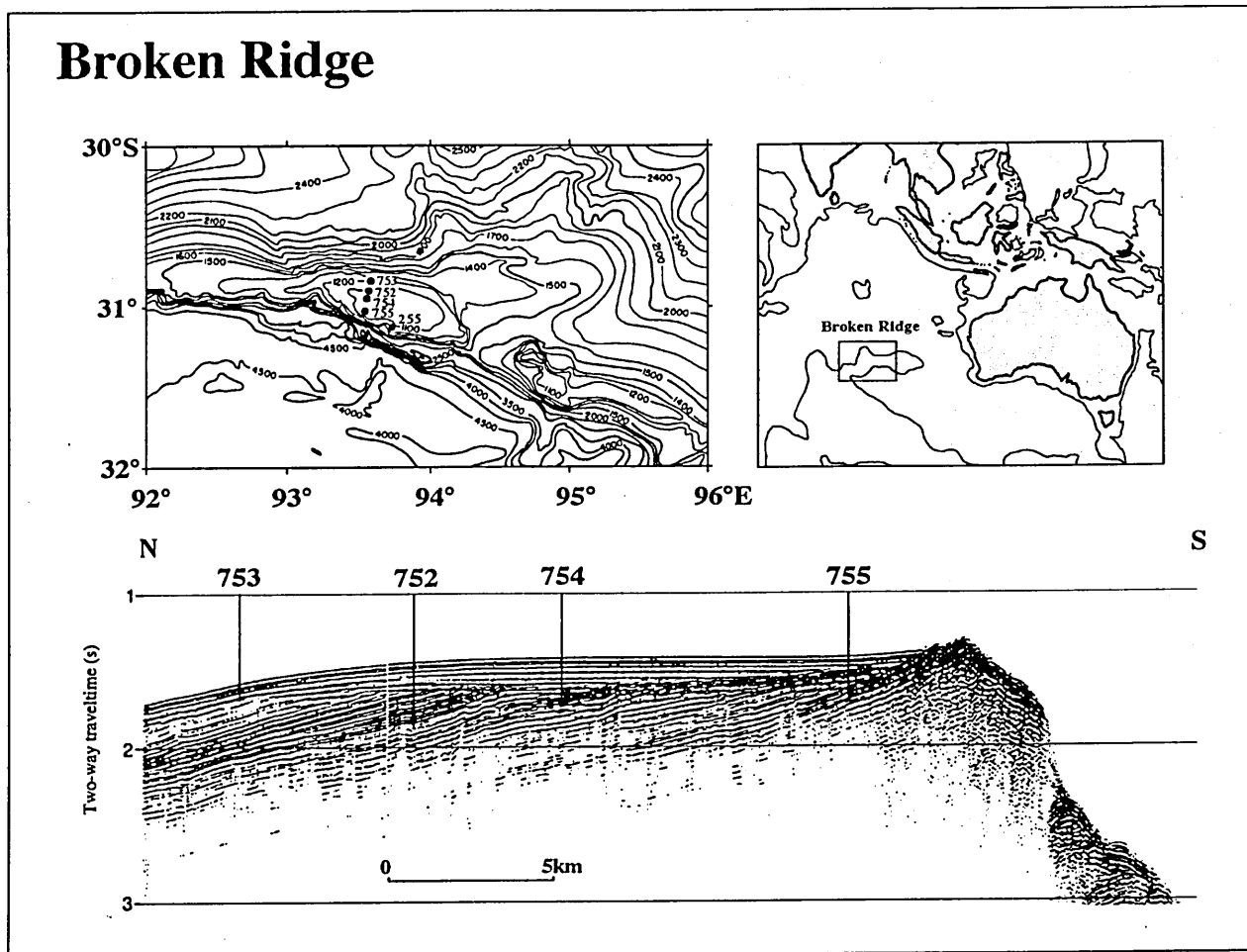


Fig. 2. Bathymetric map and single-channel seismic-reflection profile (RC2708 line 20) across Broken Ridge showing locations of ODP Leg 121 sites at Broken Ridge.

However, these studies mainly focused on the Neogene, and few Paleogene isotope data have been published. Therefore, in this study, oxygen and carbon isotope changes after the late Maastrichtian at six sites (Sites 752, 754, 756, 757, 758, and 762) within Legs 121 and 122 have been examined. These records record global events such as the sharp increase in $\delta^{18}\text{O}$ values near the middle Miocene and the Eocene / Oligocene boundary, the increase of $\delta^{18}\text{O}$ values in the Eocene (Miller et al., 1987), the chron-6 shift and the chron-16 shift of $\delta^{13}\text{C}$ values (Vincent et al., 1980; 1985), and the drastic change of $\delta^{13}\text{C}$ values across the Paleocene / Eocene boundary. These events resulted from climate change and/or changes in ocean circulation.

The purposes of this study are to reconstruct the water-mass structure and the ocean circulation of the deep sea during the Cenozoic in the Indian and South Atlantic Oceans, including the Southern Ocean which should be the main source region of deep water during the Cenozoic.

II. The isotopic record in ODP Legs 121 and 122

A. Samples

Sediment samples for isotope analysis were obtained from ODP Sites 752 and 754 (Broken Ridge), Sites 756, 757, and 758 (Ninetyeast Ridge), and Site 762 (Exmouth Plateau), in the northeast Indian Ocean (Fig. 1). Initial description of these sites have been made by Peirce, Weissel, et al. (1989) and Haq, B. U., von Rad, U., et al. (1990).

Site 752 ($30^{\circ} 53.475' \text{ S}$, $93^{\circ} 34.652' \text{ E}$) is located near the northern edge of Broken Ridge with a present water depth of 1086 m (Fig. 2). A 436-m-long section of sediments from the Pleistocene through to the upper Maastrichtian was recovered. Hole 752A was cored with an advanced hydraulic piston corer (APC) and an extended core barrel (XCB), until refusal at 308 m below seafloor (mbsf). Average core recovery of all core was 70.6%, but recovery for Cores 121-752-1H to 121-752A-11H and Cores 121-752A-26X to 121-752A-33X were 95.8% and 82.2%, respectively. Hole 752B was cored using a rotary core barrel (RCB) to a total depth of 436 mbsf, with an average recovery of 71% over the cored interval. Neogene and late Oligocene sediments were composed of foraminiferal ooze and nannofossil - foraminiferal or foraminiferal - nannofossil ooze. The Paleogene and late Maastrichtian sediments are light green or gray nannofossil calcareous chalk with fine laminations and bioturbation (Fig. 3). The lithology above the Cretaceous / Tertiary boundary at 358.75 mbsf is dark green volcanic ash

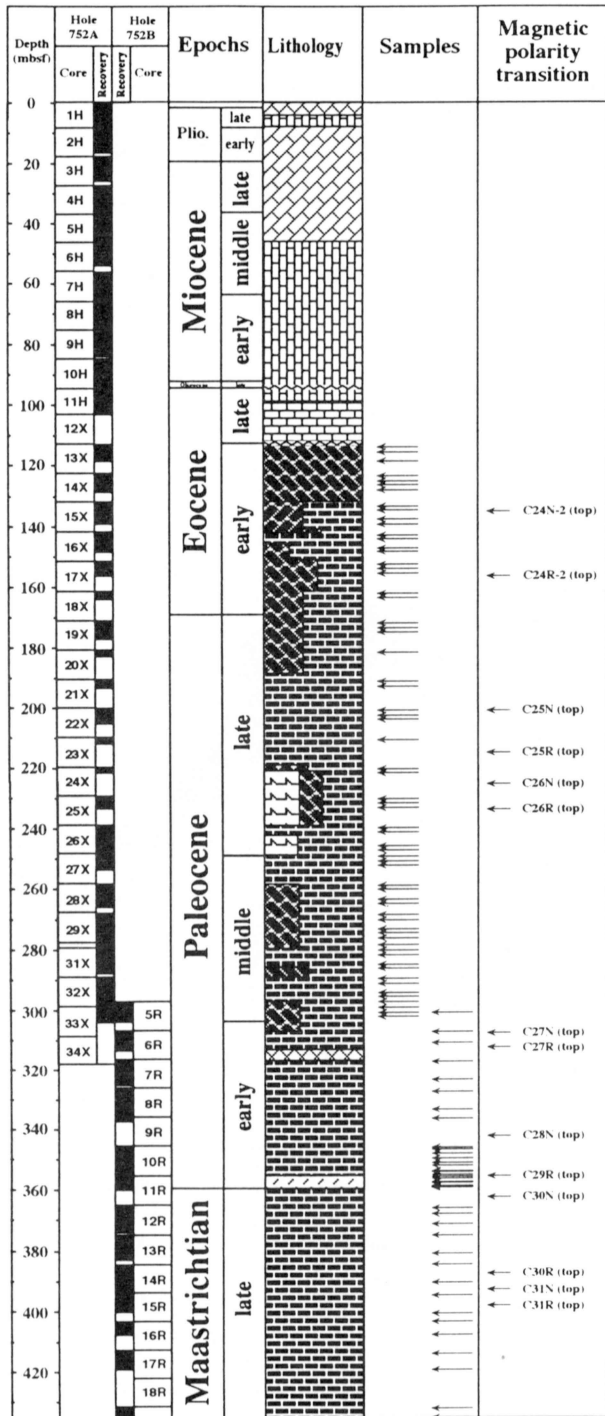


Fig. 3. Lithostratigraphy and Sample horizons at Site 752. See legend of Fig. 4.

layers (Rea et al., 1990). A middle Eocene angular unconformity is marked by coarse-grained sediments with molluscan shell fragments (Peirce, Weissel, et al., 1989; Rea et al., 1990) and foraminifer including small numbers of inner neritic species such as *Amphistegina* (Peirce, Weissel, et al., 1989). An unconformity within the nannofossil - foraminiferous or foraminiferous - nannofossil ooze was recognized during the late Eocene to late Oligocene. Sediment samples examined in this study include Samples 121-752A-

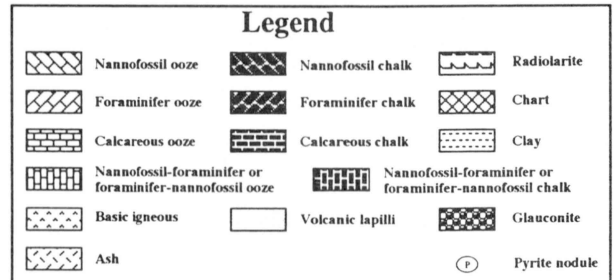
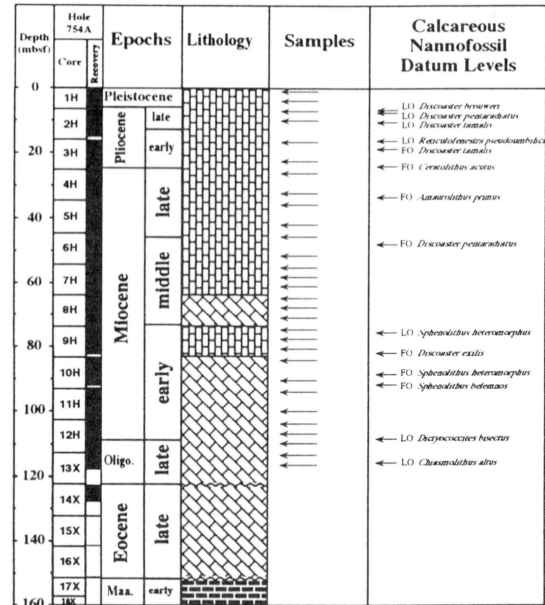


Fig. 4. Lithostratigraphy and Sample horizons at Site 754.

13X-1 through 121-752A-33X-3 (113.6-301.97; 63 samples) from the early Eocene to middle Paleocene in Hole 752A, Samples 121-752A-5R-3 through 121-752A-19R-3 (300.5-435.09 mbsf; 38 samples) from the early Paleocene to late Maastrichtian sediments (including the Cretaceous / Tertiary boundary) in Hole 752C. The location of these samples is shown in Fig 3.

Site 754 (30° 56.439'S, 93° 33.991' E) is located on the central part of Broken Ridge at a present water depth of 1075m, which is the shallowest among the sites of this study (Fig. 2). Hole 754A was cored using the APC, XCB, and Navidrill (NCB) systems to a depth of 172 mbsf. Although average core recovery was 75.5%, Cores 121-754A-1H through 121-754A-13X were almost completely recovered. Pleistocene through late Oligocene and late Eocene sediments consist of white nannofossil - foraminiferous or foraminiferous - nannofossil ooze and white - yellowish brown nannofossil ooze, which are unconformably underlain by early Maastrichtian light gray to greenish gray calcareous chalk with planar and cross-bedded laminae (Fig. 4). An angular unconformity caused by uplift during the middle Eocene was recorded (Fig. 2). A late Eocene to late Oligocene unconformity was recognized in the nannofossil-foraminiferous or foraminiferous - nannofossil ooze. At Site 754, sediment samples used in this study were Samples 121-754A-1H-1

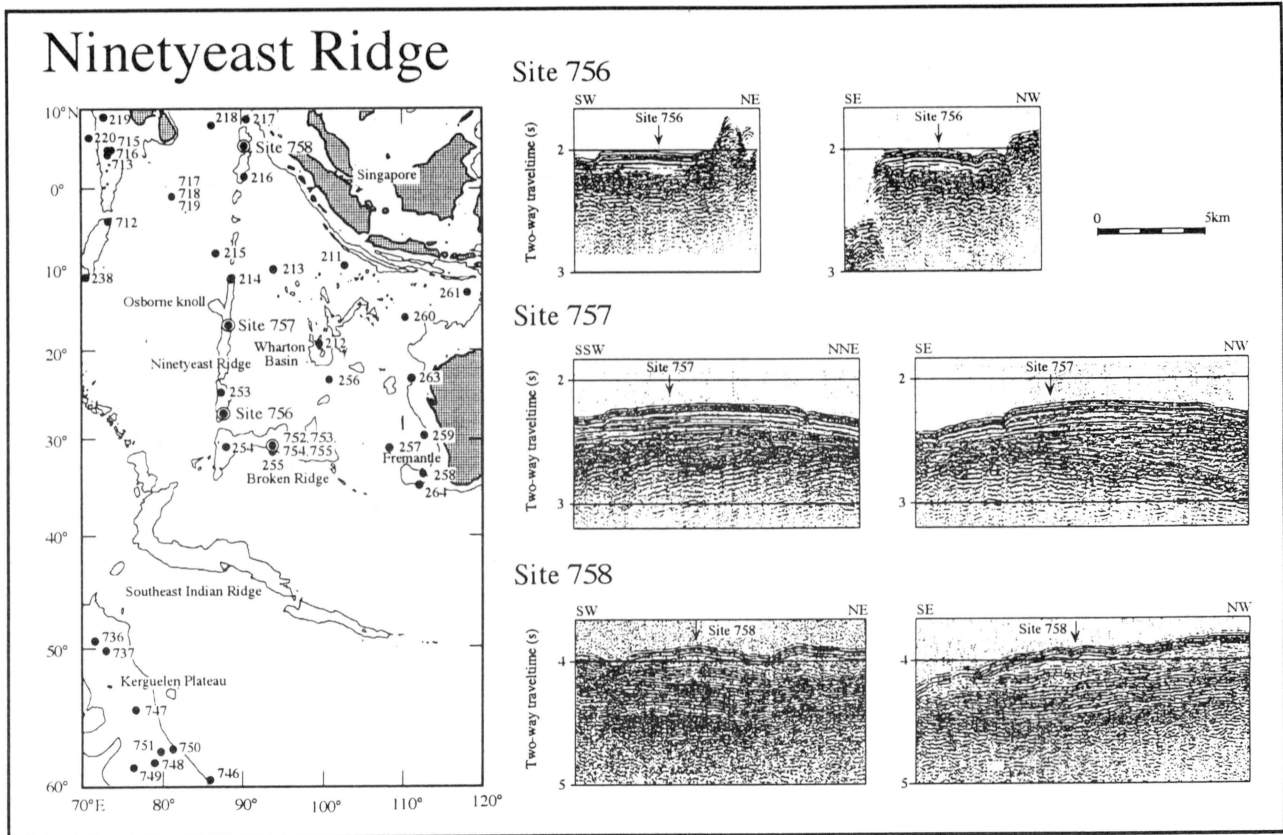


Fig. 5. Bathymetric map and seismic-reflection profile across Sites 756, 757, and 758 at Ninetyeast Ridge.

through 121-754A-13X-3 (0.7-116 mbsf; 29 samples) spanning the Neogene to Late Oligocene (Fig. 4).

Site 756 (27°21.330'S, 87°35.805' E) is located near the crest of southern end of Ninetyeast Ridge at a present water depth of 1518m (Fig. 5). Pleistocene through late Eocene sediments (228-m-thick) were recovered at this site 756. Hole 756B was cored with the APC to 104 mbsf. Hole 756C was washed to 101 mbsf with the XCB system to 150 mbsf (80.7% recovery). Pleistocene to late Eocene sediments consist of nannofossil ooze with foraminifer overlying basaltic basement. At Site 756, 33 sediment samples were analyzed comprising Samples 121-756B-1H-1 through 121-756B-11H-1 (0.7-101.3 mbsf) from the Neogene and Oligocene at Hole 756B, and Sample 121-756C-4X-1 through 121-756C-7X-5 from the early Oligocene and late Eocene at Hole 756C (Fig. 6).

Site 757 (27°21.330'S, 87°35.805' E) is near the crest of central part of Ninetyeast Ridge at a present water depth of 1652m (Fig. 5). Hole 757B was drilled to a depth of 375 m by the APC and XCB systems and recovered a section ranging from Pleistocene to late Paleocene, including the basement, with an average recovery, of 72.6%. Pleistocene through middle Eocene sediments consist of mainly nannofossil ooze, and early Eocene calcareous ooze and chalk overlying pre-Eocene volcanic ash and basement (Fig. 7). At Site 757, 55 Pleistocene to late Paleocene sediment samples were analyzed comprising Samples 121-757B-1H-2 through 121-757B-24X-4 (2.03-216.78 mbsf), (Fig. 7).

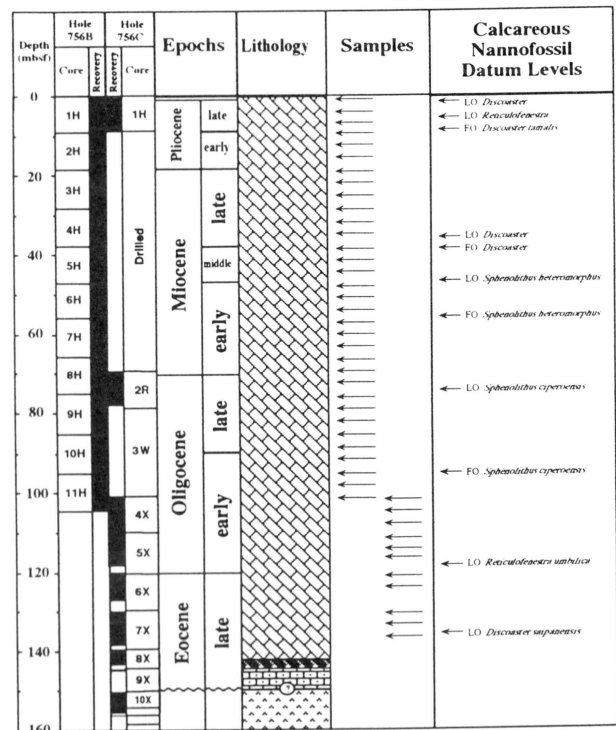


Fig. 6. Lithostratigraphy and Sample horizons at Site 756. See legend of Fig. 4.

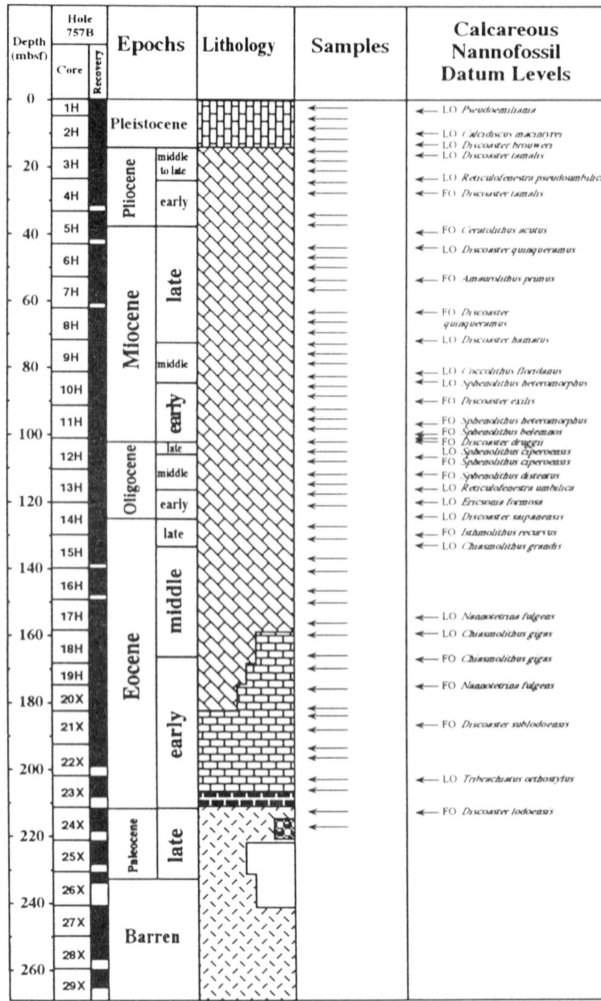


Fig. 7. Lithostratigraphy and Sample horizons at Site 757. See legend of Fig. 4.

Site 758 (5°23.049'S, 90°21.673' E) is located on the southeast side of Ninetyeast Ridge at a present water depth of 2922 m (Fig. 5), which is the deepest among Leg 121 sites. Hole 758A was cored with the APC and XCB systems to refusal at 422 mbsf, and with the RCB to a total depth of 677 mbsf. Average core recovery is 67.1%, but Cores 121-758A-1H through 121-758A-13X were completely recovered. Holocene to middle Miocene sediments are composed of nannofossil ooze with foraminifer, and the upper sediment units include terrigenous clay. Middle Miocene to Campanian sediments are mainly nannofossil and calcareous chalk, and occur across the unconformity that spans nearly the entire Eocene (Fig. 8). At Site 758, 58 sediment samples were analyzed (Samples 121-758A-1H-1 through 121-758A-31X-3, 0.75-292.65 mbsf) from the Pleistocene to late Eocene and Paleocene (Fig. 8).

Site 762 (19°53.23'S, 112°15.24' E) is located on the western part of the central Exmouth Plateau at a present water depth of 1371 m (Fig. 9). Hole 762C was drilled to a total depth of 940m by the XCB system, and recovered sediments from the Berriasian to Pleistocene (average recovery of 69.4%). The upper 182 m of foraminiferal-nannofossil and nannofossil ooze of late Oligocene and Neogene age is

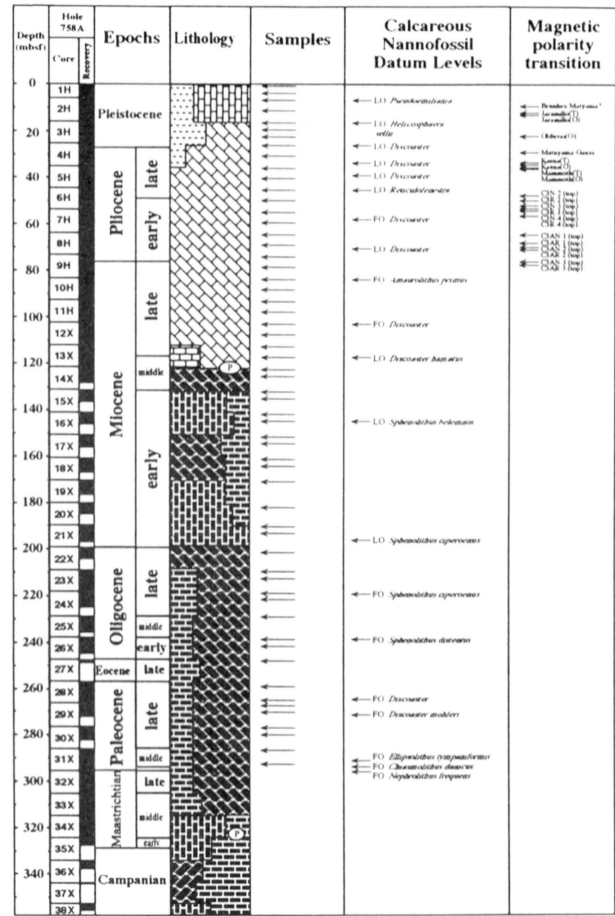


Fig. 8. Lithostratigraphy and Sample horizons at Site 758. See legend of Fig. 4.

underlain by nannofossil ooze and chalk of early Santonian to late Oligocene age (Fig. 10). At Site 762, 31 sediment samples were analyzed (Samples 122-762C-23X-1 through 122-762C-35X-1, 370.19-479.18 mbsf) from the early Eocene and late Paleocene (Fig. 10).

B. Methods

The sediment samples were processed by two methods. Loose sediments were washed with 63 μm sieve, whereas slightly consolidated sediments were treated with <3% hydrogen peroxide solution, and then washed on a 63 μm sieve (Nomura, 1991a; 1991b).

For stable isotope analysis, the benthic genus *Cibicidoides* has been measured in many oceanographic studies (Miller et al., 1989; Woodruff et al., 1990; Hodell et al., 1991; Barrera and Huber, 1991, amongst others). δ¹³C values of *Cibicidoides* are closely correlated with those of dissolved HCO₃⁻ in sea water at the sediment-water interface (Duplessy et al., 1984; Shackleton et al., 1984; Berger and Vincent, 1986; Savin and Woodruff, 1990). Because of microhabitat effects, infaunal taxa such as a *Oridorsalis* and

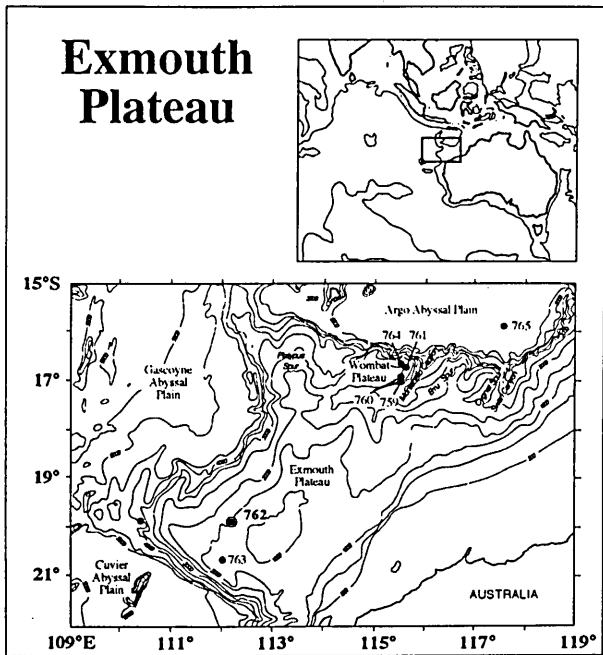


Fig. 9. Bathymetric map and seismic-reflection profile showing locations of ODP Leg 122 sites at Exmouth Plateau.

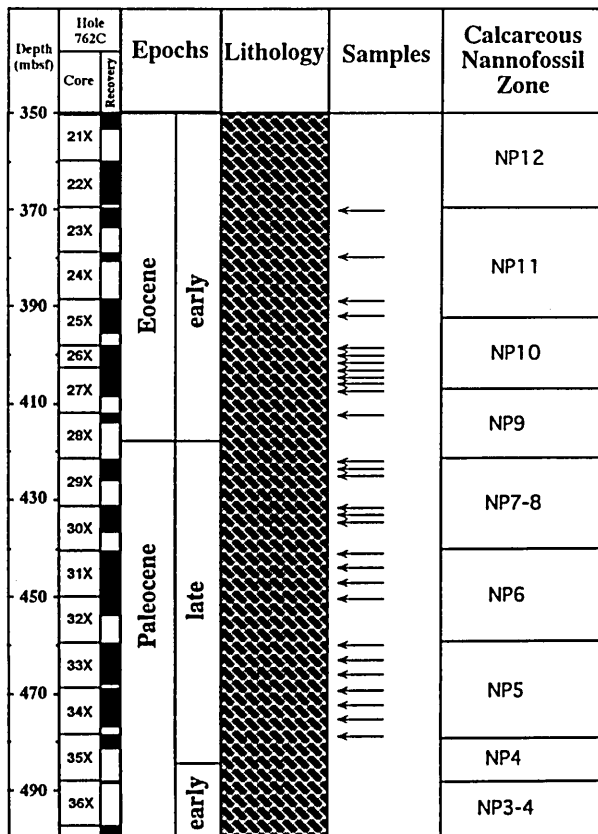


Fig. 10. Lithostratigraphy and Sample horizons at Site 762. See legend of Fig. 4.

Uvigerina are less well correlated with $\delta^{13}\text{C}$ values of dissolved HCO_3^- than epifaunal taxa such as a *Cibicidoides* (Belanger et al., 1981; Ganssen and Sarthein, 1983; Zahn et al., 1986; Savin and Woodruff, 1990; Woodruff and Savin, 1989). The species of *Cibicidoides*, however, are not common over a long stratigraphic range. For isotopic measurements, several species of the genus *Cibicidoides* were analyzed (e.g., Miller et al., 1989) or various species (e.g., Woodruff et al., 1990) at one examination. For example, the interspecific difference between *Cibicidoides kullenbergi* and *C. lamondohertyi* is $-0.22 \pm 0.12\%$ according to Woodruff et al. (1990). Therefore monospecies such as *Oridorsalis umbonatus* (which show consistent occurrence over a long range) are more useful for an isotopic study than the genus *Cibicidoides*. In this study, benthic foraminifer *Oridorsalis umbonatus* (Pleistocene to middle Eocene), *Anomalinoidea danicus* (middle Eocene to middle Paleocene), *Nuttallides truempyi* (early Eocene to late Paleocene), and *Stensioina beccariiiformis* (Paleocene to Maastrichtian) were mainly selected for stable isotopic analysis. They have been used in many other isotopic study by Vincent et al. (1985), Kennett and Stott (1990), Katz and Miller (1991), and others. To compare the isotopic values of these epifaunal species, *Cibicidoides wuellerstorfi*, *C. mundulus*, and *C. velascoensis* were measured. *Groidinoides soldanii* was also measured to correlate with data of Rea et al. (1991). Planktonic foraminifera *Subbotina* spp. (*S. triangularis*, *S. triloculinoides* and/or *S. linaperta*) were mainly used for stable isotope analysis from the Paleocene to Eocene. *Globorotalia pseudobulloides* and *Rugoglobigerna pennyi* were used for the Cretaceous / Tertiary boundary.

Benthic foraminifera for analysis were $>380 \mu\text{m}$ in this study, and three isotopic measurements were made on a single specimen in each sample. When there were few specimens $>380 \mu\text{m}$ in size, approximately 5-10 specimens in the 150-380 μm size fraction were analyzed. The differences between the averaged values of a single specimen per sample and values of 5-10 specimens were -0.041 ± 0.182 for $\delta^{18}\text{O}$ and -0.019 ± 0.129 for $\delta^{13}\text{C}$ ($N=27$) in *Stensioina beccariiiformis*, -0.115 ± 0.194 for $\delta^{18}\text{O}$ and 0.014 ± 0.143 ($N=22$) in *Nuttallides truempyi*. Approximately 5 specimens (180-350 μm size fraction) of the planktonic foraminifera *Subbotina* spp. and approximately 8 specimens (127-228 μm size fraction) of *Globorotalia pseudobulloides* and *Rugoglobigerna pennyi* were used for isotopic analysis. Other planktonic foraminifera were used according to various size fractions (Appendix A).

Specimens were put in a stainless-steel thimble and immersed in methyl alcohol. After the tests were disaggregated with a thin needle. They were ultrasonically cleaned. Stable isotope analyses were made using a Finnigan-MAT 250 mass spectrometer modified for ultra-small sample analysis, at the Shizuoka University. Oxygen and carbon isotope measurements were performed using the procedure of Wada et al. (1984; 1991). Carbonate specimens were reacted in saturated phosphoric acid, mixed solution of pyrophosphoric acid and few metaphosphoric acid (Wada et al., 1982), at 60.00°C . After the resulting CO_2 gas was purified in a glass-line, it was analyzed. The value thus obtained was converted into a value against a PDB standard by using NBS 20. The converted values are -4.18% for

$\delta^{18}\text{O}$ and -1.07‰ for $\delta^{13}\text{C}$, which is Craig's value of NBS 20 (in Blattner and Hulston, 1978). The precision for the isotope analysis is 0.02‰ for $\delta^{13}\text{C}$ and 0.05‰ for $\delta^{18}\text{O}$. Minimum volume of CO_2 gas for isotopic analysis is $2\mu\text{l}$.

C. Results

Results of isotopic analysis of benthic and planktonic foraminifera from Sites 752, 754, 756, 757, 758, and 762 are presented in Appendix A.

1. Site 752

At Site 752, oxygen and carbon isotopes were analyzed for five species; *Oridorsalis umbonatus*, *Anomalinoidea danicus*, *Nuttallides truempyi*, *Stensioina beccariiiformis*, and *Cibicidoides velascoensis*, and five planktonic foraminiferal species; *Acarinina primitiva*, *Morozovella marginodentata*, *Subbotina* spp., *Globorotalia pseudobulloides*, and *Rugoglobigerina penny*. However, analysis of *O. umbonatus* could only be made in Sample 121-752A-21X-1, 70-75cm (191.10 mbsf) from the late Paleocene, and values -0.221‰ (oxygen isotope) and 1.613‰ (carbon isotope).

Anomalinoidea danicus

Isotopes were analyzed for Samples 121-752A-13X-1, 70-75cm to 121-752A-32X-5, 70-75cm (113.60-295.40 mbsf) from the late Paleocene to early Eocene. Rare occurrences of *A. danicus* from Samples 121-752A-23X-1, 54-56cm to 121-752A-27X-3, 70-75cm (210.34-252.10 mbsf) from the upper Paleocene provide a few isotope data in this section. The isotopic records of *A. danicus* are plotted as a function of depth (mbsf) in Fig. 11.

Oxygen isotopes: Averaged $\delta^{18}\text{O}$ values of *A. danicus* between 295.40 and 252.1 mbsf gradually increase by 0.64‰ , from -1.011 to -0.390‰ . Between 239.77 and 192.60 mbsf in the upper Paleocene, $\delta^{18}\text{O}$ values are constant at about -0.6‰ . Near the Paleocene/Eocene boundary, two oxygen isotopic ratios of 191.1 mbsf and 181.4 mbsf exhibit relatively high values of -0.389 and -0.189‰ , respectively. $\delta^{18}\text{O}$ values from 174.85 to 113.60 mbsf in the lower Eocene show wide fluctuations, but gradually decrease by about 0.4‰ from -0.6 to -1.0‰ . During this period, the difference between the maximum (-0.561‰) and minimum (-1.420‰) values of $\delta^{18}\text{O}$ reaches 0.86‰ . $\delta^{18}\text{O}$ values at 152.40 mbsf and 123.30 mbsf are remarkably low (-1.420 and -1.371‰ , respectively).

Carbon isotopes: Averaged $\delta^{13}\text{C}$ values between 295.40 and 200.8 mbsf in the Paleocene gradually increase by 0.85‰ , from 1.433 to 2.315‰ . The $\delta^{13}\text{C}$ value at 200.8 mbsf is highest among the values of *A. danicus* at Site 752. $\delta^{13}\text{C}$ values at 289.40 and 252.10 mbsf are distinctively low. $\delta^{13}\text{C}$ values decrease by about 1.9‰ in the section between 200.80 and 155.40 mbsf, across the Paleocene / Eocene boundary. In the lower part of this section, $\delta^{13}\text{C}$ values dramatically decrease until a level slightly below the Paleocene / Eocene boundary. From 155.40 to 113.60 mbsf in the lower Eocene, $\delta^{13}\text{C}$ values are constant (about 0.2‰), with slight fluctuations value (max. 0.395 and min. -0.746‰). The pattern of fluctuation in $\delta^{13}\text{C}$ is similar to that of the

oxygen isotopic record. $\delta^{13}\text{C}$ values at 152.40 mbsf and 123.30 mbsf are extremely low (-0.746 and -0.648‰ , respectively), similar to $\delta^{18}\text{O}$ values.

Nuttallides truempyi

Two to five individuals were examined, because a single individual is too small ($153\text{--}380\mu\text{m}$ diameter) for isotopic analysis. However, measurements using a single individual were possible for the interval from 200.80 to 203.80 mbsf. However the difference in values between the measured individuals is small, and thus the results of both methods (using an individual and a few individuals) are almost the same. At this site, isotopic records were obtained from Samples 121-752A-13X-1, 70-75cm to 121-752A-26X-1, 97-100cm (113.60-239.77 mbsf) covering the upper Paleocene and lower Eocene. The isotopic records of *N. truempyi* are shown in Fig. 12.

Oxygen isotopes: *N. truempyi* $\delta^{18}\text{O}$ values decrease

Site 752 *Anomalinoidea danicus*

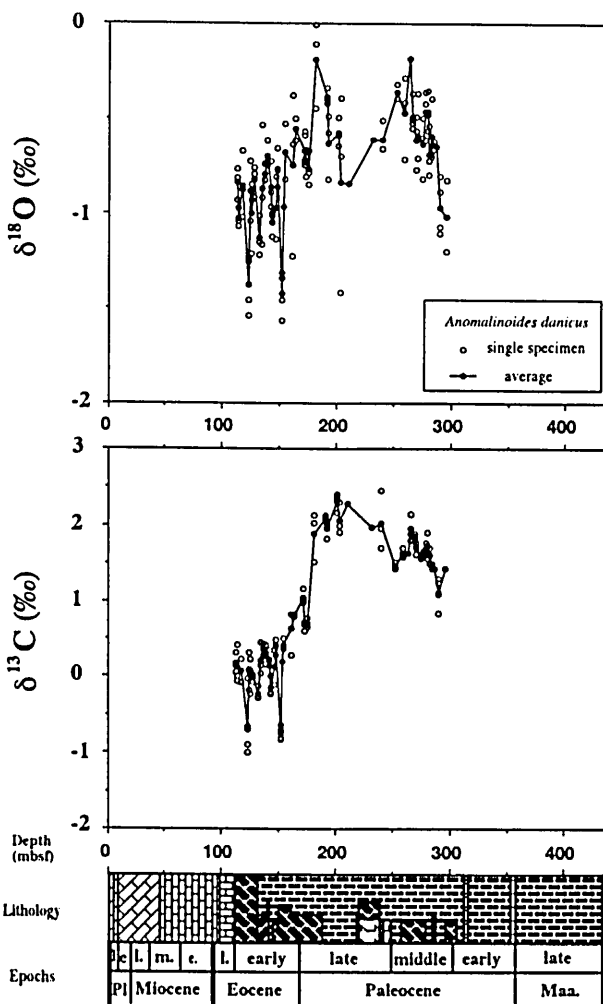


Fig. 11. Oxygen and carbon isotope records of benthic foraminifer *Anomalinoidea danicus* at Site 752.

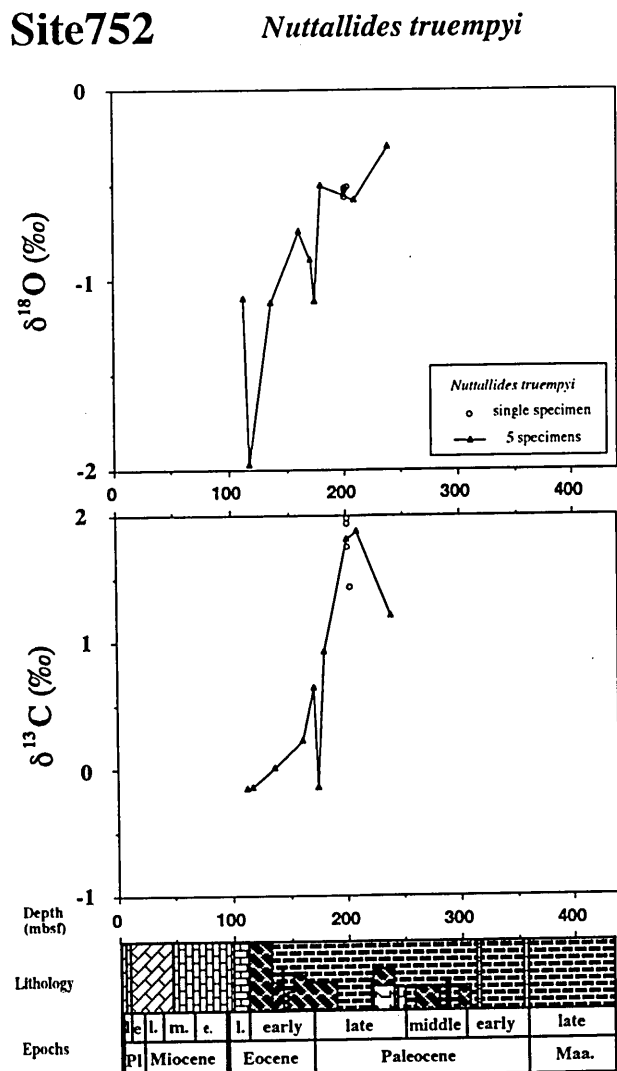


Fig. 12. Oxygen and carbon isotope records of benthic foraminifer *Nuttallides truempyi* at Site 752.

gradually by 0.2‰ from 239.77 to 181.40 mbsf in the upper Paleocene, and decrease more rapidly by 0.6‰ (-0.506 to -1.094‰) from 181.4 to 113.60 mbsf. $\delta^{18}\text{O}$ values of 174.85 mbsf (below the Paleocene / Eocene boundary) and 118.02 mbsf (above the P/E boundary) show remarkably little deviating from the general decreasing trend. The latter, in particular, records the lowest values (-1.971 ‰) among the *N. truempyi* data.

Carbon isotopes: $\delta^{13}\text{C}$ values increase from 239.77 to 210.34 mbsf in the upper Paleocene, peaking at 210.34 mbsf (1.873‰). From 210.34 to 113.60 mbsf across the Paleocene / Eocene boundary, $\delta^{13}\text{C}$ values decrease by about 2.0‰ from 1.873 to -0.149‰. $\delta^{13}\text{C}$ values decrease rapidly below the Paleocene / Eocene boundary (200.80 to 174.85 mbsf). $\delta^{13}\text{C}$ values at 174.85 mbsf are remarkably low (-0.143‰), and related to the "benthic extinction event".

Stensioina beccariiformis

Isotope analyses were performed for material from Samples 121-752A-16X-4, 59-65cm to 121-752A-33X-3,

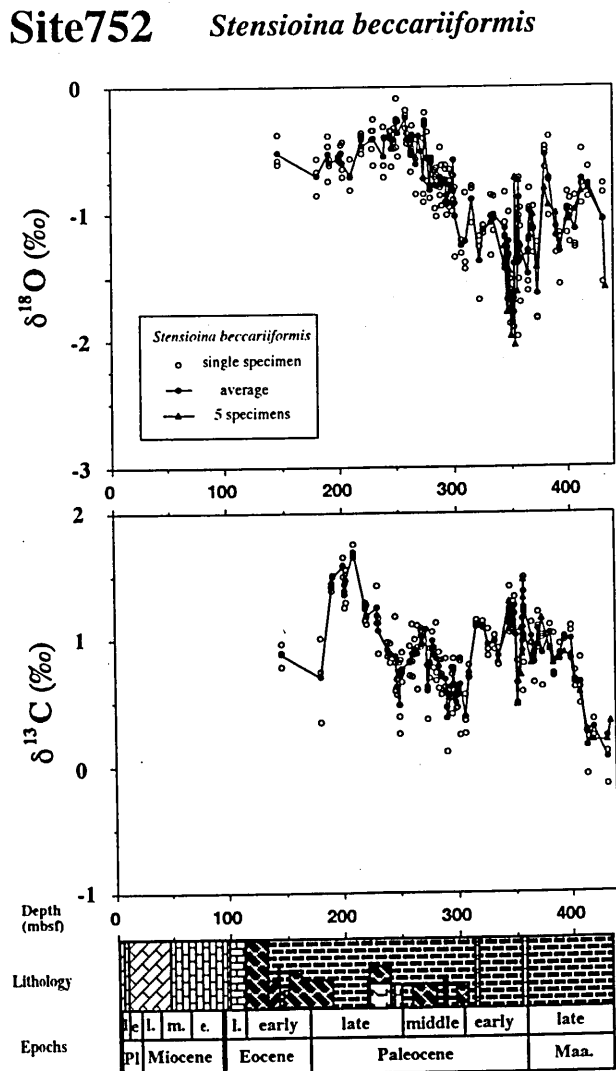


Fig. 13. Oxygen and carbon isotope records of benthic foraminifer *Stensioina beccariiformis* at Site 752.

57-60cm (147.09-301.97 mbsf) and from Samples 121-752B-5R-3, 50-53cm to 121-752B-19R-3, 49-52cm (300.50-435.09 mbsf) from the upper Maastrichtian to lower Eocene. The isotopes of *S. beccariiformis* below 335.82 mbsf were analyzed using five to ten individuals, or a single individual. No occurrence of *S. beccariiformis* slightly above Sample 121-752A-20X-1, 70-75cm (181.40 mbsf) before the Paleocene / Eocene boundary indicates that this species became extinct by the late Paleocene Benthic Event (e.g., Nomura, 1991a). However, reworked *S. beccariiformis* was found in Sample 121-752A-16X-4, 59-65cm (147.09 mbsf) above the event. The oxygen and carbon isotopic values of this sample are similar to those of 181.40 mbsf located before the "extinction event", and are high in comparison with those of *A. danicus* from the same sample. Therefore, *S. beccariiformis* at 174.85 mbsf supports the intercalation of reworked sediment. The isotopic records of *S. beccariiformis* are shown in Fig. 13.

Oxygen isotopes: Fluctuations in $\delta^{18}\text{O}$ with a general decreasing trend of 0.3‰, from -0.8 to -1.1‰, is shown

from 435.09 to 326.85 mbsf (late Maastrichtian to early Paleocene). Above the Cretaceous / Tertiary boundary, $\delta^{18}\text{O}$ values show the largest amplitude from a maximum of -0.746‰ at 358.18 mbsf, to a minimum of -1.969‰ at 351.60 mbsf. From 326.85 to 252.10 mbsf in the Paleocene section, $\delta^{18}\text{O}$ values increase from -1.120 to -0.244‰ . From a maximum value of 258.80 mbsf, $\delta^{18}\text{O}$ values gradually decrease to -0.699‰ at 181.40 mbsf, located below the Paleocene / Eocene boundary near the extinction event.

Carbon isotopes: Four $\delta^{13}\text{C}$ values between 435.09 and 413.63 mbsf in the upper Maastrichtian are constant at $\sim 0.2\text{‰}$, showing the lowest $\delta^{13}\text{C}$ values at Site 752. $\delta^{13}\text{C}$ values rapidly increase by $0.6\text{--}0.8\text{‰}$ up to 400.42 mbsf in the upper Maastrichtian. The section across the Cretaceous / Tertiary boundary (400.42 to 345.6 mbsf) shows a slight increase with a small oscillation of carbon isotopic records. However, the values just above the Cretaceous / Tertiary boundary show a large variation from 1.486‰ (358.18 mbsf) to 0.475‰ (353.10 mbsf). Averaged $\delta^{13}\text{C}$ values tend to decrease by 0.75‰ from 345.6 to 307.04 mbsf in the lower Paleocene. In this section, averaged $\delta^{13}\text{C}$ values increase from 335.82 mbsf to 316.92 mbsf, and rapidly decrease to 307.04 mbsf (0.392‰), which is the lowest value among the Paleocene $\delta^{13}\text{C}$ values. In the late early and late Paleocene, averaged $\delta^{13}\text{C}$ values tend to increase by about 1.3‰ , up to a peak value of 1.689‰ at 210.34 mbsf. However, averaged $\delta^{13}\text{C}$ values decrease between 272.92 and 247.12 mbsf in the late Paleocene. A rapid decrease of $\delta^{13}\text{C}$ value is shown from 210.34 to 181.40 mbsf, where *S. beccariiiformis* became extinct below the Paleocene / Eocene boundary.

Cibicoides velascoensis

Isotopic records were obtained from Samples 121-752A-20X-1, 70-75cm to 121-752A-33X-3, 57-60cm (181.40-301.99 mbsf) and from Samples 121-752B-5R-3, 50-53cm to 121-752B-19R-3, 49-52cm (300.50-435.09 mbsf) from the lower Maastrichtian and Paleocene. No occurrence of *C. velascoensis* above Sample 121-752A-20X-1, 70-75cm (181.40 mbsf) before the Paleocene / Eocene boundary may result from the late Paleocene Benthic Event. The rare occurrence of *C. velascoensis* through Samples 121-752A-25X-3, 79-84cm to 121-752A-30X-1, 73-76cm (232.89-278.13 mbsf) in the Paleocene provides some isotope data in this section. The isotopic record of *C. velascoensis* is shown in Fig. 14.

Oxygen isotopes: $\delta^{18}\text{O}$ values repeatedly fluctuate from 435.09 to 310.59 mbsf in the upper Maastrichtian and lower Paleocene. In this section, $\delta^{18}\text{O}$ values increase by $\sim 0.2\text{‰}$ from 435.09 to 358.18 mbsf, shift down about 0.4‰ around 360 mbsf immediately above the Cretaceous / Tertiary boundary, then show a general increasing trend of $\sim 0.1\text{‰}$ up to 310.59 mbsf. From 310.59 to 284.65 mbsf in the late early Paleocene, $\delta^{18}\text{O}$ values rapidly increase by $\sim 0.7\text{‰}$ from -1.044 to -0.345‰ . In the late Paleocene, $\delta^{18}\text{O}$ values are relatively constant ($\sim -0.4\text{‰}$).

Carbon isotopes: $\delta^{13}\text{C}$ values slightly decrease from 435.09 to 413.63 mbsf in the upper Maastrichtian, then show the lowest value among Site 752. From 413.63 to 400.32 mbsf in upper Maastrichtian, $\delta^{13}\text{C}$ values increase from 0.531 to 1.289‰ . $\delta^{13}\text{C}$ values gradually decrease up to 367.50 mbsf.

In the section from 367.50 to 347.60 mbsf just across the Cretaceous / Tertiary boundary, $\delta^{13}\text{C}$ values show a large fluctuation with a maximum of 1.719‰ (358.44 mbsf) and a minimum of 0.880‰ (353.10 mbsf). In the interval from 347.60 to 307.04 mbsf, $\delta^{13}\text{C}$ values increase by 0.3‰ between 326.85 and 322.86 mbsf; however the values decrease by 0.45‰ between 316.92 and 310.59 mbsf. $\delta^{13}\text{C}$ value at 307.04 mbsf is the lowest among the early Paleocene. In the late early and late Paleocene, $\delta^{13}\text{C}$ values tend to increase by 1.5‰ up to a peak value of 2.375‰ at 210.34 mbsf. In this section, $\delta^{13}\text{C}$ values of 245.77 mbsf are relatively low. In the latest late Paleocene, $\delta^{13}\text{C}$ values decrease rapidly by 0.9‰ from 210.34 to 181.40 mbsf and then *C. velascoensis* disappears in the section.

Subbotina spp.

Isotopic analysis is analyzed for Samples 121-752A-17X-3, 70-75cm to 121-752A-33X-3, 57-60cm (155.40-

Site 752 *Cibicoides velascoensis*

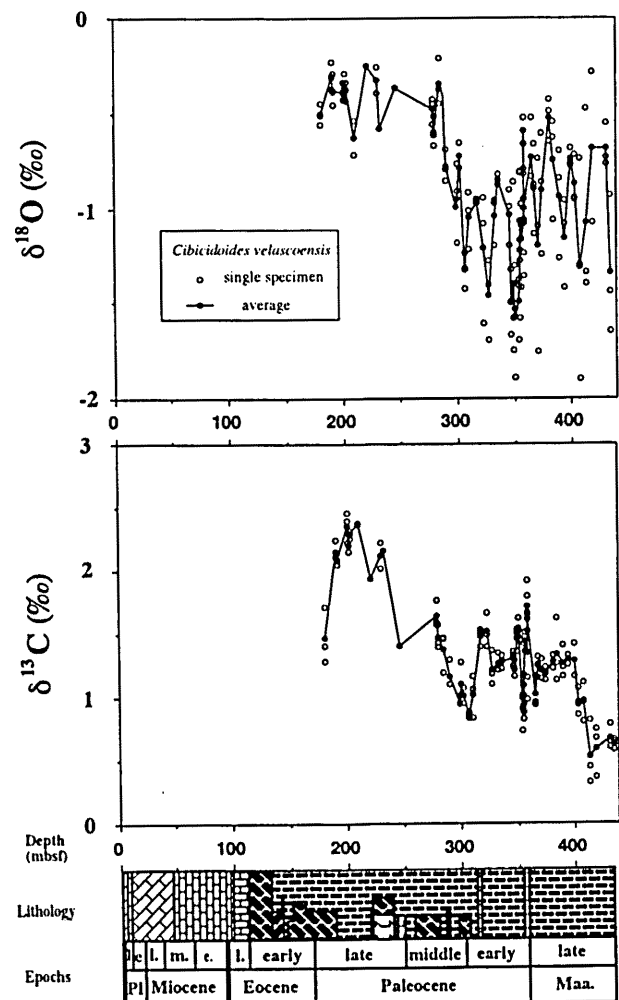


Fig. 14. Oxygen and carbon isotope records of benthic foraminifer *Cibicoides velascoensis* at Site 752.

301.97 mbsf) were made, from the lower and upper Paleocene. The isotopic records of *Subbotina* spp. are shown in Fig 15..

Oxygen isotopes: $\delta^{18}\text{O}$ values increase by about 0.6‰ from 301.97 to 258.80 mbsf. In this section, $\delta^{18}\text{O}$ values show a distinct decrease just above lower / upper Paleocene boundary with a minimum value of -1.870‰ at 274.40 mbsf. From 258.8 to 155.40 mbsf (upper Paleocene and lowermost early Eocene), $\delta^{18}\text{O}$ values fluctuate although generally decrease. $\delta^{18}\text{O}$ values increase from -1.873 to -1.324‰ in the section between 174.85 and 163.75 mbsf just across the Paleocene / Eocene boundary.

Carbon isotopes: $\delta^{13}\text{C}$ values from 301.97 to 220.20 mbsf in the lower Paleocene increase by about 1.3‰ (1.640 to 2.917‰), then decrease by 0.55‰ between 268.40 and 249.10 mbsf in the upper Paleocene section. From 220.20 to 155.40 mbsf (upper Paleocene and lowermost Eocene), $\delta^{13}\text{C}$

values decrease by about 2.2‰. The dramatic decrease in this section occurs slightly below the Paleocene / Eocene boundary.

Other planktonic foraminifera

Less common planktonic foraminiferal species, *Acarinina primitiva*, *Morozovella marginodentata*, *Globorotalia pseudobulloides*, and *Rugoglobigerina pennyi*, were analyzed for Samples 121-752A-13X-1, 70-75cm to 121-752A-17X-3, 70-75cm (113.60-155.40 mbsf, lower Eocene), Samples 121-752A-15X-1, 70-75cm to 121-752A-20X-3, 70-75cm (133.00-181.40 mbsf, lower Eocene to the uppermost Paleocene), Samples 121-752B-10R-2, 100-102cm to 121-752B-10R-7, 41-43cm (347.60-354.51 mbsf, above the Cretaceous / Tertiary boundary), and Samples 121-752B-11R-1, 112-114cm to 121-752B-12R-5, 54-57cm (358.92-370.94 mbsf, below the Cretaceous / Tertiary boundary), respectively. Two to eight individuals analyzed at

Site752

Subbotina spp.

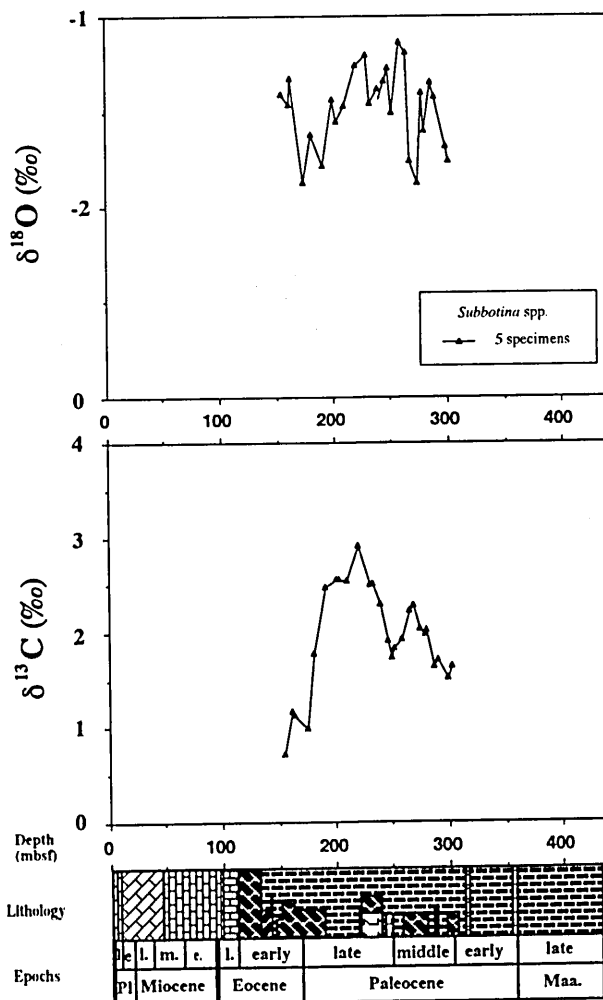


Fig. 15. Oxygen and carbon isotope records of planktonic foraminifer *Subbotina* spp. at Site 752.

Site752

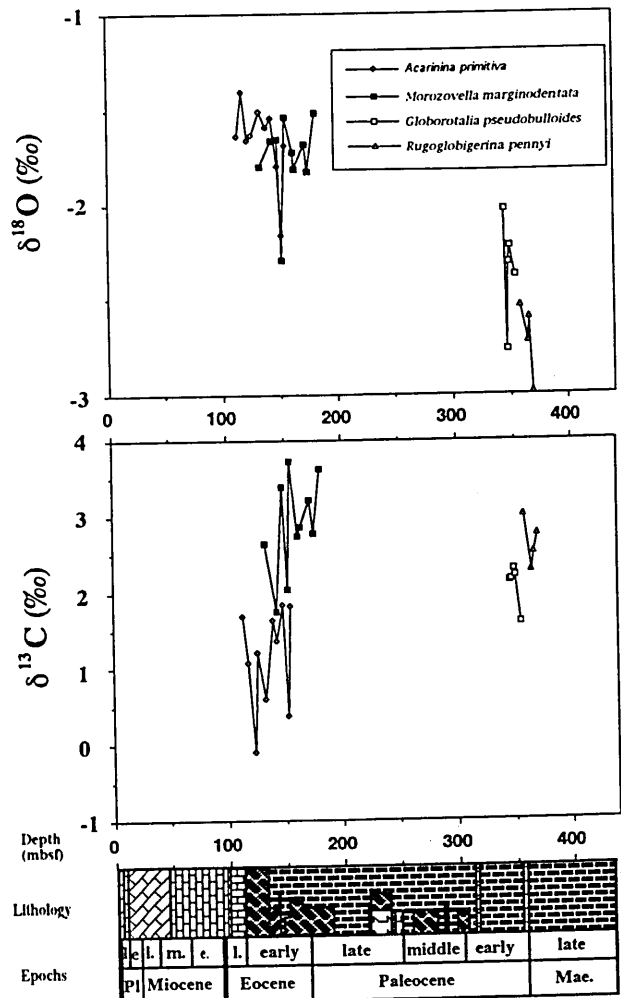


Fig. 16. Oxygen and carbon isotope records of other planktonic foraminifers at Site 752.

Site752

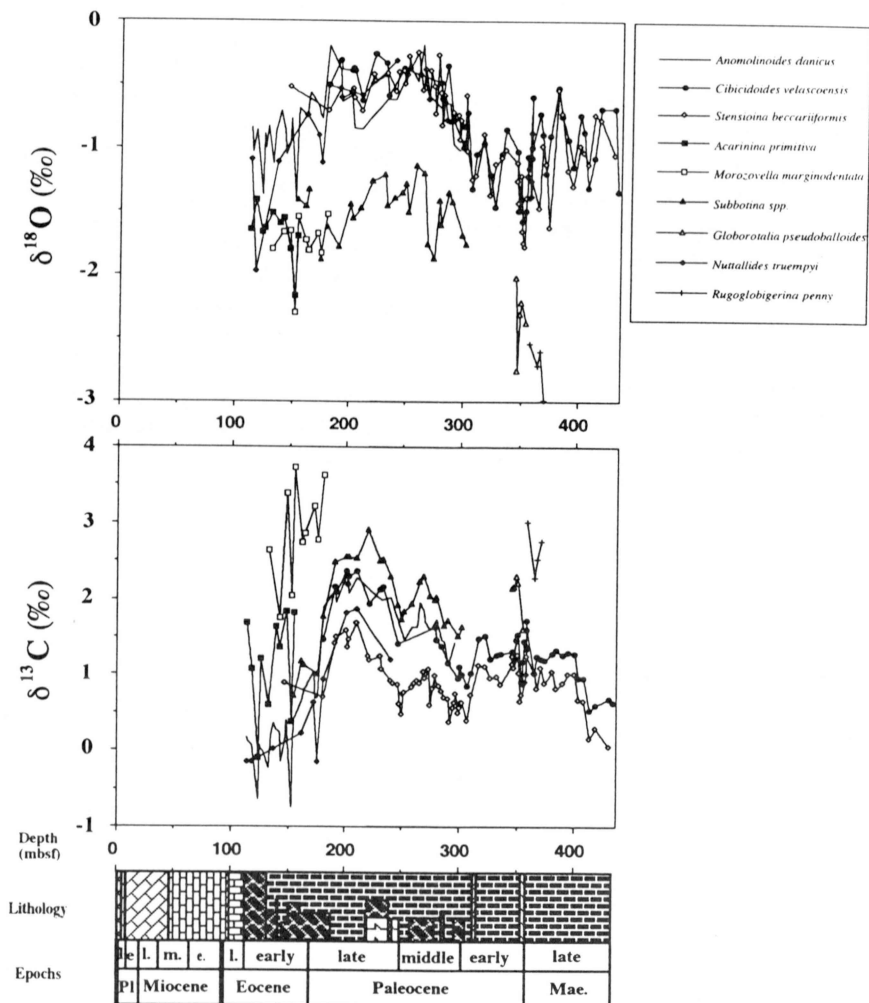


Fig. 17. Summary of oxygen and carbon isotope records at Site 752.

Site 752. *Globorotalia pseudobulloides* and *R. penny* do not overlap. The isotopic records of those species are shown in Fig. 16.

Oxygen isotopes: $\delta^{18}\text{O}$ values of *R. penny* below the Cretaceous / Tertiary boundary increase from -2.988 to -2.538‰. $\delta^{18}\text{O}$ values of *G. pseudobulloides* above the Cretaceous / Tertiary boundary also show a general increase, from -2.377 to -2.020‰. However, $\delta^{18}\text{O}$ of *G. pseudobulloides* shows remarkably low values at 347.60 mbsf. In the latest late Paleocene and early Eocene, $\delta^{18}\text{O}$ values of *M. marginodentata* are constant ($\sim -1.7\text{‰}$). In the early Eocene, $\delta^{18}\text{O}$ values of *A. primitiva* tend to increase. No significant change of *M. marginodentata* is observed across the Paleocene / Eocene boundary, which differs from the oxygen isotopic records of *Subbotina* spp. In *A. primitiva* and *M. marginodentata*, $\delta^{18}\text{O}$ values at 152.4 mbsf exhibit remarkably low values of -2.160 and -2.291‰, respectively.

Carbon isotopes: $\delta^{13}\text{C}$ values of *R. penny* below the Cretaceous / Tertiary boundary are constant at $\sim 2.6\text{‰}$. Above

the Cretaceous / Tertiary boundary, $\delta^{13}\text{C}$ values of *G. pseudobulloides* are about 2.2‰, except for a slightly low value at 354.51 mbsf. In *M. marginodentata*, $\delta^{18}\text{O}$ values are constant at $\sim 3\text{‰}$ from 181.4 to 155.4 mbsf across the Paleocene / Eocene boundary, and tend to decrease from 155.40 to 133.00 mbsf in the lower Eocene. $\delta^{13}\text{C}$ values of *A. primitiva* in the lower Eocene are constant at $\sim 1\text{‰}$. However, $\delta^{13}\text{C}$ values of *M. marginodentata* and *A. primitiva* fluctuate. At 152.40 mbsf and 123.30 mbsf, $\delta^{13}\text{C}$ values are remarkably low.

Isotopic values obtained from Site 752 are summarized in Fig. 17.

2. Site 754

At Site 754, only the benthic foraminifer *Oridorsalis umbonatus* was analyzed for oxygen and carbon isotopes. Isotope analyses were made on material from Samples 121-754A-1H-1, 70-75cm to 121-754A-13X-3, 70-75cm (0.70-

116.00 mbsf) from the Pleistocene to late Oligocene. The isotopic records of *Oridorsalis umbonatus* at Site 754 are shown in Fig. 18.

Oxygen isotopes: The oxygen isotopic records of *Oridorsalis umbonatus* are constant at $\sim 1.2\text{‰}$ from 116.00 mbsf to 89.90 mbsf in the upper Oligocene and the lower Miocene. $\delta^{18}\text{O}$ values decrease by $\sim 0.4\text{‰}$ across the early / middle Miocene boundary. $\delta^{18}\text{O}$ values are relatively low from 83.90 to 77.20 mbsf (nannofossil Zones CN3-4). In this interval, $\delta^{18}\text{O}$ values at 83.90 mbsf display the lowest value (0.933‰) among the Neogene samples. In the middle Miocene, $\delta^{18}\text{O}$ values from 77.20 to 60.80 mbsf increase significantly by 0.85‰ from 1.077‰ to 1.924‰ . $\delta^{18}\text{O}$ values are constant about 2.0‰ from 60.80 to 32.00 mbsf through the upper

middle and upper Miocene. In the early Pliocene, $\delta^{18}\text{O}$ values increase by 0.5‰ up to a small peak at 22.40 mbsf (2.467‰), and decrease by $\sim 0.3\text{‰}$. A shift of $\delta^{18}\text{O}$ (of $\sim 0.4\text{‰}$) is observed across the early / late Pliocene boundary. During the late Pliocene and Pleistocene, $\delta^{18}\text{O}$ values increase up to 2.945‰ at 0.70 mbsf near to the present.

Carbon isotopes: The carbon isotopic records show a notable increase from 113.00 to 103.30 mbsf across the Oligocene / Miocene boundary, and then attain the highest (0.938‰) among lower Miocene samples. The magnitude of this increase is 0.9‰ . $\delta^{13}\text{C}$ values decrease by 0.5‰ up to 93.60 mbsf in nannofossil Zone CN1, then begin to increase again. This increase is gradual up to 83.90 mbsf, then changes up to 80.20 mbsf within Zones CN3-4. From 80.20 to 74.20 mbsf across the Zone CN4/5 boundary in the middle Miocene, $\delta^{13}\text{C}$ values are distinctly high ($\sim 1.2\text{‰}$). During the middle Miocene, $\delta^{13}\text{C}$ values drastically decrease by about 1.1‰ from 1.2 to 0.1‰ at 67.50 mbsf, and then rapidly increase by 0.65‰ up to a weak peak at 57.80 mbsf (0.749‰). $\delta^{13}\text{C}$ values are constant at $\sim 0.5\text{‰}$ from 54.80 to 35.60 mbsf (middle and upper Miocene). $\delta^{13}\text{C}$ values display a distinct shift to lower values (decreasing by $\sim 0.5\text{‰}$) between 35.60 and 32.00 mbsf (uppermost Miocene), and show a decreasing trend up to minimum value at 22.40 mbsf (-0.341‰) in the early Pliocene (Zones CN11a-10b). This shift is correlated with the Chron-6 Carbon Shift in the latest Miocene (Vincent et al., 1980; 1985). The carbon isotopic records exhibit a slight increase up to a weak peak at 6.80 mbsf (Zone CN12d) just below the Pleistocene / Pliocene boundary, and then show a decreasing trend up to 0.70 mbsf near to the present.

Site 754 *Oridorsalis umbonatus*

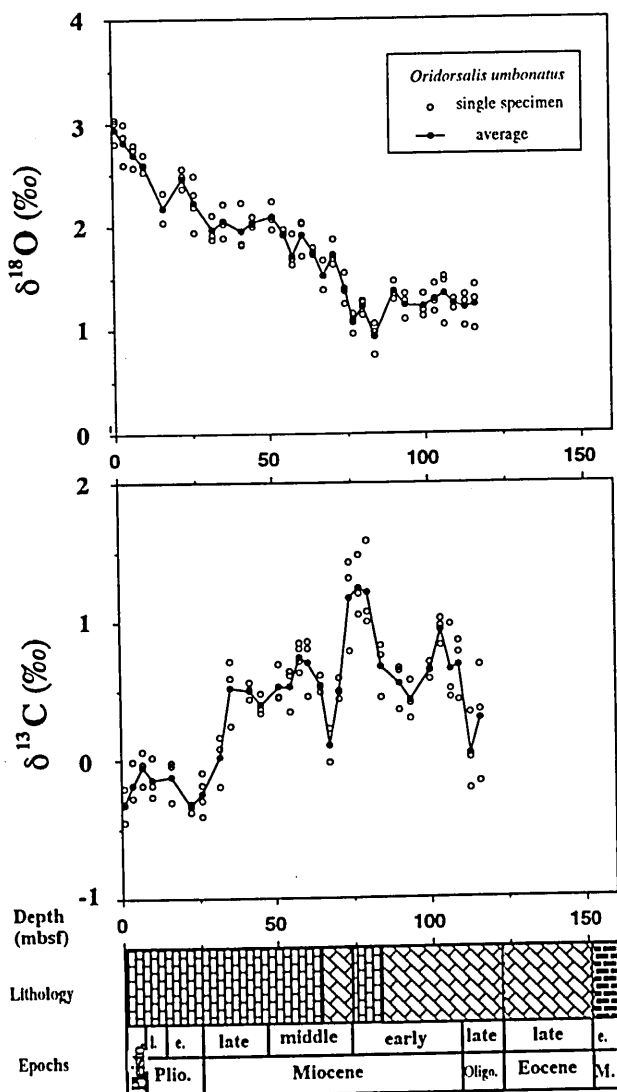


Fig. 18. Oxygen and carbon isotope records of benthic foraminifer *Oridorsalis umbonatus* at Site 754.

3. Site 756

At Site 756, only the benthic foraminifer *Oridorsalis umbonatus* was analyzed for oxygen and carbon isotopes. The analyzed material was obtained from Samples 121-756A-1H-1, 70-75cm to 121-756A-11H-5, 70-75cm (0.70-101.30 mbsf), and from Samples 121-756A-4X-1, 70-75cm to 121-756A-7X-5, 70-75cm (101.60-136.50 mbsf), from the Pleistocene to the uppermost Eocene. Fig. 19 shows the oxygen and carbon isotopic records for *Oridorsalis umbonatus* at Site 756.

Oxygen isotopes: The oxygen isotopic records show an increasing trend up to 120.90 mbsf within nannofossil Zone CP16 in the earliest Oligocene. The magnitude of this increase is 0.95‰ . During the late early Oligocene to early Miocene, $\delta^{18}\text{O}$ values show a slight increase with fluctuations of amplitude of $\sim 0.5\text{‰}$. In this interval, weak peaks in $\delta^{18}\text{O}$ value are observed at 101.60 mbsf (Zones CP17-18), 81.90 mbsf (Zone CP19), and 56.90 mbsf (Zones CN1-2) (1.622 , 1.715 , and 1.643‰ , respectively). The value of 53.60 mbsf (Zone CN3-4 just below the early / middle Miocene boundary) is the lowest among the Neogene samples. During the middle Miocene, $\delta^{18}\text{O}$ values rapidly increase by about 1.4‰ up to 2.590‰ at 41.0 mbsf (Zones CN5-6). A constant value of $\sim 2.5\text{‰}$ is observed from 41.00 to 9.20 mbsf (middle Miocene to lower Pliocene). Across the late / early Pliocene boundary, $\delta^{18}\text{O}$ values rapidly increase by about 0.8‰ from 2.335‰ (9.20 mbsf) to 3.169‰ (3.70

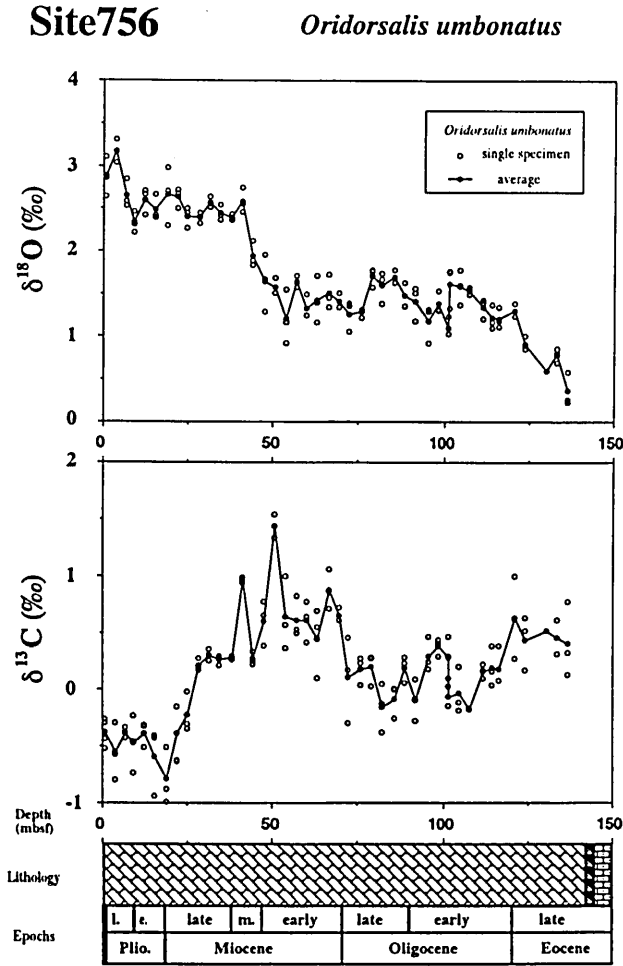


Fig. 19. Oxygen and carbon isotope records of benthic foraminifer *Oridorsalis umbonatus* at Site 756.

mbsf). Limited data from the late Pliocene and Pleistocene shows ~3.0‰.

Carbon isotopes: The carbon isotopic record shows a gradual increase up to 0.636‰ at 120.90 mbsf (Zones CP16 in the early Oligocene). In the lower Oligocene, δ¹³C values from 120.90 to 107.60 mbsf decrease by 0.8‰. δ¹³C values increase from -0.173‰ (107.60 mbsf) to 0.383‰ (98.30 mbsf) in the late early Oligocene and then reverse at the early / late Oligocene boundary, decreasing to -0.148‰ at 81.90 mbsf (Zone CP19). Across the Oligocene / Miocene boundary, δ¹³C values rapidly increase by 1.0‰ up to the peak at 66.30 mbsf (Zones CN2-1). From 62.90 to 50.60 mbsf, δ¹³C values increase by ~1.0‰ from 0.449 to 1.439‰. In this section, a notable shift is observed between 53.60 and 50.60 mbsf. During the middle Miocene, δ¹³C values significantly decrease by ~1.2‰ up to 44.00 mbsf in Zones CN5-6, and then rapidly increase by 0.7‰ up to a distinct peak at 41.00 mbsf (0.967‰), and again decrease up to 38.0 mbsf in Zone CN7. From 38.00 to 28.40 mbsf, δ¹³C values are constant (~0.25‰). In the uppermost Miocene, the carbon isotopic records between 28.40 and 18.80 mbsf

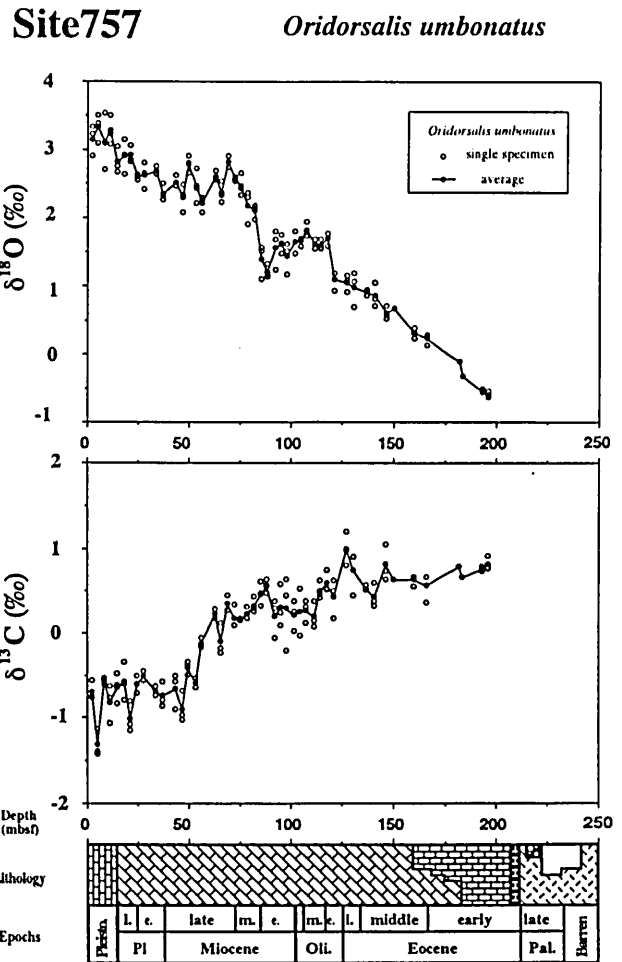


Fig. 20. Oxygen and carbon isotope records of benthic foraminifer *Oridorsalis umbonatus* at Site 757.

exhibit a remarkable negative shift that correlates with the Chron-6 Carbon Shift in the latest Miocene (Vincent et al., 1980; 1985). During the Pliocene and Pleistocene, δ¹³C increases to 0.4‰ up to 12.2 mbsf. Those values are constant at ~-0.4‰.

4. Site 757

At Site 757, oxygen and carbon isotopes were analyzed for 2 species of benthic foraminifers; *Oridorsalis umbonatus* and *Anomalinoidea danicus*, and for planktonic foraminifera *Subbotina* spp. The isotopic values *A. danicus* overlap with that of *Oridorsalis umbonatus* from Samples 121-757B-15H-5, 70-75cm to 121-757B-22X-3, 70-75cm (136.5-196.10 mbsf), and with that of *Subbotina* spp. from Samples 121-757B-15H-5, 70-75cm to 121-757B-24X-4, 48-61cm (136.5-216.78 mbsf).

Oridorsalis umbonatus

Isotopes were analyzed for Samples 121-757B-1H-2, 70-75cm to 121-757B-22X-3, 70-75cm (2.03-196.1 mbsf) from the lower Eocene to Pleistocene. Insufficient isotopic

data were obtained below the middle Eocene, because *Oridorsalis umbonatus* are rare. The isotopic records of *Oridorsalis umbonatus* are shown in Fig 20.

Oxygen isotopes: $\delta^{18}\text{O}$ values increase gradually through the Eocene, and up to 120.80 mbsf within the nannofossil Zones CP16a-b in the lowermost Oligocene. In this section, they increase up to 1.6‰. Furthermore, $\delta^{18}\text{O}$ values show a remarkable positive shift of 0.6‰ (from 1.104 to 1.710) between 120.80 and 117.20 mbsf across the Zone CP16b/c boundary. From the early Oligocene to early Miocene, $\delta^{18}\text{O}$ values decrease with small fluctuations. In this section, weak peaks are observed at 107.50 and 94.90 mbsf. In the early middle Miocene, in Zones CN3-4, $\delta^{18}\text{O}$ values decrease to the minimum value of 1.217‰ at 88.29 mbsf throughout the Neogene. From 88.20 to 68.90 mbsf, $\delta^{18}\text{O}$ values rapidly increase by ~ 1.6 ‰ from 1.217 to 2.825‰ in the middle Miocene. From the late Miocene to early Pliocene, $\delta^{18}\text{O}$ values are constant at around 2.6‰, despite a relatively large fluctuation (~ 0.6 ‰ amplitude). In this section, the peak $\delta^{18}\text{O}$ value of 2.793‰ is observed at 49.50 mbsf within Zone CN9b. In the late Pliocene and Pleistocene, $\delta^{18}\text{O}$ values increase up to ~ 3.3 ‰ near to the present. In this section, two small positive shifts are recognized; a shift of 0.30‰ between 24.30 and 20.70 mbsf across the Zone CN11b/12a boundary, and a shift of 0.45‰ between 11.20 and 14.70 mbsf across the CN12/13 boundary.

Carbon isotopes: The carbon isotopic record in the Eocene shows a trend of slight decrease up to 0.436‰ at 140.20 mbsf, then an increase up to a peak value at 126.80 mbsf (0.991‰) just below the Eocene / Oligocene boundary. The magnitude of this increase is about 0.6‰. From 126.80 to 112.20 mbsf in the lower Oligocene, $\delta^{13}\text{C}$ values decrease by ~ 0.8 ‰. From the late early Oligocene to early Miocene, $\delta^{13}\text{C}$ values are rather stable at 0.25‰. Above the early / middle Miocene boundary, a small positive shift of 0.35‰ is recognized between 91.90 and 88.20 mbsf, and then $\delta^{13}\text{C}$ values decrease to a maximum value of 0.56‰ throughout the Neogene. During the middle Miocene, $\delta^{13}\text{C}$ values gradually decrease by 0.4‰. Within Zones CN9a-8 in the late Miocene, $\delta^{13}\text{C}$ values increase up to slight peak at 68.90 mbsf (0.375‰), then immediately decrease to a lower value at 65.90 mbsf (-0.094‰), and again increase up to 62.90 mbsf. The notable decrease of $\delta^{13}\text{C}$ value from 62.90 to 46.50 mbsf in the uppermost Miocene reaches ~ 1.1 ‰, from 0.233 to -0.892‰. This decrease is correlated with the Chron-6 Carbon Shift in the latest Miocene (Vincent et al., 1980; 1985). During the Pliocene and Pleistocene, $\delta^{13}\text{C}$ values are constant at ~ 0.7 ‰. However, lower values are observed at 20.70 mbsf (-1.005‰) and 5.20 mbsf (-1.310‰).

Anomalinoidea danicus

Isotope analyses were performed on material from Samples 121-757B-15H-5, 70-75cm to 121-757B-24X-4, 58-61cm (136.50-216.78 mbsf) in lower and middle Eocene. The isotopic records of *A. danicus* are shown in Fig. 21.

Oxygen isotopes: The oxygen isotopic record displays a gradual increase up to 0.420‰ at 159.50 mbsf (nannofossil Zone CP13b). The magnitude of this increase reaches 1.74‰. In Zones CP13c and 14 in the upper middle Eocene, $\delta^{18}\text{O}$ values are constant at ~ 0.4 ‰.

Carbon isotopes: $\delta^{13}\text{C}$ values notably increase by about 1.0‰ from 0.302‰ (212.28 mbsf) to 1.340‰ (183.90 mbsf) through Zones CP10 and 11 in the lower Eocene. In the middle Eocene, $\delta^{13}\text{C}$ values exhibit a decrease of 0.4‰ up to 169.20 mbsf (0.934‰) within Zone CP13a, an increase of 0.3‰ up to a peak value at 146.20 mbsf (Zones CP14), and a decrease of 0.35‰.

Subbotina spp.

Isotope analyses were performed on five small sized tests. The oxygen and carbon isotopic values were obtained from Samples 121-757B-14H-1, 70-75cm to 121-757B-24X-4, 58-61cm (120.80-216.78 mbsf) in middle and lower Eocene. The isotopic records of *Subbotina* spp. are shown in Fig. 22.

Oxygen isotopes: $\delta^{18}\text{O}$ values increase gradually throughout the Eocene, except for a lower value at 193.10 mbsf (within nannofossil Zone CP11). The magnitude of this increase reaches 2.35‰ (from -1.323 to 1.012‰).

Carbon isotopes: $\delta^{13}\text{C}$ values increase about 0.9‰ from 212.28 to 202.70 mbsf across the Zone 10/11 boundary in the

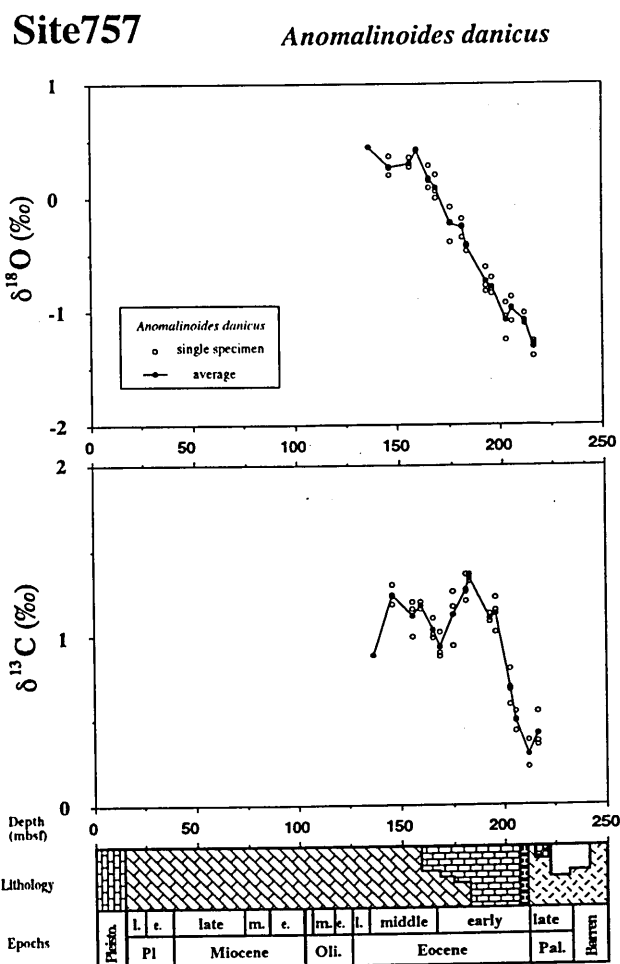


Fig. 21. Oxygen and carbon isotope records of benthic foraminifer *Anomalinoidea danicus* at Site 757.

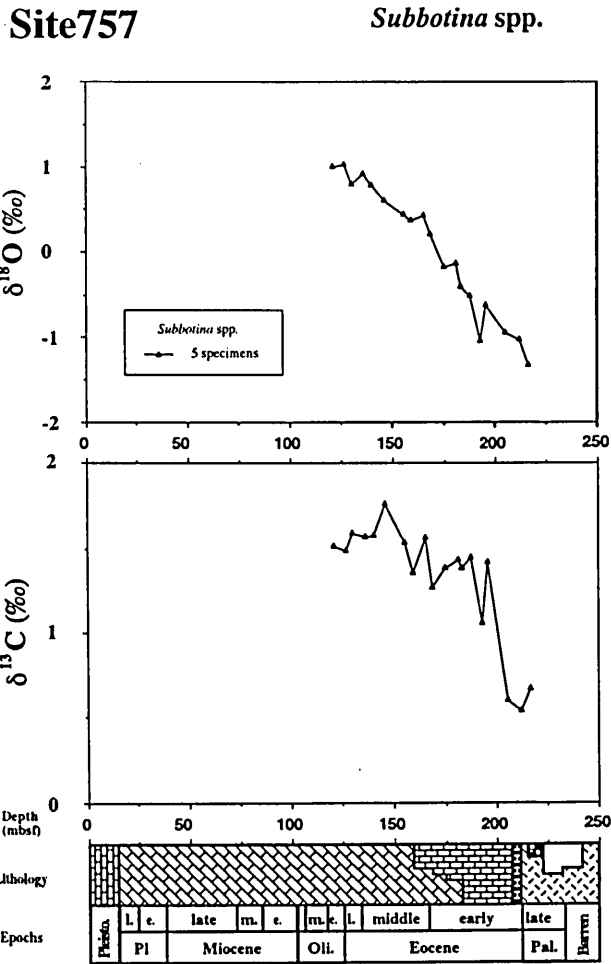


Fig. 22. Oxygen and carbon isotope records of planktonic foraminifer *Subbotina* spp. at Site 757.

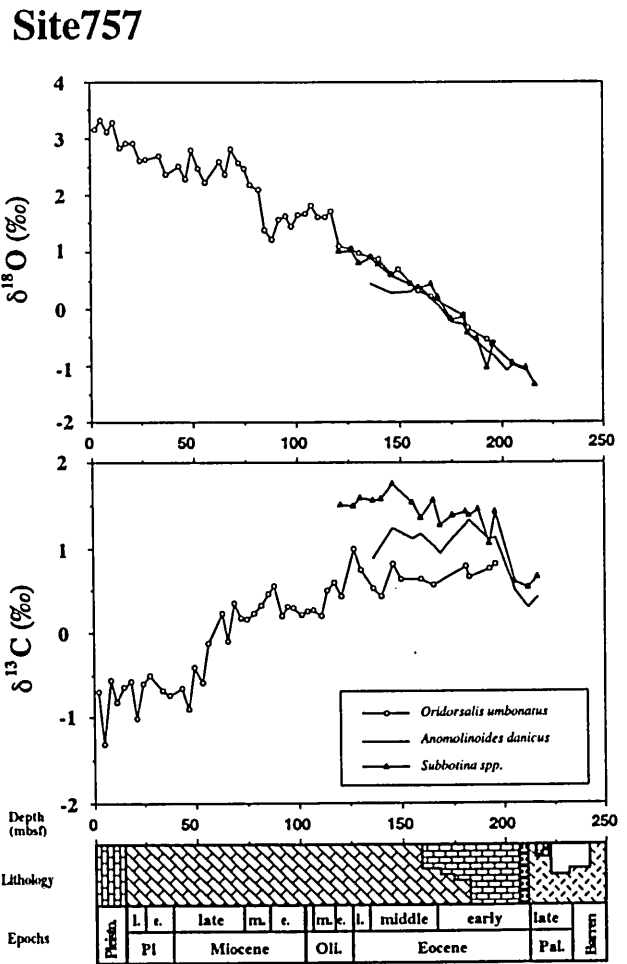


Fig. 23. Summary of oxygen and carbon isotope records at Site 757.

lower Eocene, and are constant (~1.4‰) up to 155.80 mbsf in Zone CP13b of the middle Eocene. The carbon isotopic record shows an increase up to a peak at 146.20 mbsf within Zone CP14 in the upper middle Eocene, and a gradual decrease up to 120.80 mbsf within Zones CP16a-b of the earliest Oligocene.

Isotopic values obtained from Site 757 are summarized in Fig. 23.

5. Site 758

At Site 758, oxygen and carbon isotopes were analyzed for seven benthic foraminiferal taxa, *Oridorsalis umbonatus*, *Anomalinoides danicus*, *Nuttallides truempyi*, *Stensioina beccariformis*, *Cibicidoides wuellerstorfi*, *Cibicidoides mundulus*, and *Groidinoides soldanii*, and for seven planktonic foraminiferal taxa, *Acarinina primitiva*, *A. praecursoria*, *Morozovella velascoensis*, *Subbotina pseudoecaena*, *S. eocaena*, *S. sp.1*, and *S. spp.* At this site, an unconformity has been recognized at ~250 mbsf, which lacks almost all of Eocene sequence (Peirce, Weissel, et al., 1989).

Oridorsalis umbonatus

Isotope analyses were performed for material from Samples 121-758A-1H-1, 75-80cm to 121-758A-27X-1, 75-80cm (0.75-248.05 mbsf) from the uppermost Eocene to Pleistocene. The isotopic records of *O. umbonatus* are shown in Fig. 24.

Oxygen isotopes: The oxygen isotopic records are constant at ~1.8‰, with a fluctuation from 248.05 to 135.15 mbsf through the uppermost Eocene to lower Miocene. In this interval, the amplitude of the fluctuation is ~0.4‰, and a significant negative peak (1.265‰) is observed at 190.05 mbsf above the Oligocene / Miocene boundary. δ¹⁸O values at 135.15 and 122.45 mbsf just above the early / middle Miocene boundary are relatively low (1.669 and 1.660‰, respectively). In the middle and early late Miocene, the oxygen isotopic records display a remarkable positive shift of 1.3‰ up to a weak peak at 112.85 mbsf (2.911‰). From 112.85 to 35.55 mbsf, δ¹⁸O values are constant at ~2.8‰ with smaller fluctuations (amplitude of ~0.3‰). In this section, two minor peaks are observed at 78.65 mbsf (nannofossil Zone CN9b) and 64.45 mbsf (Zones CN10-

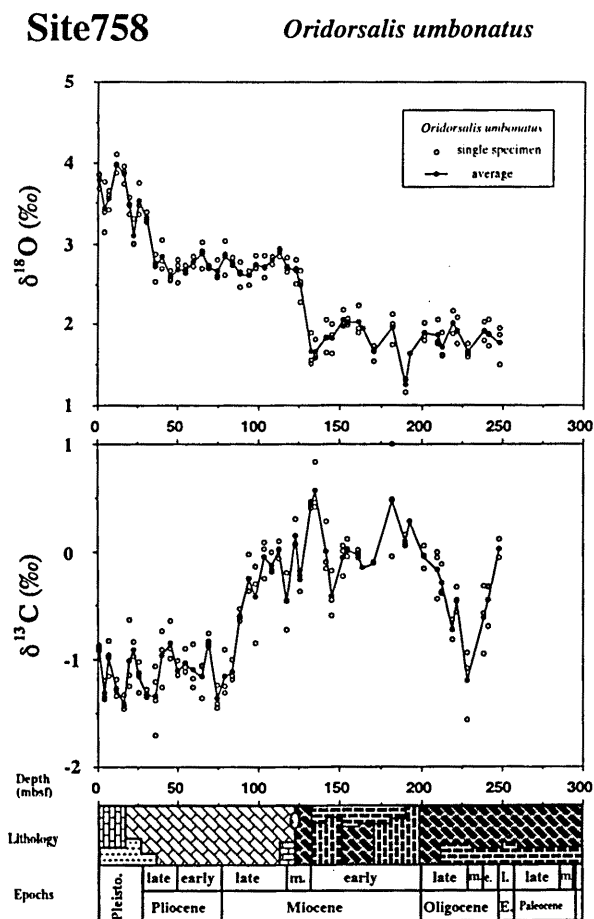


Fig. 24. Oxygen and carbon isotope records of benthic foraminifer *Oridorsalis umbonatus* at Site 758.

11a), with values of 2.850 and 2.845‰, respectively. Just across the Zone CN12 c/d boundary, a positive shift of 0.6‰ (from 2.731 to 3.322‰) in $\delta^{18}\text{O}$ value is recognized in the interval between 35.55 and 30.45 mbsf. From 30.45 to 0.75 mbsf, $\delta^{18}\text{O}$ values show increase with relatively large fluctuations. The amplitude of these fluctuations reaches $\sim 0.6\text{‰}$. An $\delta^{18}\text{O}$ peak value is observed at 11.25 mbsf just below the Zone CN14b/a boundary.

Carbon isotopes: $\delta^{13}\text{C}$ values decrease by 1.2‰ up to -1.196‰ at 228.76 mbsf (within Zone CP18), and immediately increase up to a peak value at 181.95 mbsf (0.480‰). The magnitude of this increase reaches 1.7‰. From 181.95 to 144.75 mbsf in the lower Miocene, $\delta^{13}\text{C}$ values decrease. In this section, a minor peak is observed at 154.45 mbsf (0.030‰). Across the early / middle Miocene boundary, a remarkable positive shift of $\sim 1.0\text{‰}$ occurred between 144.75 (-0.409‰) and 135.15 mbsf (0.569‰), and then $\delta^{13}\text{C}$ values are the highest among Neogene samples. In the middle Miocene, $\delta^{13}\text{C}$ values rapidly decrease by 0.9‰ from 0.569 to -0.255‰ , and increase by 0.4‰ up to a peak at 122.45 mbsf (0.151‰). From 117.35 to 93.55 mbsf in the lower upper Miocene, the carbon isotopic ratios are constant (with fluctuation) at $\sim 0.2\text{‰}$. The amplitude of

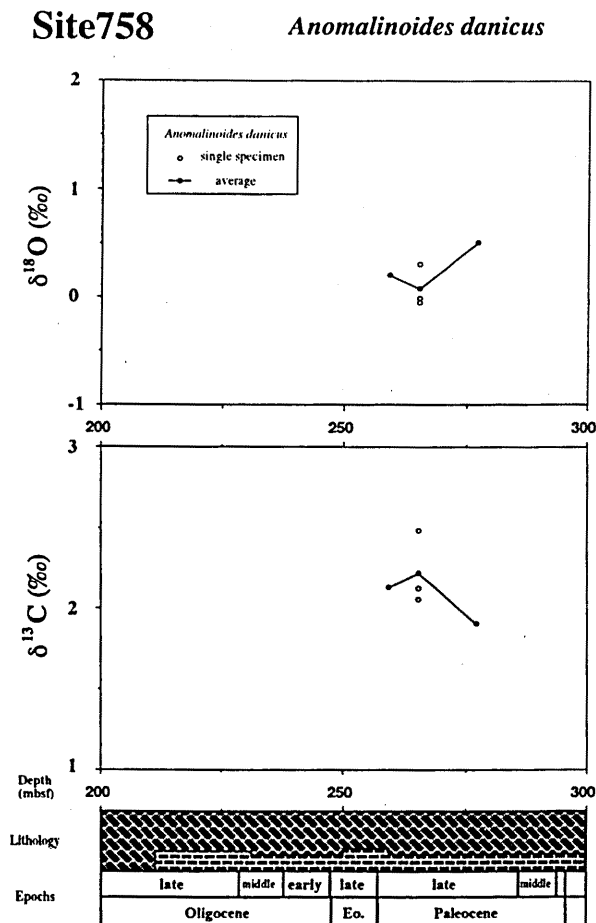


Fig. 25. Oxygen and carbon isotope records of benthic foraminifer *Anomalinoidea danicus* at Site 758.

fluctuation reaches to 0.5‰. A distinct negative shift of 1.15‰ is recognized from 93.55 (-0.247‰) to 74.15 (-1.362‰) mbsf in the uppermost Miocene. This shift is correlated with the Chron-6 Carbon Shift of the latest Miocene (Vincent et al., 1980; 1985). Throughout the Pliocene and Pleistocene, $\delta^{13}\text{C}$ values (about -1.1‰) fluctuate by 0.5‰. In this section, four minor peaks are observed at 68.95 mbsf within Zones 10-11a, 45.15 mbsf (within Zones 12a-c), 25.35 mbsf (within 13), and 6.75 mbsf (within Zones 14b-15), with values of -0.828 , -0.844 , -0.905 , and -0.981‰ , respectively.

Anomalinoidea danicus

Isotopes were analyzed for three samples from Samples 121-758A-28X-2, 75-80cm to 121-758A-30X-1, 75-80cm (265.15-277.05 mbsf) in the upper Paleocene, because of the rare occurrence of *A. danicus*. The isotopic records of *A. danicus* are shown in Fig. 25.

Oxygen and Carbon isotopes: $\delta^{18}\text{O}$ values tend to decrease, whereas $\delta^{13}\text{C}$ values tend to increase. The minimum $\delta^{18}\text{O}$ value and the maximum $\delta^{13}\text{C}$ value are 0.085 and 2.220‰, respectively.

Nuttallides truempyi

Isotope analyses were performed for two to five individuals and single specimen in each sample. The difference in values between the measured individuals is relatively small. No significant difference in the two measuring methods is recognized, except for Sample 121-758A-28X-6, 75-80 cm. At this site, the isotopic records were obtained from Samples 121-758A-28X-2, 75-80 cm to 121-758A-31X-5, 75-80 cm (259.15-292.65 mbsf) in the Paleocene. The isotopic records of *N. truempyi* are shown in Fig. 26.

Oxygen isotopes: *Nuttallides truempyi* $\delta^{18}\text{O}$ values increase up to a peak of 277.05 mbsf, and decrease with a small fluctuations.

Carbon isotopes: The carbon isotopic record shows an increase from 0.77 to 1.82‰. The magnitude of the increase is 1.1‰ in the studied section.

Stensioina beccariiformis

Isotopes were analyzed for Samples 121-758A-28X-2, 75-80cm to 121-758A-31X-5, 75-80cm (259.15-292.65 mbsf) in the Paleocene. The isotopic records of *S. beccariiformis* are shown in Fig. 27.

Oxygen isotopes: $\delta^{18}\text{O}$ values increase up to a peak at

277.05 mbsf (0.526‰) within the nannofossil Zones CP6-7 of the upper Paleocene, and decrease from 277.05 mbsf (0.526‰) to 265.15 mbsf (-0.140‰).

Carbon isotopes: $\delta^{13}\text{C}$ values increase up to a maximum value at 267.35 mbsf (2.071‰) within Zones CP6-7 in the upper Paleocene. The magnitude of this increase reaches 1.2‰.

Other Benthic foraminifers

Isotope analyses were performed for specimens from Samples 121-758A-9H-4, 75-80cm to 121-758A-27X-1, 75-80cm (78.65-248.05 mbsf) for *Groidinoides soldanii*, from Samples 121-758A-10H-1, 75-80cm to 121-758A-27X-1, 75-80cm (81.85-248.05 mbsf) for *Cibicidoides mundulus*, and from Samples 121-758A-12X-1, 75-80cm to 121-758A-14X-3, 75-80cm (103.15-125.45 mbsf) for *Cibicidoides wuellerstorfi* in the Miocene and Oligocene. The isotopic records of those species are shown in Fig. 28.

Oxygen isotopes: The trend of isotopic records of those species resembles those of *Oridorsalis umbonatus*. $\delta^{18}\text{O}$ values show a general increase. The $\delta^{13}\text{C}$ values for *Groidinoides soldanii* decrease by 0.75‰ from 248.05 to 228.76 mbsf in the lower Eocene, and are constant at ~0.4‰ up to 122.45 mbsf in the middle Miocene. During the late

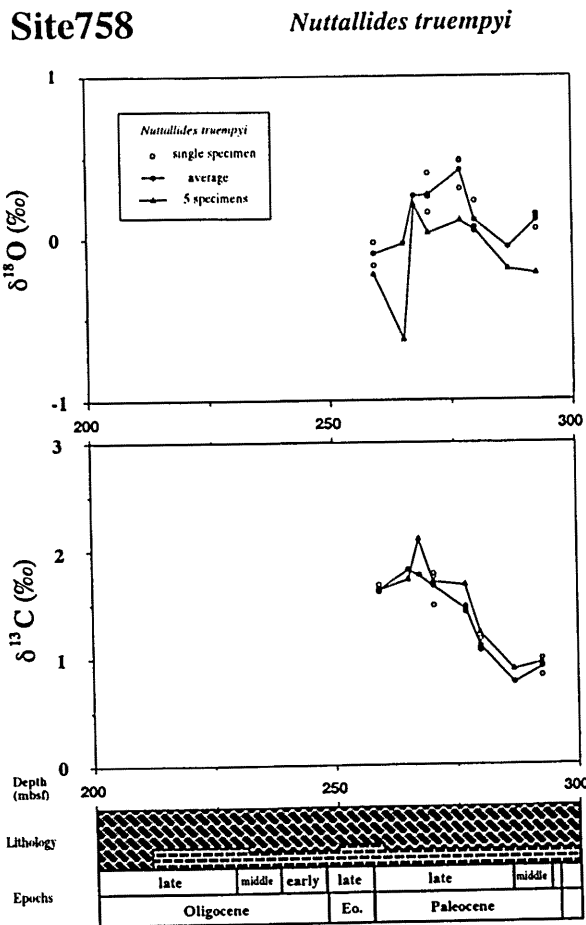


Fig. 26. Oxygen and carbon isotope records of benthic foraminifer *Nuttallides truempyi* at Site 758.

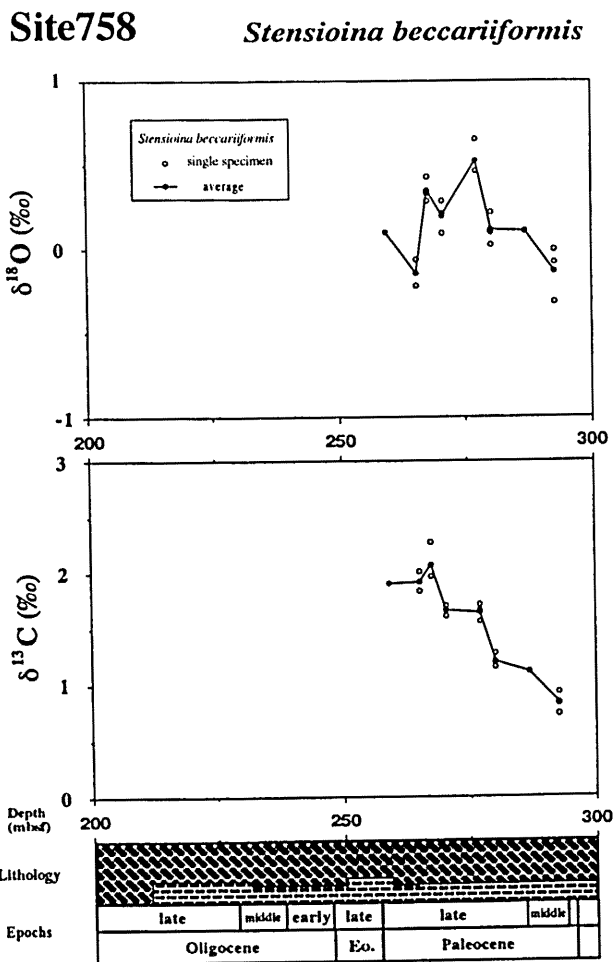


Fig. 27. Oxygen and carbon isotope records of benthic foraminifer *Stensioina beccariiformis* at Site 758.

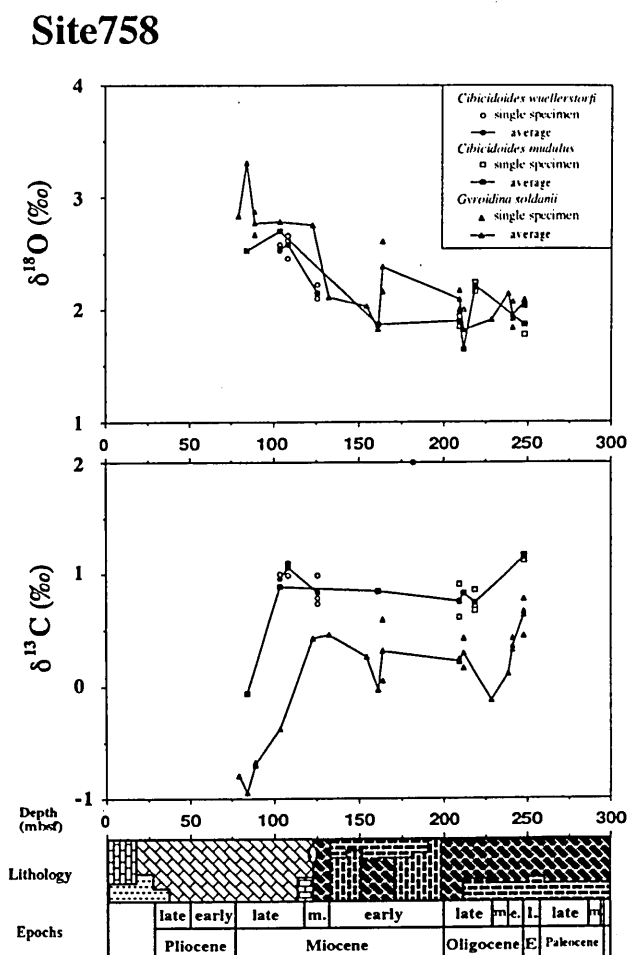


Fig. 28. Oxygen and carbon isotope records of other benthic foraminifers at Site 758.

Miocene, the *Gvroidinoides soldanii* $\delta^{13}\text{C}$ values decrease. The $\delta^{13}\text{C}$ values of *Cibicidoides mundulus* are constant ($\sim 0.8\%$) up to 103.15 mbsf within Zones CN8-9a in the upper Miocene, and exhibit a decreasing trend. The distribution of $\delta^{13}\text{C}$ in *Cibicidoides wuellerstorfi* resembles that of *Cibicidoides mundulus*.

Subbotina group

At Site 758, isotopes of the *Subbotina* group were analyzed for four taxa. Isotope analyses of *Subbotina* spp. were performed for five individuals (253-367 μm in diameter). The isotope analyses of other taxa were made on a single individual ($>380\mu\text{m}$ in diameter). The isotopic records of *Subbotina* group were obtained from Samples 121-758A-28X-2, 75-80cm to 121-758A-31X-5, 75-80cm (259.15-292.65 mbsf) in the Paleocene. The isotopic records of *Subbotina* group are shown in Fig. 29.

The oxygen isotopic record of *Subbotina* spp. shows a general decrease from 280.05 mbsf to 265.15 mbsf in the late Paleocene. $\delta^{18}\text{O}$ values at 280.05 mbsf are relatively high, and are the maximum values among the Paleocene samples. Between 367.35 and 265.15 mbsf, $\delta^{18}\text{O}$ values show a negative shift of 0.4‰. The same trend of $\delta^{18}\text{O}$ values are

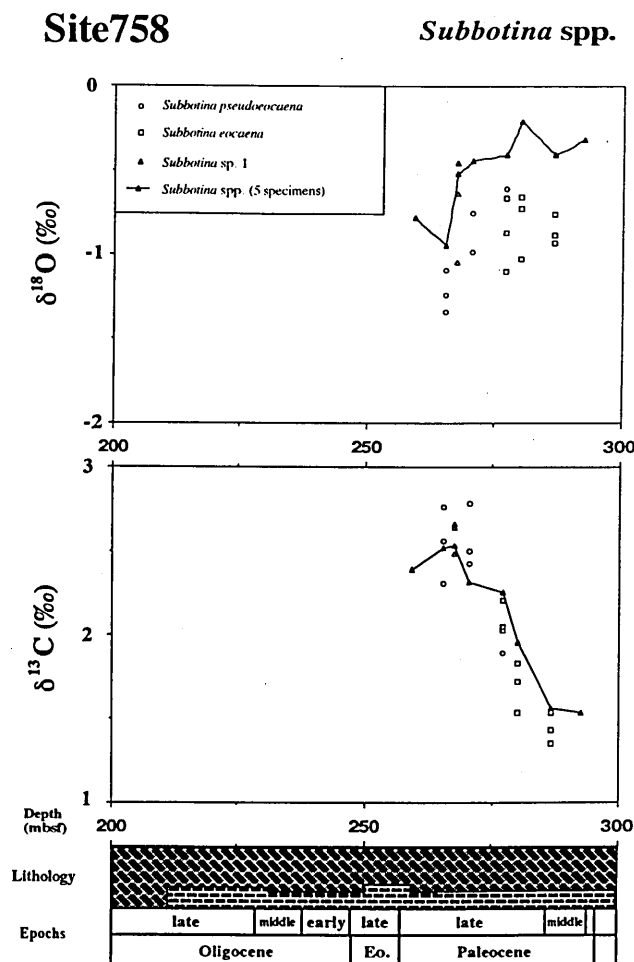


Fig. 29. Oxygen and carbon isotope records of planktonic foraminifer *Subbotina* group at Site 758.

recognized, but the identified taxa such as *Subbotina* spp. show an offset of *Subbotina* spp. toward low values by 0.3-0.5‰.

Carbon isotopes: $\delta^{13}\text{C}$ values of *Subbotina* spp. show an increase with stratigraphic height up to 267.35 mbsf (within Zones CP6-7), then decrease. The magnitude of this increase reaches 1.0‰, and the peak of this increase exhibits the maximum value (2.529‰) among the Paleocene samples. The carbon isotopic signals of other *Subbotina* show the same trend as *Subbotina* spp.

Other Planktonic foraminifers

The isotopes of *Morozovella velascoensis* were analyzed for Samples 121-758A-28X-2, 75-80cm to 121-758A-31X-1, 75-80cm (259.15-286.65 mbsf) from the upper Paleocene. *Acarinina primitiva* and *A. praecursoria* were used in Sample 121-758A-28X-2, 75-80cm (259.15) in the upper Paleocene and Sample 121-758A-31X-5, 75-80cm (292.65 mbsf) from the lower Paleocene, respectively. The isotopic records of those species are shown in Fig. 30.

Oxygen and Carbon isotopes: The oxygen isotopic records of *M. velascoensis* show a constant trend (about -1.5‰). In this section, $\delta^{18}\text{O}$ values at 265.15 mbsf are relatively low.

Site758

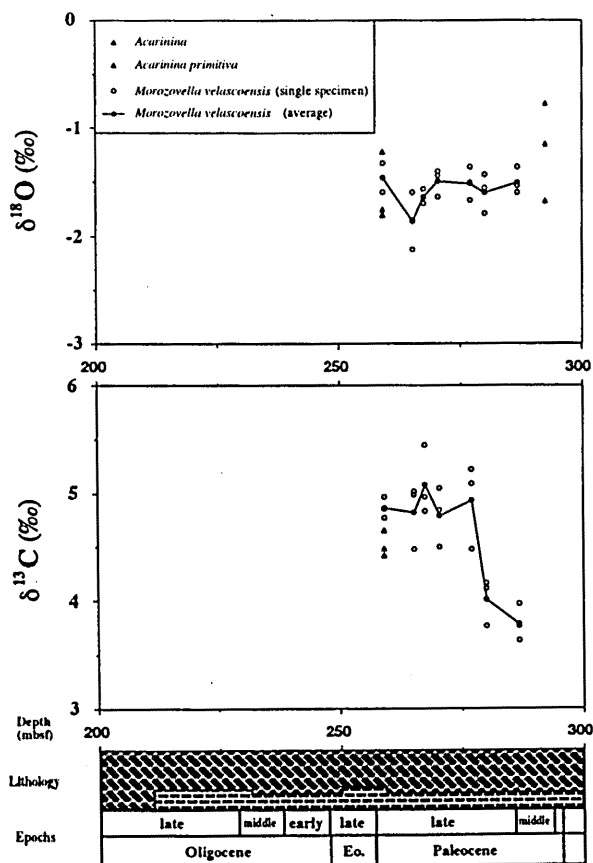


Fig. 30. Oxygen and carbon isotope records of other planktonic foraminifers at Site 758.

$\delta^{13}\text{C}$ values of *M. velascoensis* rapidly increase by $\sim 1.0\text{‰}$ up to 277.05 mbsf, and then are constant at $\sim 4.8\text{‰}$. A minor peak is observed at 267.35 mbsf (5.084), with the maximum value for the Paleocene. $\delta^{18}\text{O}$ of *Acarinina primitiva* values are close to those of *M. velascoensis*, whereas their $\delta^{18}\text{O}$ values are slightly lower than *M. velascoensis*. The isotopic ratios of *A. praecursoria* are similar to those of *Subbotina* group rather than *M. velascoensis*.

Isotopic values obtained from Site 758 are summarized in Fig. 31.

6. Site 762

At Site 762, oxygen and carbon isotopes were analyzed for four benthic foraminiferal taxa: *Anomalinoidea danicus*, *Nuttallides truempyi*, *Oridorsalis umbonatus*, and *Stensioina beccariiiformis*. At this site, isotope analyses were limited to the lower Eocene and upper Paleocene (370–490 mbsf). The analysis of *O. umbonatus* was only available in Sample 122-762A-24X-2, 77–82cm (379.77 mbsf) from the earliest Eocene, and values -0.628‰ for oxygen isotopes and -0.547‰ for carbon isotopes.

Site758

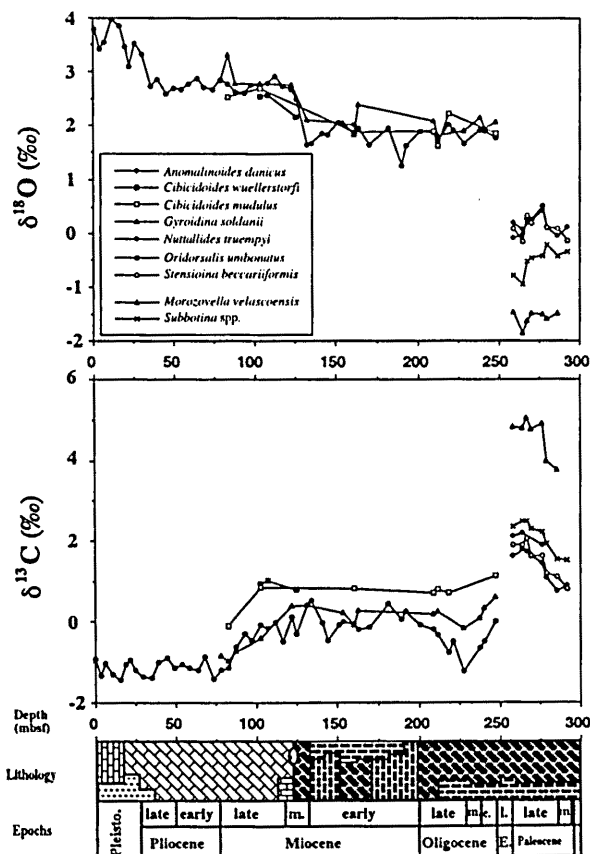


Fig. 31. Summary of oxygen and carbon isotope records at Site 758.

Anomalinoidea danicus

Isotopes were determined for Samples 122-762C-26X-1, 70–73cm to 122-762C-32X-1, 65–69cm (398.70–450.65 mbsf). The occurrence of *A. danicus* is rare in all samples. The isotopic records of *A. danicus* are shown in Fig. 32.

Oxygen isotopes: $\delta^{18}\text{O}$ values tend to increase up to 434.69 mbsf, and reaching a maximum value (0.325‰). $\delta^{18}\text{O}$ values decrease by $\sim 1.0\text{‰}$ from 434.69 to 398.16 mbsf.

Carbon isotopes: The carbon isotopic records show an increase up to 425.19 mbsf, with small fluctuations, and the maximum value of $\delta^{13}\text{C}$ is 2.623‰ . $\delta^{13}\text{C}$ values drastically decrease from 425.19–404.70 mbsf across the Paleocene / Eocene boundary. The magnitude of this decrease reaches $\sim 2.2\text{‰}$. Above 404.70 mbsf, $\delta^{13}\text{C}$ values are constant at $\sim 0.4\text{‰}$. However, a weak negative peak is observed at 400.21 mbsf (0.120‰).

Nuttallides truempyi

Isotope analyses were performed for two to five individuals and single specimen in each sample. No significant differences were recognized for the different analyses. At this site, isotopic records were obtained from

Samples 122-762C-23X-1, 69-74cm to 122-762A-35X-1, 68-73cm (370.19-497.18 mbsf). The isotopic records of *N. truempyi* are shown in Fig. 33.

Oxygen isotopes: The records of *N. truempyi* $\delta^{18}\text{O}$ show an increase, with small fluctuations up to 450.65 mbsf in the upper Paleocene. $\delta^{18}\text{O}$ values decrease by 0.9‰ (-0.138 to -1.055‰) up to 379.77 mbsf across the Paleocene / Eocene boundary. Above this, oxygen isotopic records exhibit an increasing trend.

Carbon isotopes: Carbon isotope record increase, with fluctuations up to 431.69 mbsf. The magnitude of increase is 1.0‰. In this section, A ^{13}C maximum is recognized at 447.18 mbsf (1.748‰). $\delta^{13}\text{C}$ values drastically decrease by 1.9‰ from 1.812‰ (431.69 mbsf) to -0.088‰ (406.18 mbsf) across the benthic event. From 406.18 to 379.77 mbsf, $\delta^{13}\text{C}$ records are a constant with small fluctuations at ~ 0.1 ‰. A ^{13}C minimum is observed at 400.21 mbsf (-0.487‰). The carbon isotopic records tend to increase above this level.

Stensioina beccariiiformis

Isotope analyses were performed for Samples 122-

762C-29X-1, 76-80cm to 122-762C-35X-1, 68-73cm (422.26-479.18 mbsf) in the upper Eocene. *S. beccariiiformis* does not occur above Sample 122-762C-29X-1, 76-80cm (422.26 mbsf) below the Paleocene / Eocene boundary. The isotopic records of *S. beccariiiformis* are shown in Fig. 34.

Oxygen isotopes: The oxygen isotopic record of *S. beccariiiformis* shows an increasing trend, with fluctuations, up to a maximum value (0.165‰) at 434.69 mbsf, with an increasing trend of 0.65‰. In this section, a notable peak is observed at 469.61 mbsf. From a maximum value at 434.69 mbsf, $\delta^{18}\text{O}$ values decrease to -0.525‰ at 422.26 mbsf.

Carbon isotopes: In the late Paleocene, averaged $\delta^{13}\text{C}$ values increase by ~ 1.1 ‰ with relatively large fluctuations up to a maximum value (1.946‰) at 423.69 mbsf. In this section, two remarkable peaks are recognized at 469.61 and 447.18 mbsf. $\delta^{13}\text{C}$ values rapidly decrease by 0.6‰ up to 422.26 mbsf, located just below the level of the benthic event, where *S. beccariiiformis* disappears.

Isotopic values obtained from Site 762 are summarized in Fig. 35.

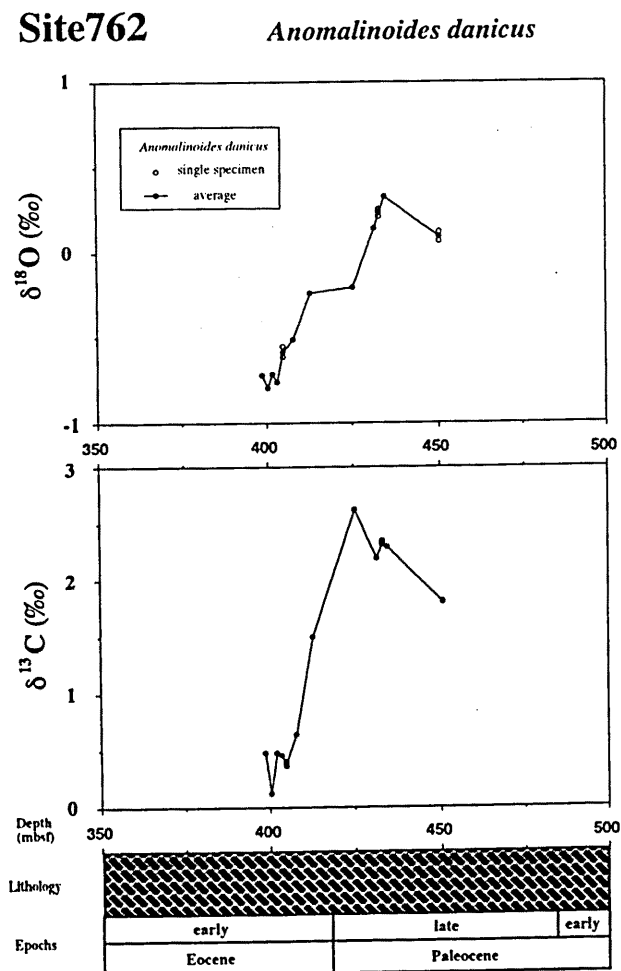


Fig. 32. Oxygen and carbon isotope records of benthic foraminifer *Anomalinoidea danicus* at Site 762.

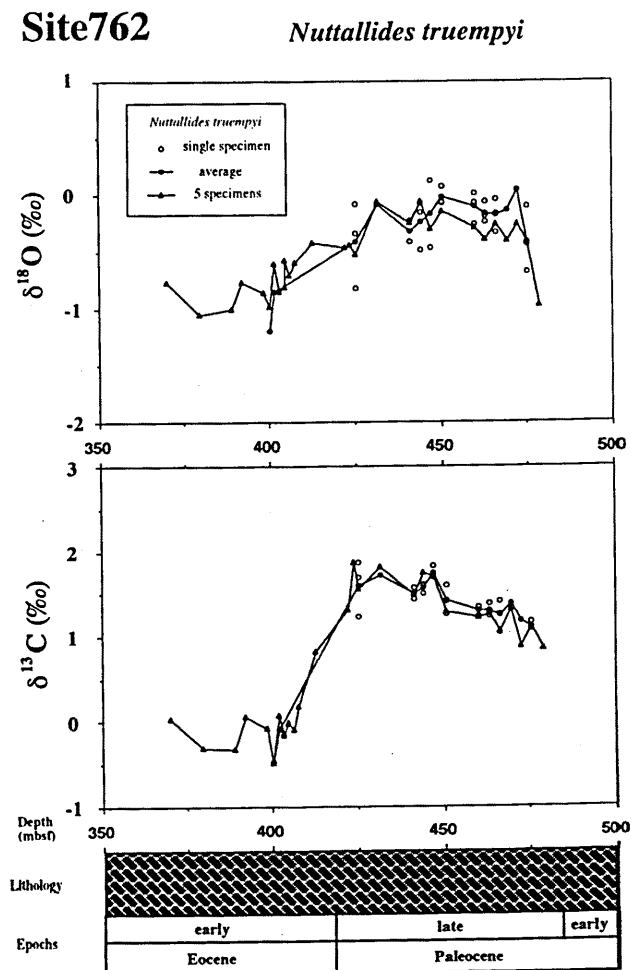


Fig. 33. Oxygen and carbon isotope records of benthic foraminifer *Nuttallides truempyi* at Site 762.

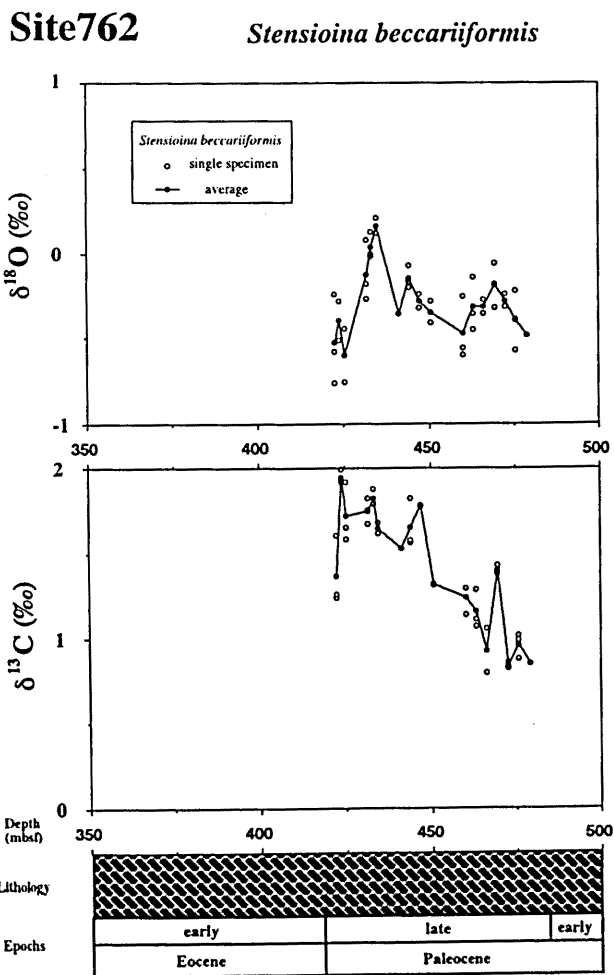


Fig. 34. Oxygen and carbon isotope records of benthic foraminifer *Stensioina beccariiiformis* at Site 762.

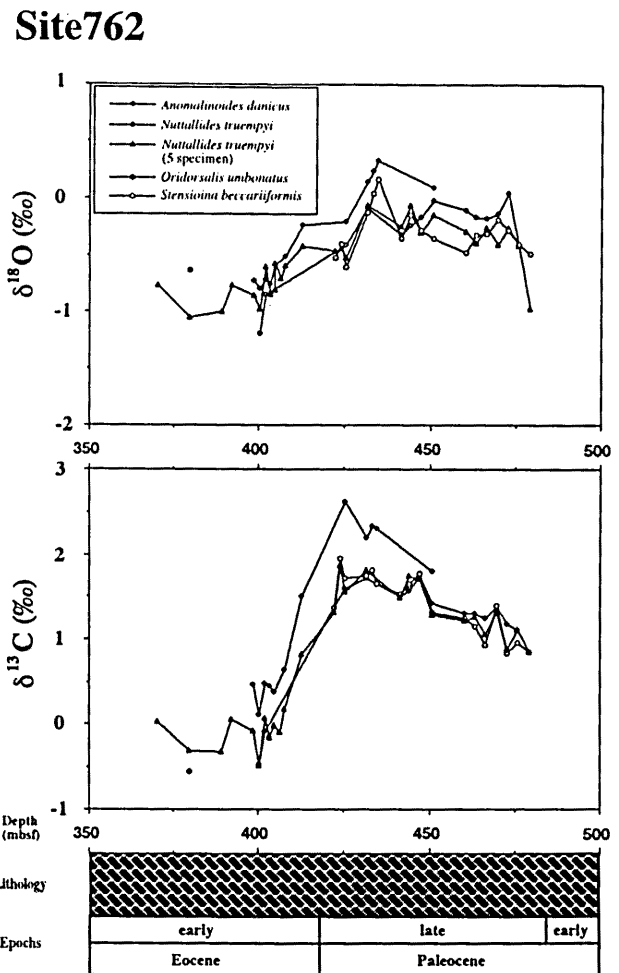


Fig. 35. Summary of oxygen and carbon isotope records at Site 762.

III. Stable isotopic paleoceanography in the South Atlantic and Indian Oceans

A. Compiled data

In and around the Indian and South Atlantic Oceans, oxygen and carbon isotopic data of Cenozoic foraminifera have been obtained at many DSDP and ODP sites. Information from these sites, including this study, is shown in Table 1. The Indian and South Atlantic Oceans are divided into six areas: the northeastern Indian Ocean (Sites 214, 215, 216, 253, 752, 754, 756, 757, 758, and 762), northwestern Indian Ocean (Sites 237, 238, 709, 714, and 716), southern Indian Ocean (Sites 538, 744, 748, 750, and 751), Central Atlantic Ocean (Sites 366, 658, 659, 665, and 667), South Atlantic Ocean (Sites 516, 517, 518, 519, 521, 522, 523, 525, 526, 527, 528, and 529), and Southern Ocean of the Atlantic (Sites 658, 689, 690, 699, 700, 702, and 704). The oxygen and carbon isotopic data of these areas are compiled in order to compare the Cenozoic paleoceanography of the Indian and South Atlantic Oceans. However, there are limited resources to compile part of the section, such as the Paleogene section of the northwestern Indian Ocean, Neogene of the Southern Ocean (Atlantic sector), and deep sea sections below 3000 m.

1. Adjustment of data

Although oxygen and carbon isotopes have been measured from foraminiferal tests at many sites, these comprise different foraminiferal species. In the studied area, the benthic foraminifer *Oridorsalis umbonatus* has been commonly measured, and *Nuttallides truempyi*, *Cibicidoides* spp., *Cibicidoides wuellerstorfi* have also been measured. For the purpose of this study, the isotopic values of sea water should be indicated by the values of different species. However, benthic species may draw in pore water and overlying water as carbon dioxide pools (Shackleton *et al.*, 1984; Woodruff *et al.*, 1990; Zachos *et al.*, 1992a; 1992c; and other), and thus different species record different values according to their paleoecology. Therefore, the oxygen and carbon isotopic data cannot be directly compared, because of interspecific variation in the isotopic values. Most workers use the adjusted values obtained by Shackleton *et al.* (1984). Barrera and Huber (1990; 1991), however, pointed out that the magnitudes of departure in the Maastrichtian and Paleogene differ from those of Shackleton *et al.* (1984). Interspecific difference vary according to region and age. Therefore, adjustment should be made at each site based on

Table 1. Sources for foraminiferal oxygen and carbon isotopic data in the studied area.

Site	Latitude	Longitude	Geographical area	Water depth (m)	Species	Range		Reference
						mbsf	Age (Ma)	
Site 214	11°20.21' S	88°43.08' E	Indian Ocean	1655	<i>Globigerinoides sacculifer</i>	86.34-200.48	5-18	Vincent et al., 1985
					<i>Dentoglobigerina altispira</i>	123.92-200.48	7-18	Vincent et al., 1985
					<i>Globoquadrina venezuelana</i>	86.34-217.95	5-23	Vincent et al., 1985
					<i>Oridorsalis umbonatus</i>	86.34-224.33	5-24	Vincent et al., 1985
Site 216	1°27.73' N	90°12.48' E	Indian Ocean	2247	<i>Globigerinoides sacculifer</i>	109.32-188.10	8-23	Vincent et al., 1985
					<i>Dentoglobigerina altispira</i>	109.32-174.39	8-21	Vincent et al., 1985
					<i>Globoquadrina venezuelana</i>	109.32-199.78	8-23	Vincent et al., 1985
					<i>Oridorsalis umbonatus</i>	109.32-199.78	8-23	Vincent et al., 1985
Site 237	7°04.99' S	58°07.48' E	Indian Ocean	1640	<i>Globigerinoides sacculifer</i>	109.82-175.42	8-20	Vincent et al., 1985
					<i>Dentoglobigerina altispira</i>	83.32-184.24	5-23	Vincent et al., 1985
					<i>Globoquadrina venezuelana</i>	83.32-200.32	5-24	Vincent et al., 1985
					<i>Oridorsalis umbonatus</i>	83.32-200.32	5-24	Vincent et al., 1985
Site 238	11°09.20' S	70°31.56' E	Indian Ocean	2844	<i>Globigerinoides sacculifer</i>	132.66-439.03	5-24	Vincent et al., 1985
					<i>Dentoglobigerina altispira</i>	132.66-398.86	5-24	Vincent et al., 1985
					<i>Globoquadrina venezuelana</i>	132.66-455.83	5-24	Vincent et al., 1985
					<i>Oridorsalis umbonatus</i>	132.66-455.83	5-24	Vincent et al., 1985
Site 253	24°52.65' S	87°21.97' E	Indian Ocean	1962	various species (Planktonic)	19.31-132.48	4-37	Oberhansli, 1986
					various species (Benthic)	19.31-132.48	4-37	Oberhansli, 1986
Site 366	5°40.7' N	19°51.1' W	Eastern Tropical Atlantic	2853	<i>Cibicides</i> spp.	114.68-433.46	8-38	Miller et al., 1989
Site 516	30°16.6' S	35°37.1' W	western South Atlantic	1313	<i>Globigerinoides sacculifer</i>	8.80-37.50	2-4	Leonard et al., 1983
					<i>Cibicidoides wuellerstorfi</i>	8.90-37.50	2-4	Leonard et al., 1983
Site 517	30°56.8' S	38°02.5' W	western South Atlantic	2963	<i>Globigerinoides sacculifer</i>	24.70-50.70	1-3	Leonard et al., 1983
					<i>Cibicidoides wuellerstorfi</i>	36.50-50.70	2-3	Hodell et al., 1983
Site 518	29°58.4' S	38°08.1' W	western South Atlantic	3944	<i>Orbulina universa</i>	28.0-43.70	2-4	Hodell et al., 1983
					<i>Cibicidoides wuellerstorfi</i>	28.0-44.0	2-4	Hodell et al., 1983
Site 519	26°08.20' S	11°39.97' W	South Atlantic	3779	<i>Globigerinoides sacculifer</i>	33.15-95.15	0-4	Weissert et al., 1984
					<i>Dentoglobigerina altispira</i>	64.25-101.00	3-4	Weissert et al., 1984
					<i>Orbulina universa</i>	101.40-112.50	4-5	Weissert et al., 1984
					<i>Cibicidoides wuellerstorfi</i>	31.65-112.50	0-5	Weissert et al., 1984
					<i>Nuttallides umbonifera</i>	31.65-101.00	0-4	Weissert et al., 1984
					<i>Orbulina universa</i>	97.20-146.80	0-9	McKenzie et al., 1984
					<i>Cibicidoides wuellerstorfi</i>	31.60-144.6	0-6	McKenzie et al., 1984
					<i>Nuttallides umbonifera</i>	31.60-121.0	5-6	McKenzie et al., 1984
					<i>Oridorsalis umbonatus</i>	115.60-122.2	4-9	McKenzie et al., 1984
					<i>Globigerinoides sacculifer</i>	13.24-37.37	1-4	Weissert et al., 1984
					<i>Oridorsalis umbonatus</i>	31.33-41.60	3-5	Weissert et al., 1984
Site 521	26°04.43' S	10°15.87' W	South Atlantic	4141	<i>Cibicidoides wuellerstorfi</i>	12.40-45.31	1-6	Weissert et al., 1984
					<i>Nuttallides umbonifera</i>	13.92-45.31	1-6	Weissert et al., 1984
					<i>Globigerinoides conglobatus</i>	11.11-16.03	1-2	Weissert et al., 1984
Site 522	26°06.84' S	05°06.78' W	South Atlantic	4457	<i>Oridorsalis umbonatus</i>	16.58-30.15	2-6	Weissert et al., 1984
					<i>Cibicidoides wuellerstorfi</i>	16.58-30.15	2-6	Weissert et al., 1984
					<i>Nuttallides umbonifera</i>	11.11-30.15	1-6	Weissert et al., 1984
					various species (Planktonic)	66.45-145.52	26-36	Poore and Matthews, 1984
					various species (Benthic)	66.45-145.52	26-36	Poore and Matthews, 1984
					<i>Globoquadrina venezuelana</i>	130.7-145.4	35-36	Oberhansli et al., 1984
					<i>Catapsydrax dissimilis</i>	130.20-146.80	35-36	Oberhansli et al., 1984
Site 523	28°33.13' S	02°15.08' W	South Atlantic	4572	<i>Stilostomella</i> spp.	130.20-146.80	35-36	Oberhansli et al., 1984
					<i>Oridorsalis umbonatus</i>	119.20-140.40	34-36	Oberhansli et al., 1984
					<i>Globigerinoides ruber</i>	5.75-25.45	1-3	Weissert et al., 1984
					<i>Oridorsalis umbonatus</i>	9.63-24.70	1-3	Weissert et al., 1984
					<i>Cibicidoides wuellerstorfi</i>	5.75-25.45	1-3	Weissert et al., 1984
					<i>Nuttallides umbonifera</i>	9.63-24.70	1-3	Weissert et al., 1984
					various species (Planktonic)	85.90-184.00	34-48	Oberhansli et al., 1984
Site 524	29°29.05' S	03°30.74' E	South Atlantic	4806	various species (Benthic)	85.90-184.00	34-48	Oberhansli et al., 1984
					various species (Planktonic)	9.50-117.30	55-56	Oberhansli et al., 1984
					<i>Stensiolina beccariiiformis</i>	98.3	61	Oberhansli et al., 1984
					<i>Nuttallides truempyi</i>	9.50-28.10	55-56	Oberhansli et al., 1984
					<i>Oridorsalis umbonatus</i>	9.50-18.90	55-56	Oberhansli et al., 1984
					<i>Stensiolina beccariiiformis</i>	164.57-230.19	61-67	He et al., 1984
					<i>Oridorsalis umbonatus</i>	164.57-230.19	61-67	He et al., 1984
Site 525	29°04.24' S	02°59.12' E	South Atlantic	2467	various species (Benthic)	0.8-470.91	0-68	Shackleton et al., 1984
					various species (Planktonic)	99.37-433.72	8-62	Shackleton et al., 1984
Site 526	29°07.36' S	03°08.28' E	South Atlantic	1054	various species (Benthic)	6.71-217.53	1-37	Shackleton et al., 1984
					various species (Planktonic)	30.80-214.71	4-37	Shackleton et al., 1984
Site 527	28°02.49' S	01°45.80' E	South Atlantic	4428	various species (Benthic)	40.94-282.20	4-67	Shackleton et al., 1984
					various species (Planktonic)	146.77-283.14	50-67	Shackleton et al., 1984
Site 528	28°31.49' S	02°19.44' E	South Atlantic	3800	various species (Benthic)	8.87-315.99	1-58	Shackleton et al., 1984
					various species (Planktonic)	8.87-315.99	1-58	Shackleton et al., 1984
Site 529	28°55.83' S	02°46.08' E	South Atlantic	3035	various species (Benthic)	0.11-268.84	0-58	Shackleton et al., 1984
					various species (Planktonic)	0.11-268.84	0-58	Shackleton et al., 1984

Table 1. (continued).

Site	Latitude	Longitude	Geographical area	Water depth (m)	Species	Range		Reference					
						mbsf	Age (Ma)						
Site 658	20°44.95' N	18°34.87' W	Eastern Tropical Atlantic	2263	<i>Globorotalia inflata</i>	0.24-136.07	0-2	Samthein and Tiedemann, 1989					
					<i>Cibicides wuellerstorfi</i>	0.24-293.01	0-4	Samthein and Tiedemann, 1989					
					<i>Uvigerina auberiana</i>	151.73-293.01	2-4	Samthein and Tiedemann, 1989					
					<i>Uvigerina peregrina</i>	151.73-293.01	2-4	Samthein and Tiedemann, 1989					
Site 659	18°04.63' N	21°01.57' W	Eastern Tropical Atlantic	3071	<i>Globorotalia inflata</i>	0.47-32.59	0-1	Samthein and Tiedemann, 1989					
					<i>Cibicides wuellerstorfi</i>	0.47-32.92	0-1	Samthein and Tiedemann, 1989					
Site 662	1°23.41' S	11°44.35' W	Eastern Tropical Atlantic	3824		117.39-136.09	1-2	Karlin et al., 1989					
Site 663	1°11.87' S	11°52.71' W	Eastern Tropical Atlantic	3708	<i>Globigerinoides ruber</i>	4.81-32.21	0-1	Rucklman and Jancecek, 1989					
Site 665	2°57.07' N	19°40.07' W	Eastern Tropical Atlantic	4752	<i>Cibicides</i> spp.	41.06-57.93	2	Curry and Miller, 1989					
Site 667	4°34.15' N	21°54.68' W	Eastern Tropical Atlantic	3529	<i>Cibicides</i> spp.	220.16-375.16	19-31	Miller et al., 1989					
Site 689	64°31' S	03°06' E	Weddell Sea	2080	various species (Planktonic)	102.10-233.10	31-66	Stott et al., 1990					
					various species (Planktonic)	227.51-261.8	64-68	Stott and Kennett, 1990					
					<i>Archaeoglobigerina australis</i>	246.87-291.30	68-76	Barrera and Huber, 1990					
					<i>Globigerinelloides multispinatus</i>	246.87-294.50	68-77	Barrera and Huber, 1990					
					<i>Abathomphalus mayaroensis</i>	246.87-256.38	68-69	Barrera and Huber, 1990					
					<i>Nuttallides</i> spp.	165.90-236.47	18-67	Kennett and Stott, 1990					
					<i>Cibicides</i> spp.	93.15-233.54	18-66	Kennett and Stott, 1990					
					<i>Nuttallides truempyi</i>	246.87-256.38	68-69	Barrera and Huber, 1990					
					<i>Stensioina beccariiiformis</i>	246.87-294.50	68-77	Barrera and Huber, 1990					
					<i>Coryphostoma incrassata</i>	246.87-269.50	66-71	Barrera and Huber, 1990					
					Site 690	65°10' S	01°12' E	Weddell Sea	2914	various species (Planktonic)	93.95-247.76	38-66	Stott et al., 1990
										various species (Planktonic)	224.74-242.94	63-65	Stott and Kennett, 1990
										<i>Archaeoglobigerina australis</i>	263.01-314.52	68-72	Barrera and Huber, 1990
<i>Globigerinelloides multispinatus</i>	258.00-316.62	67-72	Barrera and Huber, 1990										
<i>Abathomphalus mayaroensis</i>	258.00-277.87	67-70	Barrera and Huber, 1990										
<i>Nuttallides truempyi</i>	53.00-206.06	26-59	Kennett and Stott, 1990										
<i>Cibicides</i> spp.	50.75-212.82	20-60	Kennett and Stott, 1990										
<i>Nuttallides truempyi</i>	258.00-283.81	68-71	Barrera and Huber, 1990										
<i>Stensioina beccariiiformis</i>	263.01-316.62	68-72	Barrera and Huber, 1990										
<i>Coryphostoma incrassata</i>	258.00-300.40	67-71	Barrera and Huber, 1990										
Site 698	51°27.51' S	33°05.96' W	South Atlantic	2128						<i>Nuttallides truempyi</i>	45.20-73.40	55-58	Katz and Miller, 1991
										<i>Cibicides</i> spp.	4.86-81.71	52-59	Katz and Miller, 1991
Site 699	51°32.537' S	30°40.619' W	South Atlantic	3708						<i>Nuttallides truempyi</i>	442.94-498.32	52-57	Katz and Miller, 1991
Site 700	51°31.977' S	30°16.688' W	South Atlantic	3598	<i>Nuttallides truempyi</i>	131.20-329.18	50-65	Katz and Miller, 1991					
Site 702	50°56.786' S	26°22.117' W	South Atlantic	3083	<i>Cibicides</i> spp.	67-290.49	46-61	Katz and Miller, 1991					
					<i>Nuttallides truempyi</i>	32.23-274.10	39-59	Katz and Miller, 1991					
Site 704	46°52.8' S	7°25.3' E	South Atlantic	2532	<i>Neogloboquadrina puckylerma</i>	0.41-233.41	0-5	Hodell and Ciesielski, 1991					
					<i>Cibicides</i> spp.	2.28-53.90	0-1	Hodell and Ciesielski, 1991					
					various species (Benthic)	214.21-384.71	4-9	Muller et al., 1991					
Site 709	03°54.9' S	60°33.1' E	Mascarene Plateau	3041	<i>Globigerinoides sacculifer</i>	13.35-36.50	1-4	Shackleton and Hall, 1990					
					<i>Globigerinoides ruber</i>	26.30-53.96	3-5	Shackleton and Hall, 1990					
					<i>Globigerinoides sacculifer</i>	25.70-103.40	2-5	Vincent et al., 1991					
					<i>Dentoglobigerina altispira</i>	30.20-171.00	3-19	Vincent et al., 1991					
					<i>Globoquadrina venezuelana</i>	133.80-215.45	13-26	Vincent et al., 1991					
Site 714	05°03.6' S	73°47.2' E	Mascarene Plateau	2038	various species (Benthic)	54.2-200.5	5-24	Woodruff et al., 1990					
					<i>Globigerinoides sacculifer</i>	0.15-19.48	0-5	Droxler et al., 1990					
					<i>Globigerinoides sacculifer</i>	26.5-70.3	7-12	Boersma and Mikkelsen, 1990					
Site 716	04°56.0' S	73°17.0' E	Mascarene Plateau	544	various species (Benthic)	26.5-205.0	7-24	Boersma and Mikkelsen, 1990					
					<i>Globigerinoides sacculifer</i>	0.20-40.90	0-1	Droxler et al., 1990					
Site 738	62°42.54' S	82°47.25' E	Kerguelen Plateau	2253	various species (Planktonic)	23.00-377.23	35-65	Barrera and Huber, 1991					
					<i>Cibicides</i> spp.	18.42-33.86	35-37	Barrera and Huber, 1991					
					<i>Nuttallides truempyi</i>	61.00-380.70	39-66	Barrera and Huber, 1991					
					<i>Stilostomella subspinosa</i>	23.00-80.00	35-40	Barrera and Huber, 1991					
					<i>Stensioina beccariiiformis</i>	286.02-377.80	57-65	Barrera and Huber, 1991					
					Site 744	61°34.66' S	80°35.46' E	Kerguelen Plateau	2317	<i>Globigerinita juvenilis</i>	100.17-109.67	26-27	Barrera and Huber, 1991
										<i>Chiloguembelina cubensis</i>	118.2-175.07	31-39	Barrera and Huber, 1991
<i>Globorotaloides suteri</i>	82.63-175.07	19-39	Barrera and Huber, 1991										
<i>Cibicides</i> spp.	28.11-175.07	9-39	Barrera and Huber, 1991										
<i>Nuttallides</i> spp.	27.11-77.09	9-18	Barrera and Huber, 1991										
<i>Stilostomella subspinosa</i>	81.13-175.07	18-39	Barrera and Huber, 1991										
<i>Cibicides</i> spp.	38.05-63.55	10-15	Woodruff and Chamber, 1991										
Site 747	54°48.68' S	76°47.64' E	Central Kerguelen Plateau	1695	<i>Cibicides</i> spp.	55.24-57.75	13-14	Woodruff and Chamber, 1991					
					<i>Cibicides kullenbergi</i>	51.20-77.75	11-18	Woodruff and Chamber, 1991					
Site 748	58°26.45' S	78°58.89' E	Central Kerguelen Plateau	1290	<i>Cibicides</i> spp.B	50.23-52.70	10-11	Woodruff and Chamber, 1991					
					<i>Cibicides</i> spp.	33.50-141.90	8-26	Wright and Miller, 1992					
					<i>Subbotina angiporoides</i>	104.28-127.00	33-37	Zachos et al., 1992a					
					<i>Subbotina lin</i> spp. ?	171.90-147.90	40-44	Zachos et al., 1992a					

Table 1. (continued).

Site	Latitude	Longitude	Geographical area	Water depth (m)	Species	Range		Reference					
						mbsf	Age (Ma)						
Site 748	58°26.45' S	78°58.89' E	Central Kerguelen Plateau	1290	<i>Chiloguembelina cubensis</i>	104.50-176.00	33-45	Zachos et al., 1992a					
					<i>Chiloguembelina</i> spp.	154.28-389.22	42-59	Zachos et al., 1992a					
					<i>Cibicidoides</i>	67.57-350.80	22-58	Zachos et al., 1992a					
					<i>Gyroldina</i> spp.	67.57-389.22	22-59	Zachos et al., 1992a					
					<i>Stensioina beccariiiformis</i>	379.26-388.53	59	Zachos et al., 1992a					
Site 750	57°35.54' S	81°14.42' E	Kerguelen Plateau	2031	<i>Nuttallides umbonifera</i>	77.10-96.10	25-30	Zachos et al., 1992a					
					<i>Heterohelix globulosa</i>	349.73-356.28	66	Zachos et al., 1992b					
					<i>Eoblobigerina eobulloides</i>	349.73-349.93	66	Zachos et al., 1992b					
					<i>Globigerinelloides</i> spp.	349.73-350.25	66	Zachos et al., 1992b					
					<i>Nuttallides truempyi</i>	348.08-356.06	65-66	Zachos et al., 1992b					
Site 751	57°43.56' S	79°48.89' E	Kerguelen Plateau	1634	various species (Benthic)	40.28-165.2	4-19	Mackensen et al., in press					
					<i>Cibicidoides</i> spp.	0.80-102.05	0-35	Rea et al., 1991					
					Site 752	30°53.48' S	93°34.65' E	Broken Ridge	1086	<i>Cibicidoides</i> spp.	0.78-125.78	0-28	Rea et al., 1991
					Site 756	27°21.33' S	87°35.80' E	Ninetyeast Ridge	1518	<i>Gyroldinooides</i> spp.	0.80-140.20	0-37	Rea et al., 1991
Site 757	17°01.46' S	88°10.90' E	Ninetyeast Ridge	1652									
					Site 758	05°23.05' S	90°21.67' E	Ninetyeast Ridge	2935	<i>Uvigerina</i> spp.	2.13-123.90	0-37	Rea et al., 1991
Site 758	05°23.05' S	90°21.67' E	Ninetyeast Ridge	2935									
					<i>Dentoglobigerina altispira</i>	39.85-146.02	2-18	Vincent et al., 1991					
					<i>Globoquadrina venezuelana</i>	63.55-146.02	4-18	Vincent et al., 1991					
					<i>Globigerinoides sacculifer</i>	0.01-34.91	0-2	Farrell and Janecek, 1991					

data of interspecific differences previously compiled by many workers as well as in the present study.

In this study, the isotopic composition of various foraminiferal species is adjusted to that of *O. umbonatus*. This species was measured in the largest number of studied sites, occurs over a long time range, and has a stable magnitude of departure from the isotopic composition of the *Cibicidoides* group. The original δ values of *O. umbonatus* are converted into the δ values of dissolved CO₂ of bottom water by 0‰ for $\delta^{18}\text{O}$ and by +1.0‰ for $\delta^{13}\text{C}$ (Shackleton et al., 1984). For the Early to middle Paleogene, the isotopic composition is adjusted to that of *N. truempyi* and the δ values of *N. truempyi* are converted into the δ values of dissolved CO₂ of bottom water by +0.330‰ for $\delta^{18}\text{O}$ and by +1.082‰ for $\delta^{13}\text{C}$. If no measurements are given for *O. umbonatus*, the isotopic compositions of various species are adjusted to those of one taxon (*Cibicidoides* group, such as *C. wuellerstorfi*), and the data are converted to the δ values of dissolved CO₂ of bottom water. Adjusted values used in this study are shown in Table 2.

The δ values of planktonic foraminifers are adjusted to those of *Globigerinoides sacculifer* because this species has been measured at many sites and the ecology of this species is well known. An ecological investigation by plankton net reveals that *G. sacculifer* mostly inhabits water shallower than 50m (Be, 1977). Therefore, this study assumes the isotopic composition of *G. sacculifer* as an indication of the δ values of dissolved CO₂ of surface water shallower than 50 m depth. The adjustments of planktonic foraminifera are established by the same method as benthic foraminifera, and most of the adjustments are calculated based on the data of Shackleton et al. (1984). Some of the adjustments, however, are established by indirect differences, because the range of

planktonic foraminifer is often short. The isotopic compositions of Paleogene planktonic foraminifera are adjusted to those of *Subbotina* spp., which has been measured in many studies. In the *Subbotina* group, the intergeneric difference and the isotopic fluctuation according test size is relatively small (Shackleton et al., 1985; Stott et al., 1990). The δ values of *Subbotina* spp. in the Paleogene are converted into the δ values of *G. sacculifer* by -0.610‰ for $\delta^{18}\text{O}$ and by +0.850‰ for $\delta^{13}\text{C}$, which are indirectly obtained from the interspecific difference studied by Shackleton et al. (1984). The δ values adjusted from *Subbotina* spp., however, may not indicate that of dissolved CO₂ of surface water (50 m depth), because the δ values adjusted from *Subbotina* spp are >1‰ lower than those of the *Morozovella* group and the *Acarinina* group. The adjusted values of planktonic foraminifers are shown in Table 2.

2. Chronology of the samples

The ages of the studied section are defined by geomagnetic polarity events and the biostratigraphy of calcareous nannofossils. The numerical ages of nannofossil events have been originally determined by correlation with the geomagnetic polarity time scale.

The magnetic polarity time scale used in this study is based on Berggren et al. (1985a, 1985b, 1985c). The time scale of nannofossil species events has been proposed in many studies (Thiersten et al., 1977; Backman and Shackleton, 1983; Backman and Pestiaux, 1986; Berggren et al., 1985a; 1985b; 1985c; Backman, 1987; Gartner, 1977; Zijdeveld et al., 1986; Poore et al., 1983; Lohman, 1986; Barton and Bloemendal, 1986; Baldauf et al., 1987; Clement and Robinson, 1987; Takayama and Sato, 1987; Rahman and Roth, 1989; Rio et al., 1990; Gartner, 1990; Okada, 1990;

Table 2. Values to adjust to δ value of the bottom or surface (about 50m) DIC in the studied area.

Species	Site 752 Site 758 Site 754 Site 762 Site 756		Site 214 Site 238 Site 216 Site 237		Site 253		Site 516 Site 517 Site 518		Site 521 Site 519 Site 522 Site 523		Site 527 Site 525 Site 528 Site 526 Site 529		Site 667 Site 662 Site 658 Site 665 Site 659		
	$\delta^{18}\text{O}$	$\delta^{13}\text{C}$	$\delta^{18}\text{O}$	$\delta^{13}\text{C}$	$\delta^{18}\text{O}$	$\delta^{13}\text{C}$	$\delta^{18}\text{O}$	$\delta^{13}\text{C}$	$\delta^{18}\text{O}$	$\delta^{13}\text{C}$	$\delta^{18}\text{O}$	$\delta^{13}\text{C}$	$\delta^{18}\text{O}$	$\delta^{13}\text{C}$	
Benthic foraminifera (bottom water)															
<i>Anomalinoidea danicus</i>	0.178	0.547													
<i>Bulimina jarvisi</i>												-0.080	0.460		
<i>Bulimina</i> spp.												-0.140	0.445		
<i>Cassidulina cornuta</i>															
<i>Cibicides kullenbergi</i>									0.336	0.036		0.615	0.070		
<i>Cibicides lamondohertyi</i>															
<i>Cibicides mundulus</i>															
<i>Cibicides</i> sp.B															
<i>Cibicides</i> sp.C															
<i>Cibicides</i> spp.					0.396	0.210						0.280	0.290	0.280	0.290
<i>Cibicides wuellerstorfi</i>					0.490	0.240		0.208	-0.437	0.208	-0.437	0.795	-0.105	0.795	-0.105
<i>Gavilinkella</i> spp.												0.167	0.737		
<i>Globocassidulina</i> spp.												-0.230	0.441		
<i>Globocassidulina subglobosa</i>									-0.094	0.616					
<i>Cyroidina</i> spp.												0.110	0.280		
<i>Nuttallides</i> spp.												0.418	0.548		
<i>Nuttallides truempyi</i>	0.330	1.082													
<i>Nuttallides umbonifera</i>									0.147	-0.222					
<i>Oridorsalis umbonatus</i>	0.000	1.000	0.000	1.000	0.000	1.000			0.000	1.000		0.000	1.000		
<i>Planulina bradyi</i>															
<i>Planulina renzi</i>												0.870	-0.085		
<i>Rectuvigerina spinea</i>															
<i>Stensioina beccariformis</i>	0.214	1.203													
<i>Stilostomella</i> spp.												0.035	0.515		
<i>Uvigerina</i> spp.												-0.036	0.498		
Planktonic foraminifera (surface water)															
<i>Acarinina nitida</i>												-0.449	0.239		
<i>Acarinina primitiva</i>	-0.511	0.367													
<i>Acarinina</i> spp.												-0.580	0.052		
<i>Calapsydrax echinus</i>												-0.110	-0.133		
<i>Calapsydrax</i> spp.												-0.649	0.999		
<i>Calapsydrax unicus</i>												-0.822	1.244		
<i>Chiloguembelina wilcoxensis</i>												-0.439	1.317		
<i>Dentoglobigerina allispira</i>	-0.184	0.140	0.102	-0.060								-0.290	0.600		
<i>Globigerina angulicostata</i>												-0.229	0.849		
<i>Globigerina apertura</i>												0.080	0.285		
<i>Globigerina barbomensis</i>												-0.569	1.219		
<i>Globigerina carpudena</i>												-0.540	0.580		
<i>Globigerina euaperta</i>												-0.729	0.679		
<i>Globigerina globularis</i>												-0.689	1.159		
<i>Globigerina gortanii</i>												-1.220	0.696		
<i>Globigerina neperthes</i>												-0.221	0.349		
<i>Globigerina obesa</i>												-0.811	0.879		
<i>Globigerina osachensis</i>												-0.649	0.689		
<i>Globigerina pseudoampliapertura</i>												-0.754	0.944		
<i>Globigerina selli</i>												-0.719	0.579		
<i>Globigerina tripartita</i>												-0.529	0.859		
<i>Globigerina venezuelana</i>	-1.759	1.065	-1.115	1.027	-1.115	1.027									
<i>Globigerina winkleri</i>												-0.540	0.430		
<i>Globigerinoides ruber</i>									0.264	-0.011		-0.285	0.095	0.264	
<i>Globigerinoides sacculifer</i>	0.000	0.000			0.000	0.000						0.000	0.000		
<i>Globigerinoides seigliei</i>												-0.110	1.230		
<i>Globigerinoides subquadratus</i>												0.140	0.230		
<i>Globigerinoides trilobus</i>					-0.550	0.415									
<i>Globigerinatheka index</i>												-0.490	0.380		
<i>Globigerinatheka mexicana</i>												-0.309	0.266		
<i>Globigerinatheka</i> spp.												-0.355	0.450		
<i>Globoquadrina dehisrens</i>					-0.100	0.785						-0.100	0.785		
<i>Globoquadrina pradehisrens</i>												-0.679	0.579		
<i>Globoquadrina</i> spp.												-0.639	0.869		
<i>Globoquadrina transdehisrens</i>												-0.779	0.564		
<i>Globorotalia cerroazulensis</i>												-0.330	0.952		
<i>Globorotalia conoidea</i>												-0.571	0.679		
<i>Globorotalia inflata</i>												-1.465	0.198	-1.465	0.198
<i>Globorotalia peripheroronda</i>												-0.340	0.655		
<i>Globorotalia sikensis</i>												-0.659	0.999		
<i>Globorotalia truncatulinoides</i>												-1.335	0.285		
<i>Globorotaloides miozea</i>												-0.401	0.869		
<i>Globorotaloides suteri</i>															
<i>Haukenina</i> spp.												-0.466	0.840		
<i>Morozovella acuta</i>												0.408	-0.263		
<i>Morozovella aragonensis</i>												0.296	-0.860		
<i>Morozovella formosa</i>												-0.249	-0.331		
<i>Morozovella lehneri</i>												-0.059	-0.221		
<i>Morozovella marginodentata</i>	-0.464	-1.139													
<i>Morozovella pseudomenardii</i>												-0.652	1.485		
<i>Morozovella rex</i>												-0.034	-0.221		
<i>Morozovella subbotinae</i>												-0.371	0.251		
<i>Morozovella velascoensis</i>												-0.294	0.330		
<i>Planoglobulina</i> spp.												-0.234	-0.064		
<i>Subbotina</i> spp.	-0.610	0.850										-0.610	0.850		
<i>Tarborotalia increbescens</i>												-0.463	0.759		
<i>Tarborotalia</i> spp.												-0.594	0.899		

Backman *et al.*, 1990; and others). Among them, the biostratigraphic time scale proposed by Berggren *et al.* (1985a, 1985b) has been accepted as the standard time scale (e.g., Peirce, Weissel, *et al.*, 1989). Backman *et al.* (1990), however, pointed out that low-latitude nannofossil biochronology differs from the standard of mid-latitude biostratigraphy established by Berggren *et al.* (1985a, 1985b).

Based on their information the biochronology of nannofossil species events is constructed as follows; a) the basic biostratigraphic time scale follows that of Berggren *et al.* (1985a), b) nannofossil species events near the studied area are included if they are not contradictory in order, and c) species events described since Berggren *et al.* (1985a, 1985b) are considered. Thus, the biochronology applied in the time interval are: Pleistocene (Takayama and Sato, 1987), late Pliocene (Backman and Shackleton, 1983), early Pliocene to late Miocene (Rio *et al.*, 1990), middle Miocene to early Miocene (Backman *et al.*, 1990), Oligocene to middle Eocene (Okada, 1990), and the early Eocene to Paleocene (Berggren *et al.*, 1985a; 1985b; 1985c). In addition to the species events proposed by Martini (1971) and Okada and Bukry (1980), zonal schemes (Backman, J., Duncan, R. A., *et al.*, 1988; Peirce, Weissel, *et al.*, 1989; Rio *et al.*, 1990) are also considered. These compiled nannofossil events are shown in Table 3.

The numerical age of a sampled horizon is calculated by assuming a constant sedimentation rate between the two stratigraphic levels. The geomagnetic polarity events and nannofossil event levels at each site used in this study are shown in Tables 4 and 5, respectively.

3. Paleodepth

Paleodepths in this study were reconstructed by published data based on a "backtrack method" (Sites 214, 215, 216, 237, 238, 709, 752, 754, 756, and 758: Zachos *et al.*, 1992c; Sites 738 and 744: Barrera and Huber, 1991; Site 253: Kidd and Davies, 1978; Site 519: Finger, 1984; Sites 521, 522, and 523: Hsü *et al.*, 1984; Sites 525, 526, 527, 528, and 529: Moore, Jr. *et al.*, 1984; Sites 366 and 667: Miller *et al.*, 1989; Sites 698, 699, 700, and 702: Katz and Miller, 1991). The paleodepth at Site 748 through the Paleogene was upper bathyal (water depth: ~1000 m; Mackensen and Berggren, 1991). The paleodepth at Site 748 was calculated by assuming linear change from this depth to the present depth (1290 m). In case of an unknown subsidence curve, the present depth was used for the paleodepth for relatively young ages.

4. Paleolatitude and Paleolongitude

Paleolatitude and paleolongitude of the Indian Ocean site are estimated from migration velocity calculated from backtracked paleocoordinate (Zachos *et al.*, 1992c). Paleolatitude and paleolongitude of the South Atlantic Ocean site are estimated from migration velocity given by Scotese *et al.* (1988), and based on sea-floor spreading isochrons (Larson *et al.*, 1985). The paleolatitude and paleolongitude at each site are shown in Table 6.

B. General trend of isotopic records in the northern Indian Ocean

In the northern Indian Ocean, distribution pattern of carbon and oxygen isotopic ratios in surface (~50 m in depth) and bottom waters throughout the Cenozoic are illustrated in Figs. 36-39.

The oxygen isotopic records of bottom water around the Cretaceous / Tertiary boundary show a negative shift of 0.4‰ from -0.8 to -1.2‰. During the Paleocene, $\delta^{18}\text{O}$ values increase by 1.3‰ (66-61 Ma), and decrease down to a minimum value (-0.6‰) in the earliest Eocene (56 Ma). $\delta^{18}\text{O}$ values throughout the Eocene gradually increase by 1.6‰ from the minimum value in the earliest Eocene. From the Oligocene, distinct positive shifts are recognized three times. These shifts are observed immediately above epoch or subepoch boundaries, and accompanied by decreasing $\delta^{18}\text{O}$ values before the shifts. The first shift is recognized around the Eocene / Oligocene boundary, and the second shift in the middle Miocene. The net magnitudes of the first and second shifts are 0.7‰, 0.8‰, respectively. In the interval between the first and second shifts, the oxygen isotopic ratios vary from 1.0 to 2.5 ‰, with a gradual increase. Although a discontinuity is found near the Oligocene / Miocene boundary, this is probably caused by limited data. The third shift is recognized in the late Pliocene, and the net magnitude of this shift is 0.8‰. The oxygen isotopic ratios in interval between the second and third shifts are constant with a variation of 1.8-3.3‰. However, two remarkable ^{18}O maxima are recognized around 6 and 8 Ma. In the interval between the third shift and the present, $\delta^{18}\text{O}$ values show a gradual increase.

During the Paleocene, the increase of $\delta^{18}\text{O}$ values in surface water is observed up to 61 Ma, and subsequently a ^{18}O minimum value (-2.0‰) is recognized in the earliest Eocene (58 Ma). The same pattern of isotopic change is found in oxygen isotopic records of bottom water. The oxygen isotopic records of surface water, however, exhibit a smaller magnitude of change (only 0.5 ‰) than those of bottom waters. The difference of $\delta^{18}\text{O}$ value between Sites 752 and 758 is relatively large (about 0.8‰) at that time. The ^{18}O minimum of surface water during the earliest Eocene (58 Ma) delayed by about 2 Ma than that of bottom waters. From 61 to 37 Ma, the distribution of $\delta^{18}\text{O}$ values are parallel between surface and bottom water around 1.8‰. During the Eocene, $\delta^{18}\text{O}$ values increase by 2.5‰ (from -2.1 to 0.4‰) in the entire water column, with the increase of surface $\delta^{18}\text{O}$ values being especially rapid from the early to early middle Eocene. As a result, the $\delta^{18}\text{O}$ difference between surface and bottom water is reduced to about 0.6‰ after the early middle Eocene. From the Oligocene to early Miocene, $\delta^{18}\text{O}$ values are constant around -0.2‰. No remarkable shift at the Oligocene / Eocene boundary is recognized. Then, $\delta^{18}\text{O}$ values of surface water are again separated from those of bottom water at this time, and the difference is about 1.5‰ during the Oligocene. During the Miocene, $\delta^{18}\text{O}$ values decrease from 20 Ma to 16 Ma, increase to a peak value (~0.2‰) at 12 Ma, and again decrease to about -1.0‰. However, $\delta^{18}\text{O}$ values at Site 253 tend to increase from 0.5 to 1.5 ‰ at that time. Hence, the scatter of $\delta^{18}\text{O}$ value increases, and the highest degree of scatter (reaching 3‰) is recorded at 6 Ma. Around the Miocene / Pliocene boundary, $\delta^{18}\text{O}$ values at Site 253 rapidly decrease, and as a

Table 3. Cenozoic calcareous nannofossil datum levels and corresponding zonal boundaries of Okada and Bukry (1980) and Martini (1971) with age estimates.

Event	Species	Zone (base)		Age (Ma)	Reference	Event	Species	Zone (base)		Age (Ma)	Reference
		Bukry (1980)	Martini (1971)					Bukry (1980)	Martini (1971)		
Increase	<i>Emiliana huxleyi</i>			0.085	1	FO	<i>Discoaster druggii</i>	CN1c	NN2	23.6	10
LO	<i>Helicosphaera inversa</i>			0.15	2	LO	<i>Dictyococcites bisectus</i>	CN1a	NN1	23.7	3
FO	<i>Emiliana huxleyi</i>	CN15	NN21	0.275	1, 3	LO	<i>Sphenolithus ciperoensis</i>			23.7	3
LO	<i>Pseudoemiliana lacunosa</i>	CN14b	NN20	0.460	1, 3	LO	<i>Crassidiscus backmanii</i>			24.8	11
FO	<i>Helicosphaera inversa</i>			0.48	2	FO	<i>Crassidiscus backmanii</i>			25.0	11
Acme top	<i>Reticulofenestra</i> sp. A			0.83	2	LO	<i>Sphenolithus distentus</i>	CP19b	NP25	26.0	11
FO	<i>Gephyrocapsa parallela</i>			0.89	2	LO	<i>Chiasmolithus altus</i>			28.2	3
Increase	<i>Gephyrocapsa oceranica</i>			0.90	4	FO	<i>Sphenolithus ciperoensis</i>	CP19a	NP24	30.2	3
Acme top	<i>Gephyrocapsa</i> (small)			0.93	5	FO	<i>Sphenolithus distentus</i>	CP18		31.2	11
LO	<i>Gephyrocapsa</i> (large)			1.10	2	LO	<i>Sphenolithus aff. distentus</i>			32.4	11
LO	<i>Helicosphaera sellii</i>			1.19	2	LO	<i>Reticulofenestra umbilica</i>	CP17	NP23	34.2	11
FO	<i>Gephyrocapsa</i> (large)			1.36	2	LO	<i>Bramletteius serraculoides</i>			34.2	11
LO	<i>Calcidiscus macintyreii</i>			1.45	3, 6	LO	<i>Ericsonia obruta</i>			34.4	12
FO	<i>Gephyrocapsa oceanica</i>			1.57	2	LO	<i>Ericsonia formosa</i>	CP16c	NP22	34.9	12
FO	<i>Gephyrocapsa caribbeanica</i>			1.66	2	LO	<i>Isthmolithus recurvus</i>			34.9	3
LO	<i>Discoaster triradiatus</i>			1.89	7	LO	<i>Hayella suturiformis</i>			35.1	11
LO	<i>Discoaster brouweri</i>	CN13a	NN19	1.91	2	FO	<i>Sphenolithus aff. distentus</i>			35.1	11
Increase	<i>Discoaster triradiatus</i>			2.07	7	LO	<i>Chiasmolithus titus</i>			35.8	11
LO	<i>Discoaster pentaradiatus</i>	CN12d	NN18	2.35	6, 8	Acme top	<i>Ericsonia subdisticha</i>	CP16b		35.9	3
LO	<i>Discoaster surculus</i>	CN12c	NN17	2.41	6	LO	<i>Reticulofenestra oamaruensis</i>			36.0	13
LO	<i>Discoaster tamalis</i>	CN12b		2.65	6, 7	Increase	<i>Ericsonia obruta</i>			36.1	12
LO	<i>Discoaster variabilis</i>			2.90	5	LO	<i>Discoaster saipanensis</i>	CP16a	NP21	36.7	3, 12
LO	<i>Sphenolithus</i> spp.			3.45	6	LO	<i>Discoaster barbadiensis</i>			36.8	11
LO	<i>Sphenolithus neobabies</i>			3.51	8	LO	<i>Criocentrum reticulatum</i>			37.0	11
LO	<i>Sphenolithus abies</i>			3.56	2	FO	<i>Isthmolithus recurvus</i>	CP15b	NP19	37.8	3, 11
LO	<i>Reticulofenestra pseudoumbilica</i>	CN12a	NN16	3.56	6	FO	<i>Chiasmolithus oamaruensis</i>			39.8	3
FO	<i>Discoaster tamalis</i>	CN11b		3.8	3	LO	<i>Chiasmolithus grandis</i>	CP15a	NP18	40.0	3
LO	<i>Amaurolithus tricomiculatus</i>			3.7	3	LO	<i>Campylosphaera dela</i>			40.6	11
FO	<i>Pseudoemiliana lacunosa</i>			4.05	8	LO	<i>Sphenolithus spiniger</i>			41.4	11
FO	<i>Discoaster asymetricus</i>			4.1	3	LO	<i>Sphenolithus furcatolithoides</i>			41.4	11
LO	<i>Amaurolithus delicatus</i>			4.11	9	FO	<i>Dictyococcites bisectus</i>			41.4	11
LO	<i>Amaurolithus primus</i>	CN11a	NN15	4.37	9	FO	<i>Reticulofenestra reticulata</i>			42.1	13
LO	<i>Ceratolithus acutus</i>			4.43	9	LO	<i>Chiasmolithus solitus</i>	CP14b	NP17	43.4	11
FO	<i>Ceratolithus rugosus</i>	CN10c	NN13	4.66	9	LO	<i>Discoaster bifax</i>			43.4	11
FO	<i>Ceratolithus acutus</i>	CN10b		4.85	10	LO	<i>Nannotetrina alata</i>			43.4	11
LO	<i>Triquetrorhabdulus rugosus</i>			4.90	10	FO	<i>Reticulofenestra umbilica</i>			43.5	11
LO	<i>Ceratolithus armatus</i>			5.06	5	LO	<i>Cruciplacolithus staurion</i>			43.5	11
LO	<i>Discoaster quinquenarius</i>	CN10a	NN12	5.26	5	LO	<i>Nannotetrina fulgens</i>	CP14a		45.4	3
LO	<i>Amaurolithus amplificus</i>			5.33	9	FO	<i>Discoaster bifax</i>			46.6	11
LO	<i>Discoaster berggrenii</i>			5.80	5	LO	<i>Chiasmolithus gigas</i>	CP13c		47.0	3
FO	<i>Amaurolithus amplificus</i>			6.02	9	FO	<i>Sphenolithus furcatolithoides</i>			48.2	11
FO	<i>Amaurolithus primus</i>	CN9b		6.70	10	FO	<i>Chiasmolithus gigas</i>	CP13b		49.0	11
FO	<i>Discoaster quinquenarius</i>			7.46	10	FO	<i>Nannotetrina fulgens</i>	CP13a	NP15	49.8	3
FO	<i>Discoaster berggrenii</i>	CN9a	NN11	8.00	9	LO	<i>Discoaster lodoensis</i>			50.4	12
LO	<i>Discoaster neohamatus</i>			8.10	8	FO	<i>Reticulofenestra inflata</i>			52.0	3
FO	<i>Discoaster neorectus</i>	CN8b		8.5	3	FO	<i>Discoaster sublodoensis</i>	CP12	NP14	52.6	3
FO	<i>Discoaster loeblichii</i>			8.5	3	LO	<i>Tribracliatius orthostylus</i>	CP11	NP13	53.7	3
LO	<i>Discoaster hamatus</i>	CN8a	NN10	8.67	10	FO	<i>Discoaster lodoensis</i>	CP10	NP12	55.3	3
LO	<i>Catinaster</i> spp.			8.77	10	LO	<i>Tribracliatius contortus</i>	CP9b	NP11	56.3	3
FO	<i>Discoaster neohamatus</i>			8.96	10	FO	<i>Discoaster diastypus</i>			56.5	3
LO	<i>Catinaster coalitus</i>			9.00	5	FO	<i>Tribracliatius orthostylus</i>			56.6	3
FO	<i>Catinaster calyculus</i>			10.00	3	LO	<i>Fasciculolithus</i> spp.	CP9a	NP10	57.8	3
LO	<i>Coccolithus miopelagicus</i>			10.23	2	FO	<i>Tribracliatius bramlettei</i>			57.8	3
FO	<i>Discoaster hamatus</i>	CN7	NN9	10.5	10	FO	<i>Campylosphaera eodela</i>			58.2	3
FO	<i>Catinaster coalitus</i>	CN6	NN8	11.1	10	FO	<i>Discoaster multiradiatus</i>	CP8	NP9	59.2	3
FO	<i>Discoaster pentaradiatus</i>			12.0	3	FO	<i>Discoaster mobilis</i>			59.8	3
FO	<i>Discoaster kugleri</i>		NN7	12.2	10	FO	<i>Heliolithus reidellii</i>	CP7	NP8	60.0	3
LO	<i>Coronocyclas nitescens</i>			12.7	9	FO	<i>Discoaster mohleri</i>	CP6	NP7	60.4	3
LO	<i>Coccolithus floridanus</i>	CN5b		13.1	3, 10	FO	<i>Heliolithus kleinpellii</i>	CP5	NP6	61.6	3
LO	<i>Sphenolithus heteromorphus</i>	CN5a	NN6	13.6	10	FO	<i>Ellipsolithus tympaniformis</i>			62.0	3
FO	<i>Discoaster exilis</i>			15.4	3	FO	<i>Fasciculolithus</i> spp.	CP4	NP5	62.0	3
LO	<i>Helicosphaera ampliapertura</i>	CN4	NN5	16.0	3, 10	FO	<i>Ellipsolithus macellus</i>	CP3	NP4	63.7	3
Acme top	<i>Discoaster deflandrei</i>			16.1	9	FO	<i>Prinsius martinii</i>			63.8	3
FO	<i>Discoaster signus</i>			16.1	9	FO	<i>Chiasmolithus danicus</i>	CP2	NP3	64.8	3
FO	<i>Sphenolithus heteromorphus</i>	CN3		18.4	10	FO	<i>Cruciplacolithus tenuis</i>	CP1b	NP2	65.9	3
LO	<i>Sphenolithus belemnios</i>		NN4	18.8	10	FO	<i>Biantholithus sparsus</i>	CP1a	NP1	66.4	3
LO	<i>Triquetrorhabdulus carinatus</i>		NN3	19.5	10	FO	<i>Nephrolithus frequens</i>			69.0	
FO	<i>Sphenolithus belemnios</i>	CN2		20.0	10	LO	<i>Reticulofenestra levis</i>			69.0	

Note: FO = first occurrence, LO = last occurrence. The references refer to the age column and represent (1) Thierstein et al. (1977); (2) Takayama and Sato (1987); (3) Berggren et al. (1985a, 1985b, 1985c); (4) Gartner (1977); (5) Gartner (1990); (6) Backman and Shackleton (1983); (7) Backman and Pestioux (1986); (8) Rahman and Roth (1989); (9) Rio et al. (1990); (10) Backman et al. (1990); (11) Okada (1990); (12) Backman (1987); (13) Wei and Wise (1990).

Table 4. (continued).

Event	Species	Zone (base)			Age (Ma)	Site 750	Site 752	Site 754	Site 756	Site 757	Site 758	Site 762*
		Okada and Bukry (1980)	Martini (1971)									
FO	<i>Emiliania huyley</i>	CN15	NN21	0.275								
LO	<i>Pseudoemiliania lacunosa</i>	CN14b	NN20	0.460					2.3-3.8	6.0-7.5	4.4	
LO	<i>Helicosphaera sellii</i>			1.19						15.6-17.1		
FO	<i>Gephyrocapsa</i>			1.36								
LO	<i>Calcidiscus macintyre</i>			1.45					9.0-10.5			
FO	<i>Gephyrocapsa oceanica</i>			1.57								
LO	<i>Discoaster tetradiatus</i>			1.89								
LO	<i>Discoaster broweri</i>	CN13a	NN19	1.91		4.3	6.1-6.6	0.47-1.97	12.0-13.5	25.2-26.7	27.9	
Increase	<i>Discoaster tetradiatus</i>			2.07								
LO	<i>Discoaster pentaradiatus</i>	CN12d	NN18	2.35			6.6-8.1			33.3-34.8		
LO	<i>Discoaster surculus</i>	CN12c	NN17	2.41								35.9
LO	<i>Discoaster tamalis</i>	CN12b		2.65			9.6-11.1		15.5-17.0			
LO	<i>Discoaster variabilis</i>			2.90						37.8-39.3		
LO	<i>Sphenolithus</i>			3.45								
LO	<i>Reticulofenestra pseudoumbilica</i>	CN12a	NN16	3.56		10.3	15.7-16.2	4.50-6.00	23.0-23.6	44.4-45.9	64.4	
FO	<i>Discoaster tamalis</i>	CN11b		3.8			17.7-19.2	8.00-8.50	26.6-28.1	57.0-58.5		
LO	<i>Amaurolithus tricorniculatus</i>			3.7								
LO	<i>Amaurolithus primus</i>	CN11a	NN15	4.37								
FO	<i>Ceratolithus rugosus</i>	CN10c	NN13	4.66								92.9
FO	<i>Ceratolithus acutus</i>	CN10b		4.85		19.1	22.7-25.3		38.2-39.7			
LO	<i>Discoaster quinqueringus</i>	CN10a	NN12	5.26					42.8-44.3	69.7-71.2	99.4	
LO	<i>Discoaster berggrenii</i>			5.80								
FO	<i>Amaurolithus primus</i>	CN9b		6.70		25.1	31.8-34.9		52.5-54.0	83.1-84.6		
FO	<i>Discoaster quinqueringus</i>			7.46					62.2-63.6	102.4-103.9		
FO	<i>Discoaster berggrenii</i>	CN9a	NN11	8.00								111.9
FO	<i>Discoaster neorectus</i>	CN8b		8.5								
LO	<i>Discoaster hamatus</i>	CN8a	NN10	8.67				34.70-35.67	71.2-71.8	116.6-118.1		
FO	<i>Discoaster neohamatus</i>			8.96								
FO	<i>Discoaster hamatus</i>	CN7	NN9	10.5				37.30-38.80				118.4
FO	<i>Catiraster coaditus</i>	CN6	NN8	11.1								
FO	<i>Discoaster pentaradiatus</i>			12.0			46.5-48.5					
FO	<i>Discoaster kugleri</i>		NN7	12.2								
LO	<i>Coronocyclas nitescens</i>			12.7								
LO	<i>Coccolithus floridanus</i>	CN5b		13.1					80.9-81.5			
LO	<i>Sphenolithus heteromorphus</i>	CN5a	NN6	13.6		58.1	73.5-77.0	45.40-46.90	83.0-84.5		125.9	
FO	<i>Discoaster exilis</i>			15.4			80.0-83.2		89.0-90.5			
LO	<i>Helicosphaera ampliaperta</i>	CN4	NN5	16.0								
Aeme top	<i>Discoaster deflandrei</i>			16.1								
FO	<i>Discoaster signus</i>			16.1								
FO	<i>Sphenolithus heteromorphus</i>	CN3		18.4		73.1	86.7-89.7	54.70-56.20	95.7-97.2			
LO	<i>Sphenolithus belemnos</i>		NN4	18.8						144.0-145.5	133.5	
LO	<i>Triquetrorhabdulus carinatus</i>		NN3	19.5								
FO	<i>Sphenolithus belemnos</i>	CN2		20.0			89.7-92.9		98.7-100.2			
FO	<i>Discoaster druggii</i>	CN1e	NN2	23.6					100.2-100.8			
LO	<i>Dicryococites bisectus</i>	CN1a	NN1	23.7		91.7	107.1-108.6					144.9
LO	<i>Sphenolithus ciperoensis</i>			23.7				73.70-74.70	100.8-102.3	195.3-196.8		
LO	<i>Sphenolithus distentus</i>	CP19b	NP25	26.0								
LO	<i>Chiasmolithus albus</i>			28.2		93.4	116.8-113.8					
FO	<i>Sphenolithus ciperoensis</i>	CP19a	NP24	30.2				94.20-96.10	105.3-106.8	218.3-219.8	152.9	
FO	<i>Sphenolithus distentus</i>	CP18		31.2					110.5-112.0	237.7-239.2		
LO	<i>Reticulofenestra umbilica</i>	CP17	NP23	34.2				117.70-119.20	115.0-116.5			162.4
LO	<i>Ericsonia formosa</i>	CP16c	NP22	34.9					119.5-120.1			173.4
LO	<i>Isthmolithus recurvus</i>			34.9								
LO	<i>Reticulofenestra oamaruensis</i>			36.0								
LO	<i>Discoaster saipanensis</i>	CP16a	NP21	36.7				134.90-136.40	123.1-124.6			200
FO	<i>Isthmolithus recurvus</i>	CP15b	NP19	37.8		104.3			129.1-129.8			222
FO	<i>Chiasmolithus oamaruensis</i>			39.8								
LO	<i>Chiasmolithus grandis</i>	CP15a	NP18	40.0					131.8-133.3			246
FO	<i>Reticulofenestra reticulata</i>			42.1								
LO	<i>Chiasmolithus solitus</i>	CP14b	NP17	43.4								272.5
FO	<i>Reticulofenestra umbilica</i>			43.5								
LO	<i>Nannotetrina fulgens</i>	CP14a		45.4					153.6-155.1			
LO	<i>Chiasmolithus gigas</i>	CP13c		47.0					158.4-159.9			
FO	<i>Chiasmolithus gigas</i>	CP13b		49.0					166.0-167.5			
FO	<i>Nannotetrina fulgens</i>	CP13a	NP15	49.8					174.7-175.3			307.5
FO	<i>Discoaster subloedenis</i>	CP12	NP14	52.6					185.7-187.2			329.5
LO	<i>Tribrachiatum orthostylus</i>	CP11	NP13	53.7					202.0-203.5			334.5
FO	<i>Discoaster lodoensis</i>	CP10	NP12	55.3		114.3			211.7-213.2			369.5
LO	<i>Tribrachiatum costatum</i>	CP9b	NP11	56.3								393
FO	<i>Discoaster diastypus</i>			56.5								
FO	<i>Tribrachiatum orthostylus</i>			56.6								
LO	<i>Fasciculithus</i>	CP9a	NP10	57.8								407
FO	<i>Tribrachiatum bramlettei</i>			57.8								
FO	<i>Discoaster multiradiatus</i>	CP8	NP9	59.2				202.05		264.1-265.4		421.5
FO	<i>Heliolithus reidelii</i>	CP7	NP8	60.0				219.5				
FO	<i>Discoaster mohleri</i>	CP6	NP7	60.4						271.1-271.6		434
FO	<i>Heliolithus klebyellii</i>	CP5	NP6	61.6				251.4				459.5
FO	<i>Ellipsolithus tympaniformis</i>			62.0	309.55-309.77	297.7				290.4-291.6		
FO	<i>Fasciculithus</i>	CP4	NP5	62.0								480
FO	<i>Ellipsolithus macellus</i>	CP3	NP4	63.7								489.5
FO	<i>Prabus marshallii</i>			63.8				318.7				
FO	<i>Chiasmolithus danicus</i>	CP2	NP3	64.8	346.42-346.44	345.1				292.1-295.6		529
FO	<i>Cruciplacolithus tenuis</i>	CP1b	NP2	65.9				353.46				552.5
FO	<i>Biantholithus sparsus</i>	CP1a	NP1	66.4	350.0	358.5						556
FO	<i>Nephrolithus frequens</i>			69.0						295.8-295.9		
LO	<i>Reticulofenestra levins</i>			69.0	394.6-394.4							

Table 6. Paleolatitude and paleolongitude for Indian and South Atlantic Ocean DSDP and ODP sites.

Age (Ma) Latitude Longitude			Age (Ma) Latitude Longitude			Age (Ma) Latitude Longitude			Age (Ma) Latitude Longitude		
Site 214			Site 750			Site 689			Site 529		
0	-11.3368	88.7180	0	-57.5920	81.2395	0.0	-64.5167	3.1000	0.0	-28.9305	2.7680
10	-15.9456	84.8116	10	-57.3442	78.9036	5.9	-66.2000	4.9456	5.9	-31.8078	2.7354
20	-20.5250	80.8046	20	-57.7088	78.1143	23.0	-69.0000	6.9186	23.0	-35.9078	1.9362
35	-27.3647	75.6525	35	-58.4260	77.3241	37.7	-71.0000	11.9846	37.7	-39.1078	-0.0243
45	-30.3489	71.2536	45	-57.5245	75.5062	59.2	-67.1000	5.1890	59.2	-37.7078	-1.0619
55	-37.6345	66.5541	55	-56.2898	73.9504	66.2	-67.2000	3.4593	66.2	-39.7078	-2.7915
60	-43.5605	63.1810	60	-56.5576	73.2808						
66	-53.9952	55.9089	66	-58.2030	72.1747	Site 690			Site 516		
						0.0	-65.1667	1.2000	0.0	-30.2767	-35.6183
						5.9	-66.6000	3.3355	5.9	-31.8000	-34.0447
Site 216			Site 751			23.0	-69.2000	4.9584			
0	1.4622	90.2080	0	-57.7260	79.8148	37.7	-71.5000	10.3713	Site 517		
10	-3.2994	87.3183	10	-57.4782	77.4789	59.2	-67.5000	4.1512	0.0	-30.9467	-38.0417
20	-7.9936	83.9027	20	-57.8428	76.6896	66.2	-67.4000	1.6143	5.9	-32.7000	-36.6901
35	-15.1726	80.1364	35	-58.5600	75.8994						
45	-18.3758	76.4942	45	-57.6585	74.0815	Site 698			Site 518		
55	-25.8665	72.7479	55	-56.4238	72.5257	0.0	-51.4585	-33.0993	0.0	-29.9733	-38.1350
60	-31.9171	70.1932	60	-56.6916	71.8561	37.7	-58.0000	18.5531	5.9	-31.7000	-36.4601
66	-42.8533	65.7410	66	-58.3370	70.7500	59.2	-56.1000	9.2249			
						66.2	-57.1000	7.6105	Site 519		
Site 237			Site 752						0.0	-26.1367	-11.6662
0	-7.0832	58.1247	0	-30.8913	93.5775	Site 699			5.9	-28.2000	-12.3067
10	-8.1770	56.7058	10	-35.8046	88.0336	0.0	-51.5423	-30.6770			
20	-10.0566	55.1970	20	-40.4972	83.0759	37.7	-58.0000	16.1332	Site 521		
35	-13.5454	53.8467	35	-47.4277	75.2207	59.2	-56.1000	6.8033	0.0	-26.0738	-10.2645
45	-14.8039	52.7407	45	-49.7730	69.4555	66.2	-56.9000	4.8430	5.9	-28.1000	-10.3514
55	-15.9938	52.3575	55	-49.0639	67.5327						
60	-17.3310	51.9533	60	-49.4415	67.3313	Site 700			Site 522		
66	-22.4280	49.7591	66	-51.1781	67.0640	0.0	-51.5330	-30.2781	0.0	-26.1140	-5.1130
Site 238			Site 754			37.7	-58.0000	15.3265	5.9	-28.1000	-5.6358
0	-11.1535	70.5260	0	-30.9407	93.5658	59.2	-56.1000	6.2268	23.0	-33.3000	-6.6880
10	-14.1274	66.4282	10	-35.8529	88.0160	66.2	-56.9000	4.1512	37.7	-35.0000	8.0666
20	-17.6872	62.1526	20	-40.5448	83.0540						
35	-22.4291	56.7264	35	-47.4727	75.1873	Site 702			Site 523		
45	-23.8565	52.5880	45	-49.8163	69.4157	0.0	-50.9464	-26.3686	0.0	-28.5522	-2.2513
			55	-49.1061	67.4905	37.7	-57.7000	12.0999	5.9	-30.9000	-1.8403
			60	-49.4835	67.2886	59.2	-55.5000	2.6521	23.0	-34.7000	-3.4593
			66	-51.2198	67.0189	66.2	-56.7000	1.6143	37.7	-37.3000	-4.6095
									59.2	-36.4000	-5.7655
Site 709			Site 756			Site 704			Site 658		
0	-3.9120	60.5517	0	-27.3548	87.5973	0.0	-46.8800	7.4217	0.0	20.7492	-18.5812
10	-4.8220	58.9752	10	-31.8147	82.1515	5.9	-48.9000	7.8211	5.9	-19.0000	-18.7476
20	-6.4862	57.2805	20	-36.2513	77.1224	23.0	-53.6000	6.9186			
35	-9.6658	55.6776	35	-42.6070	69.4599	Site 525			Site 366		
45	-10.7405	54.4558	45	-45.2435	63.6150	0.0	-29.0712	2.9853	0.0	5.6783	-19.8517
55	-11.7185	53.9289	55	-52.0742	56.5491	5.9	-31.9485	2.9528	5.9	3.9000	-22.0831
60	-12.9602	53.4716	60	-57.6124	50.7719	23.0	-36.0485	2.1535	23.0	0.4000	-19.6028
66	-15.7036	52.4378	66	-66.4441	34.9935	37.7	-39.2485	0.1930	37.7	-2.1000	-19.8207
						59.2	-37.8485	-0.8445			
						66.2	-39.8485	-2.5742	Site 667		
Site 738			Site 757						0.0	4.5692	-21.9113
0	-62.7092	82.7875	0	-17.0233	88.1802	Site 526			5.9	1.0000	-19.7827
10	-62.4770	80.2168	10	-21.5704	83.7710	0.0	-29.1227	3.1380	23.0	-1.8000	-22.2550
20	-62.8313	79.5958	20	-26.0973	79.4448	5.9	-32.0000	3.1054	37.7	-3.9000	-22.4712
35	-63.5433	78.9116	35	-32.7713	73.5277	23.0	-36.1000	2.3062			
45	-62.6609	76.7674	45	-35.6424	68.6980	37.7	-39.3000	0.3457			
55	-61.4366	74.9877	55	-42.7977	63.3564	59.2	-37.9000	-0.6919			
60	-61.6973	74.4524	60	-48.6259	59.3740	66.2	-39.9000	-2.4215			
66	-63.3389	73.4704	66	-58.6726	49.9854						
Site 744			Site 758			Site 527					
0	-61.5777	80.5910	0	5.3842	90.3612	0.0	-28.0415	1.7633			
10	-61.3221	78.0938	10	0.6002	87.7646	5.9	-30.9188	1.7308			
20	-61.6922	77.4088	20	-4.1143	84.5032	23.0	-35.0188	0.9315			
35	-62.4115	76.6534	35	-11.3592	81.0910	37.7	-38.2188	-1.0290			
45	-61.4999	74.6470	45	-14.6066	77.6362	59.2	-36.8188	-2.0665			
55	-60.2574	72.9944	55	-22.1319	74.0700	66.2	-38.8188	-3.7962			
60	-60.5309	72.4101	60	-28.1985	71.6449						
66	-62.1785	71.2999	66	-39.2081	67.6717	Site 528					
						0.0	-28.5248	2.3240			
						5.9	-31.4022	2.2914			
Site 748			Site 762			23.0	-35.5022	1.4922			
0	-58.4408	78.9815	0	-19.8872	112.2540	37.7	-38.7022	-0.4683			
10	-58.1638	76.6303	10	-25.8986	108.4818	59.2	-37.3022	-1.5059			
20	-58.5497	75.8385	20	-31.0841	105.0744	66.2	-39.3022	-3.2355			
35	-59.2735	75.0103	35	-39.5450	100.8771						
45	-58.3423	73.2131	45	-42.8307	96.9645						
55	-57.0886	71.7107	55	-42.9795	95.3122						
60	-57.3704	71.0416	60	-43.3728	95.3188						
66	-59.0207	69.8372	66	-45.1141	96.0148						

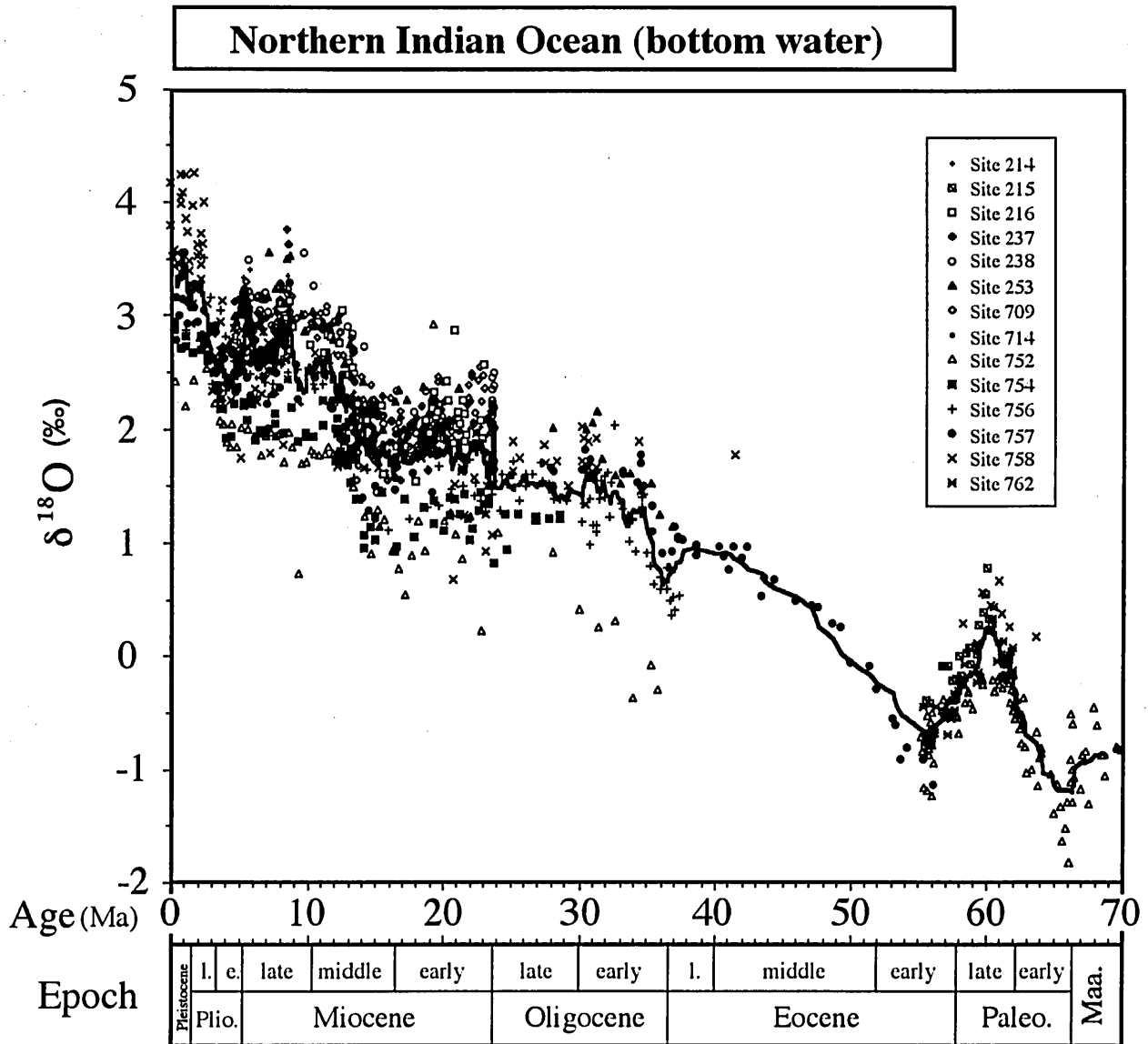


Fig. 36. Composite oxygen isotope record of bottom water in the Northern Indian Ocean region. All data corrected to dissolved inorganic carbon (DIC) of bottom water (See III-A-1 section). The smoothed curve is obtained by 15-point running average.

consequence there is less scatter in $\delta^{18}\text{O}$ values. Afterward, the general trends of surface $\delta^{18}\text{O}$ value are close to constant, although they exhibit a large fluctuation ($\sim 1.8\text{‰}$). From the Miocene, the general trend of surface oxygen isotopic record shows the opposite pattern to that of bottom water, and consequently the $\delta^{18}\text{O}$ difference between surface and bottom water is expanded. Near the present-day, this difference reaches 4.5‰ .

The carbon isotopic values of bottom water tend to increase about 2.5‰ from 70 to 65 Ma. Across the K/T boundary, no remarkable change in $\delta^{13}\text{C}$ values is found, however they fluctuate. Around the late / early Paleocene boundary (63 Ma), $\delta^{13}\text{C}$ shows a minimum value ($\sim -1.8\text{‰}$).

The carbon isotopic ratios rapidly increase from this minimum value to a maximum ($\sim 3.0\text{‰}$) at 60 Ma within the late Paleocene, which shows the highest values among the Cenozoic data. Across the Paleocene / Eocene boundary, they exhibit a remarkable negative shift by 2.5‰ to the minimum value zone in the earliest Eocene (56-54 Ma). $\delta^{13}\text{C}$ values in the minimum zone are largely fluctuated, and are around 0.6‰ . $\delta^{13}\text{C}$ values increase by 1.2‰ from 55 to 53 Ma and remain around 1.6‰ from the late early to late Eocene (53-37 Ma). Around the Oligocene / Eocene boundary, they increase slightly to 1.8‰ . From the Oligocene, two remarkable carbon isotopic negative shifts are observed in the earliest Oligocene (33 Ma) and the latest Miocene (6 Ma, Chron-6

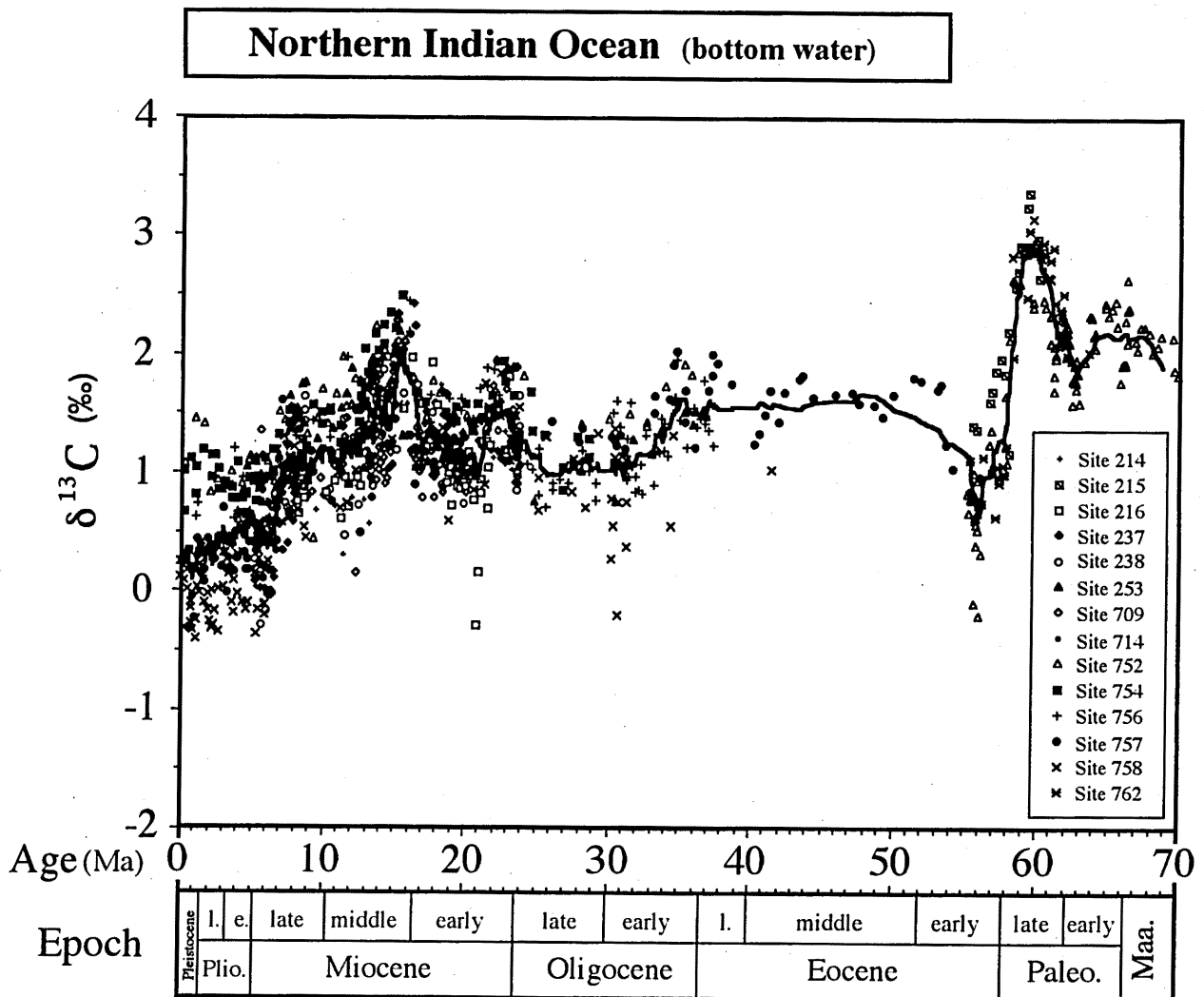


Fig. 37. Composite carbon isotope record of bottom water in the Northern Indian Ocean region. All data corrected to DIC of bottom water (See III-A-1 section). The smoothed curve is obtained by 15-point running average.

Shift: Vincent *et al.*, 1985). The magnitudes of these shifts are $\sim 0.7\text{‰}$ and $\sim 0.8\text{‰}$, respectively. During the interval between these two shifts, $\delta^{13}\text{C}$ values are constant at $\sim 1.0\text{‰}$ with a range of 1.0‰ . In this interval, pronounced peaks are recognized in the earliest Miocene (23 Ma) and the middle Miocene (15 Ma, Monterey Excursion: Vincent and Berger, 1985). $\delta^{13}\text{C}$ ratios begin to increase from the Oligocene / Miocene boundary reaching a maximum value ($\sim 1.5\text{‰}$) at 23 Ma, and subsequently decreasing at 20 Ma. It ratios increase again from 17 Ma, reaching a maximum value ($\sim 2.2\text{‰}$) at 15 Ma, and decreasing to 1.0‰ at 12 Ma (the peak of the middle Miocene). The magnitude of the middle Miocene peak is higher than that of the earliest Miocene. After the shift in the latest Miocene, $\delta^{13}\text{C}$ values are generally constant around 0.5‰ , but the degree of scatter is relatively high ($\sim 1.5\text{‰}$).

The carbon isotopic ratios of surface water vary in a similar pattern to that of bottom waters during the Paleocene and Eocene. During the Paleocene, they are higher by $\sim 0.5\text{‰}$

than those of the bottom water. $\delta^{13}\text{C}$ differences between surface and bottom water increases to 0.8‰ at the earliest Eocene. From the middle Eocene to Oligocene, $\delta^{13}\text{C}$ values remain around 2.5‰ without any remarkable shifts. Consequently, the surface to bottom $\delta^{13}\text{C}$ difference expands to 1.1‰ in the late Oligocene. $\delta^{13}\text{C}$ values from the Oligocene / Miocene boundary to the early Miocene (20 Ma) decrease to a minimum value ($\sim 1.5\text{‰}$). This minimum value is close to the present value, and at this time, the $\delta^{13}\text{C}$ difference between surface to bottom water is small ($\sim 0.4\text{‰}$). During the Miocene, pronounced peaks are recognized at 15 and 8 Ma. For the peaks at 15 and 8 Ma, the values are $\sim 2.7\text{‰}$ and $\sim 2.3\text{‰}$, respectively. The peak at 15 Ma is remarkable in the carbon isotopic record of bottom water, whereas the peak at 8 Ma is weak, and consequently the surface to bottom difference increases more than 1.0‰ at this time. Although $\delta^{13}\text{C}$ values in surface water show a negative

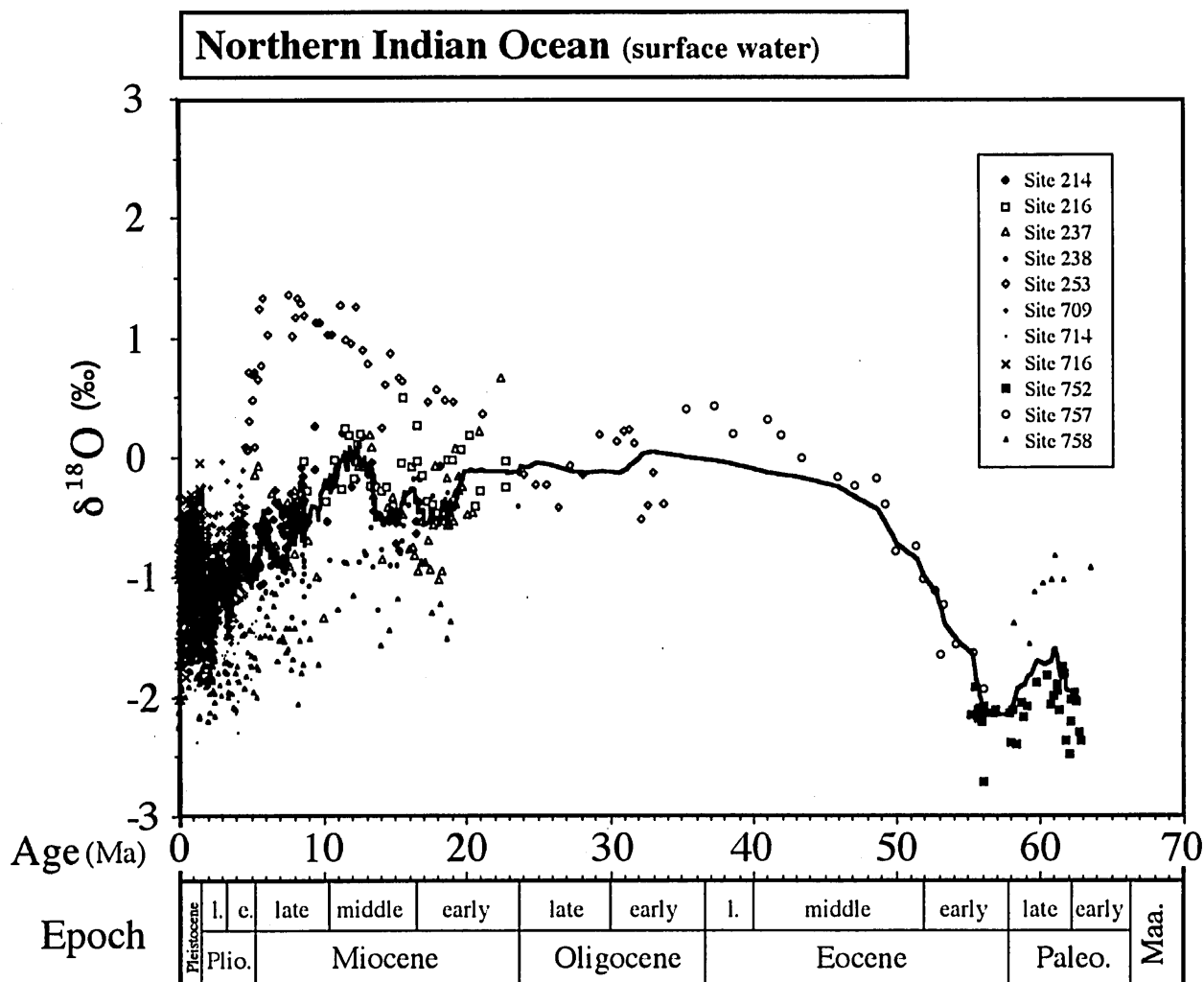


Fig. 38. Composite oxygen isotope record of surface water in the Northern Indian Ocean region. All data corrected to DIC of surface water (about 50 m below sea; See III-A-1 section). The smoothed curve is obtained by 15-point running average.

shift of 0.5‰ at ~6 Ma, the magnitude of this shift is smaller than that of the bottom waters. After the shift at 6 Ma, $\delta^{13}\text{C}$ values are constant around 1.4‰ with a low degree of scatter (~0.8‰). In this section, a weak peak (~0.3‰) is observed at 3 Ma. The carbon isotopic ratios of surface water are about 1.0‰ higher than those of bottom water.

C. Correlation of oxygen and carbon isotopes between the Indian Ocean and the South Atlantic Ocean

The isotopic records of the northern Indian Ocean are correlated with those of the Indian and South Atlantic Oceans. The general trends of carbon and oxygen isotopes in these ocean through the Cenozoic are illustrated in Figs. 40-53.

Oxygen isotopic record: During the late Maastrichtian, $\delta^{18}\text{O}$ values of bottom water in the Southern Ocean (Atlantic sector) are ~1.7‰ higher than those of the northern Indian Ocean. From 66 to 61 Ma within the Paleocene, although

$\delta^{18}\text{O}$ values in the northern Indian Ocean rapidly increase, the increases in the South Atlantic and southern Indian Oceans are gradual. Consequently, the difference of $\delta^{18}\text{O}$ values between the northern Indian Ocean and the other oceans decreases to 0.4‰ in the $\delta^{18}\text{O}$ maximum value at 61 Ma. Through the late Paleocene to Eocene, the same trends as those of the northern Indian Ocean are observed in all oceans. The $\delta^{18}\text{O}$ difference throughout the ocean is small. However, $\delta^{18}\text{O}$ values in the South Atlantic Ocean tend to be slightly higher than those of the other ocean. In the shift at the Oligocene / Eocene boundary, the net magnitude in the northern Indian Ocean is remarkably small (~0.6‰). The southern Indian Ocean exhibits a shift of ~0.8‰, the South Atlantic Ocean is ~0.6‰, the Southern Ocean (Atlantic sector) is ~1.0‰, and the Central Atlantic Ocean is ~1.0‰. This shift is generally large in the Antarctic Ocean, and tends to decrease northwards at this time. During the Oligocene, different trends in oxygen isotopes of bottom water are

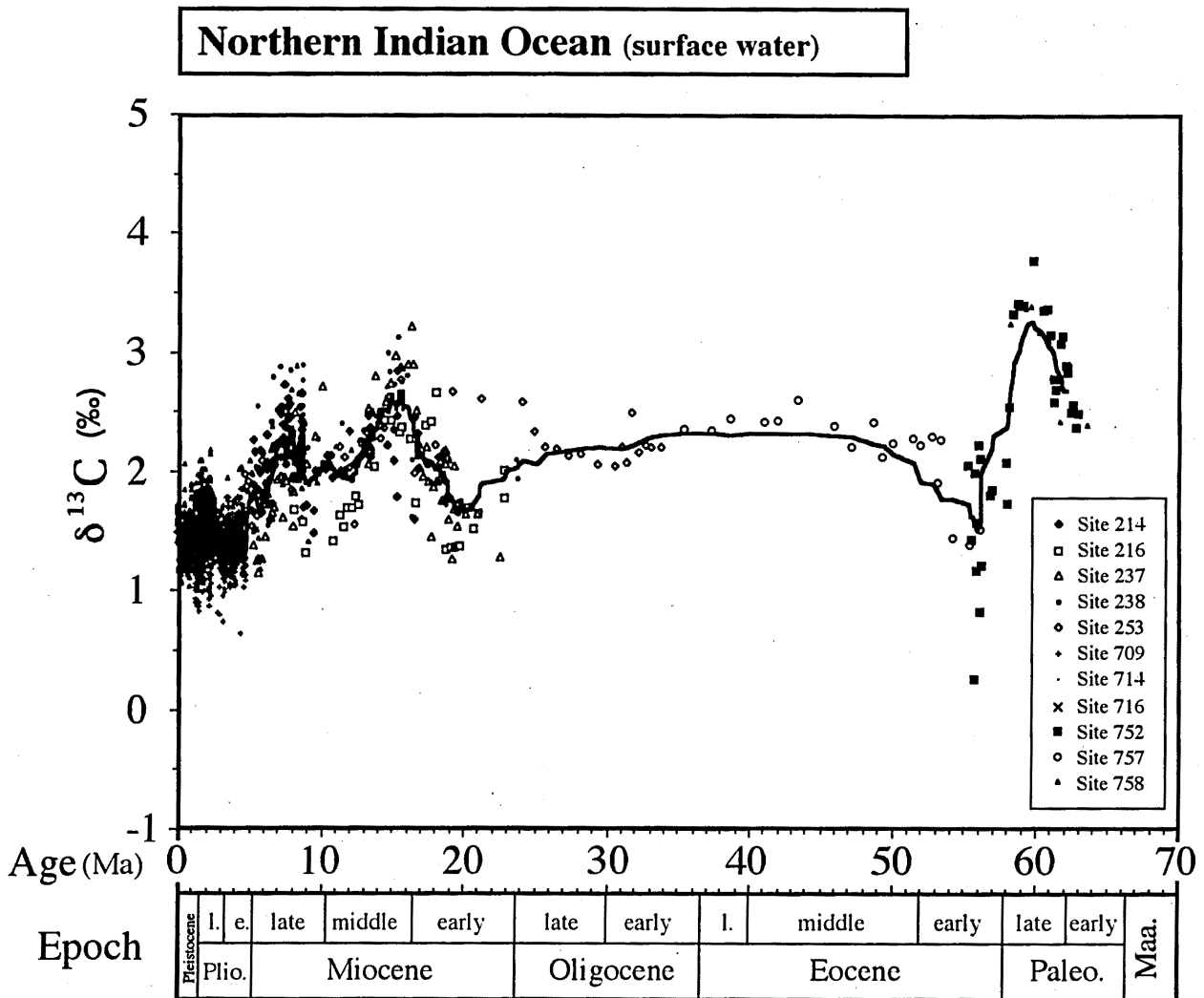


Fig. 39. Composite carbon isotope record of surface water in the Northern Indian Ocean region. All data corrected to DIC of surface water (about 50 m below sea; See III-A-1 section). The smoothed curve is obtained by 15-point running average.

recognized in each ocean. $\delta^{18}\text{O}$ values in the Southern Ocean (Atlantic sector) increase. In the northern Indian and Central Atlantic Oceans, they gradually increase, while they gradually decrease in the southern Indian and South Atlantic Oceans. Then, $\delta^{18}\text{O}$ values in the northern Indian, South and Central Atlantic Oceans become $\sim 0.7\text{‰}$, which is lower than those of the southern Indian and Southern Oceans. The shift in the middle Miocene is largest in the southern Indian Ocean ($\sim 0.9\text{‰}$). The net magnitudes of this shift in the northern Indian and South Atlantic Oceans are $\sim 0.8\text{‰}$. In the Central Atlantic Ocean, the shift is remarkable small ($\sim 0.6\text{‰}$). The decrease in $\delta^{18}\text{O}$ value before the shift is clearly observed in the southern Indian Ocean. The shift in the late Pliocene is most conspicuous in the Central Atlantic Ocean, and then the magnitude of shift reaches 1.3‰ . The net magnitudes in other oceans are 1.1 to 0.6‰ , and the northern Indian Ocean records the lowest magnitude. The magnitude of this shift tends to decrease southwards in the Atlantic Ocean side, and

toward the Indian Ocean. Thus, $\delta^{18}\text{O}$ values appear to be highest in the Central Atlantic Ocean from the late Pliocene onwards.

From the Paleocene to Eocene, the oxygen isotopic records in surface water exhibit a similar pattern in all oceans. $\delta^{18}\text{O}$ values during the Paleocene, however, are different in each region. $\delta^{18}\text{O}$ values in the northern Indian Ocean are smallest at $\sim 2.0\text{‰}$. The Southern Ocean and South Atlantic Oceans are $\sim 1.5\text{‰}$, which is larger than that of the northern Indian Ocean. In the South Atlantic Ocean, $\delta^{18}\text{O}$ values are larger at $\sim 0.8\text{‰}$ compared with the northern Indian Ocean. The magnitude of increase during the early Eocene is large in the northern Indian Ocean, and similar $\delta^{18}\text{O}$ values are recognized in the Indian and South Atlantic Oceans during the middle and late Eocene. In this interval, the surface to bottom $\delta^{18}\text{O}$ difference is also small ($< 1\text{‰}$). A sharp increase at the Oligocene / Eocene boundary is recognized in the southern

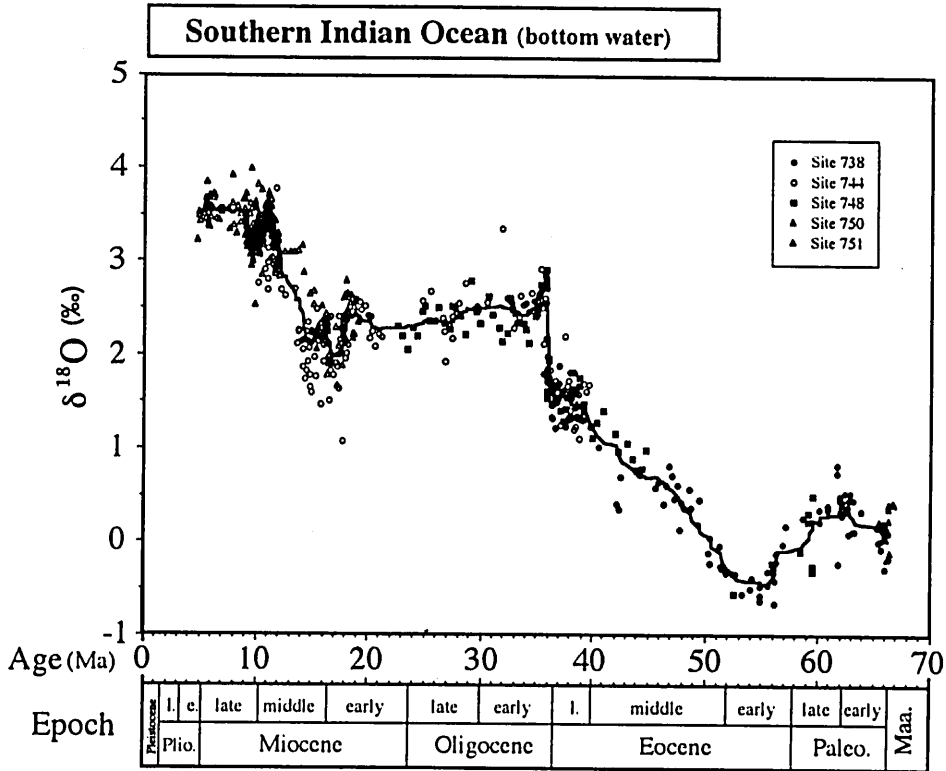


Fig. 40. Composite oxygen isotope record of bottom water in the Southern Indian Ocean region. All data corrected to DIC of bottom water (See III-A-1 section). The smoothed curve is obtained by 15-point running average.

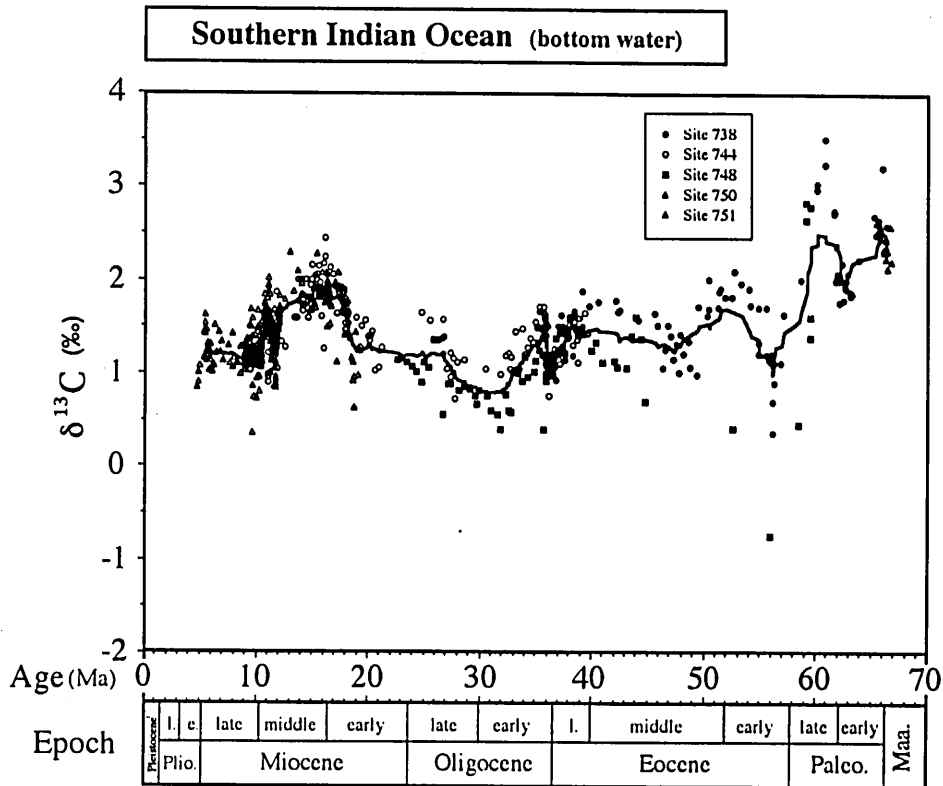


Fig. 41. Composite carbon isotope record of bottom water in the Southern Indian Ocean region. All data corrected to DIC of bottom water (See III-A-1 section). The smoothed curve is obtained by 15-point running average.

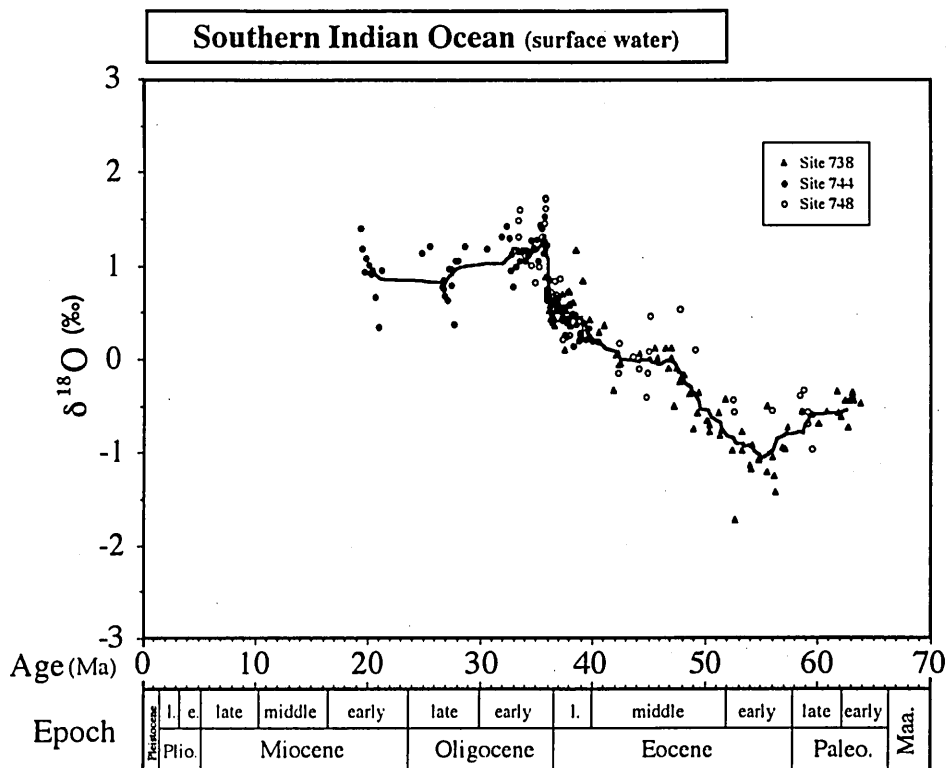


Fig. 42. Composite oxygen isotope record of surface water in the Southern Indian Ocean region. All data corrected to DIC of surface water (about 50 m below sea; See III-A-1 section). The smoothed curve is obtained by 15-point running average.

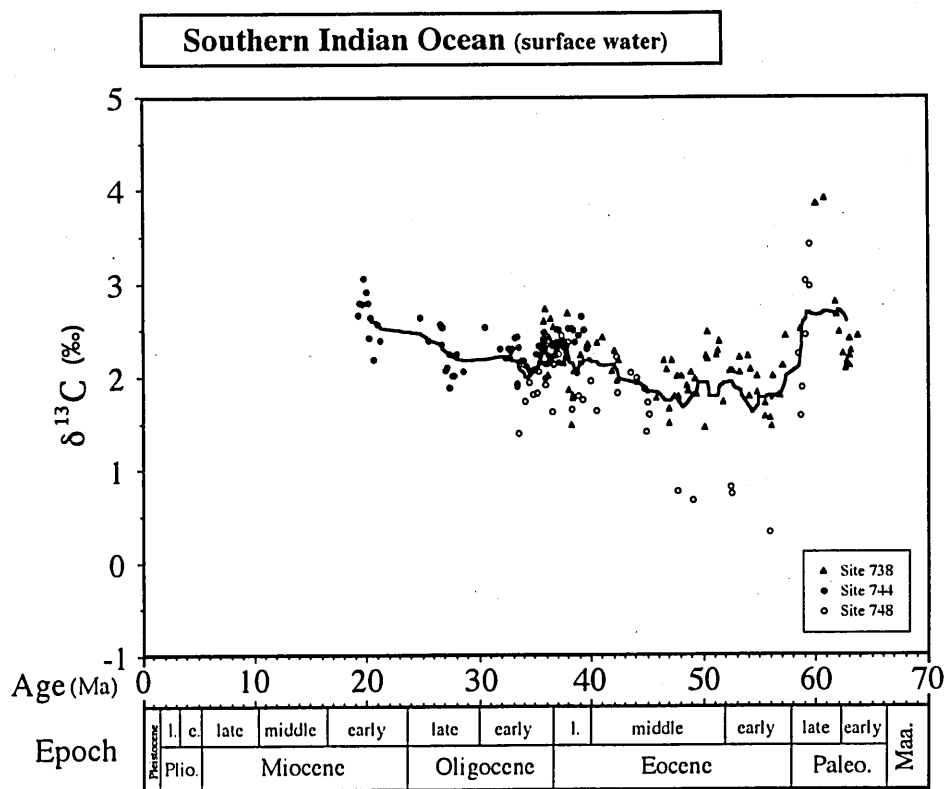


Fig. 43. Composite carbon isotope record of surface water in the Southern Indian Ocean region. All data corrected to DIC of surface water (about 50 m below sea; See III-A-1 section). The smoothed curve is obtained by 15-point running average.

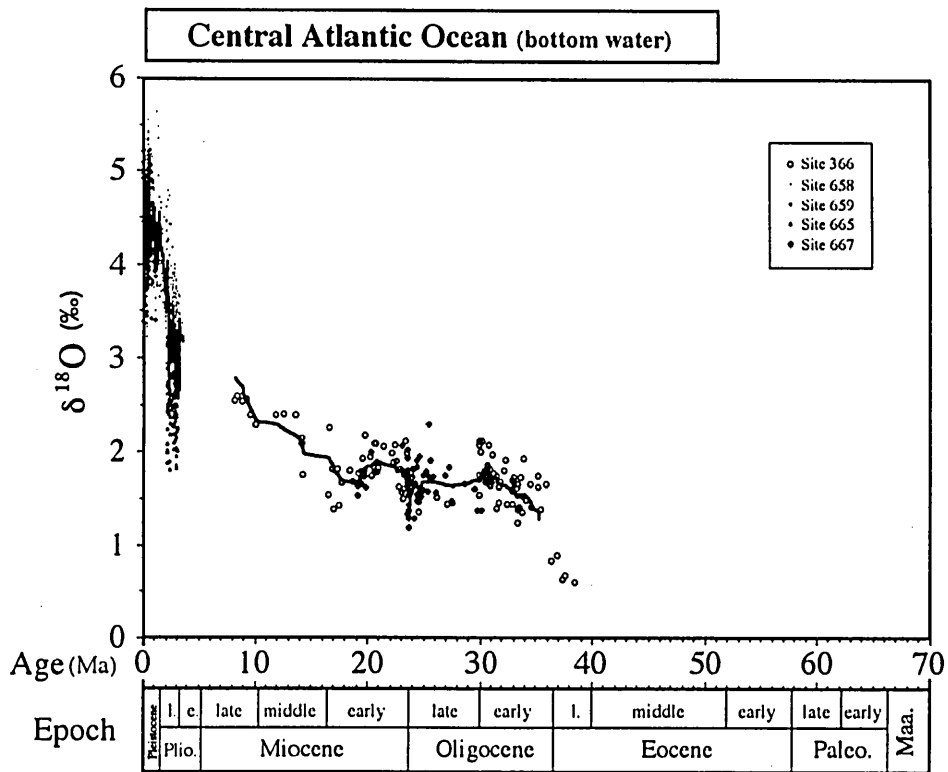


Fig. 44. Composite oxygen isotope record of bottom water in the Central Atlantic Ocean region. All data corrected to DIC of bottom water (See III-A-1 section). The smoothed curve is obtained by 15-point running average.

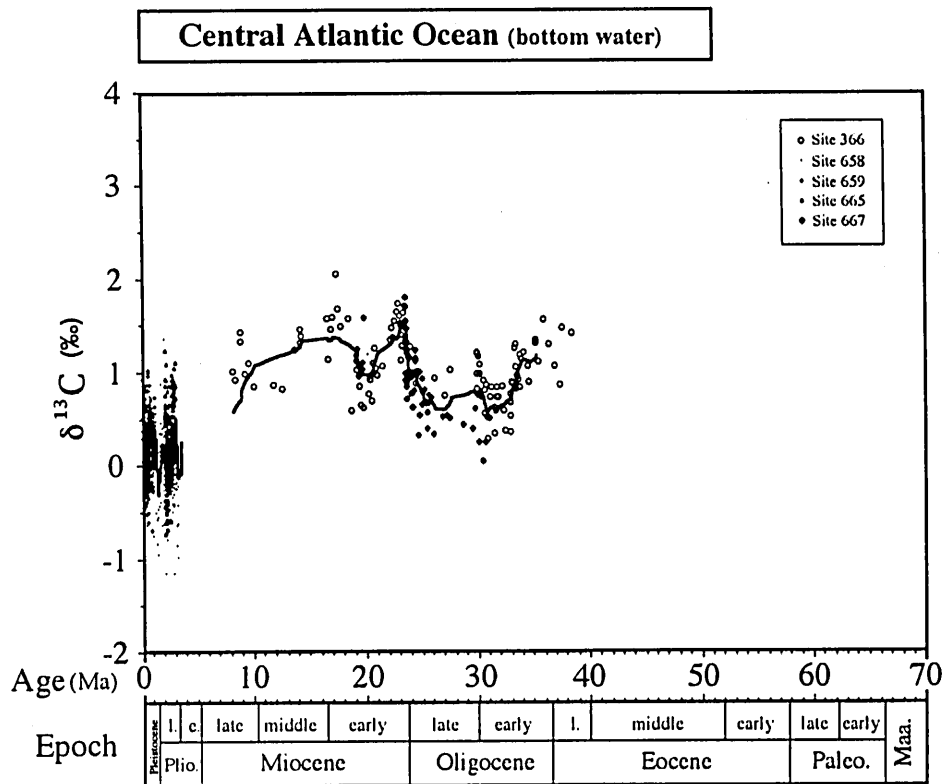


Fig. 45. Composite carbon isotope record of bottom water in the Central Atlantic Ocean region. All data corrected to DIC of bottom water (See III-A-1 section). The smoothed curve is obtained by 15-point running average.

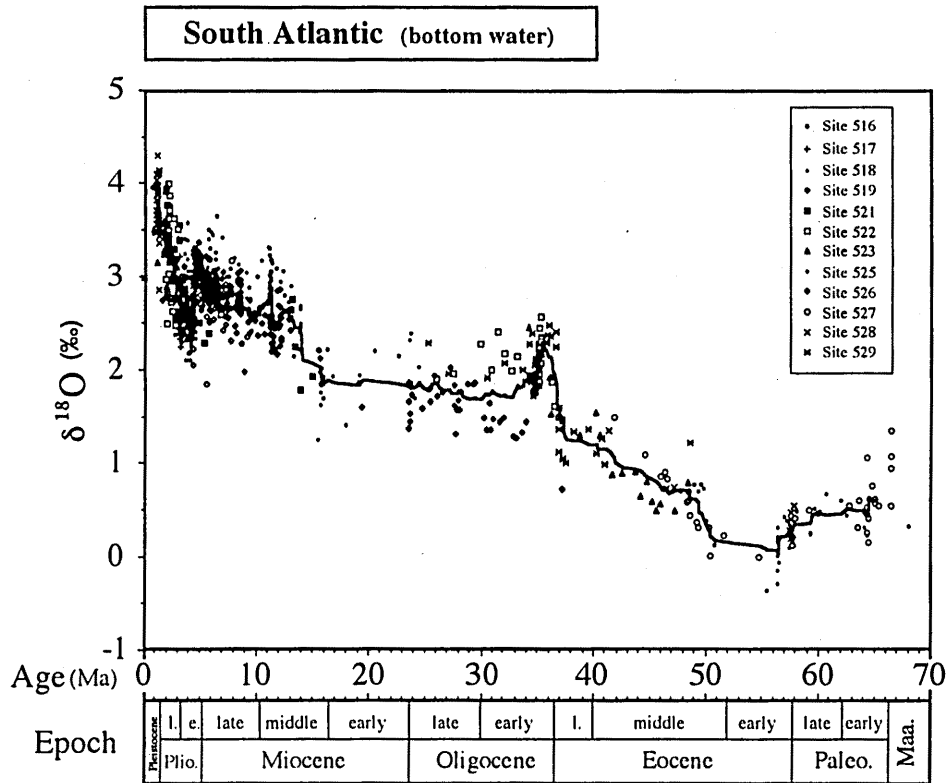


Fig. 46. Composite oxygen isotope record of bottom water in the South Atlantic Ocean region. All data corrected to DIC of bottom water (See III-A-1 section). The smoothed curve is obtained by 15-point running average.

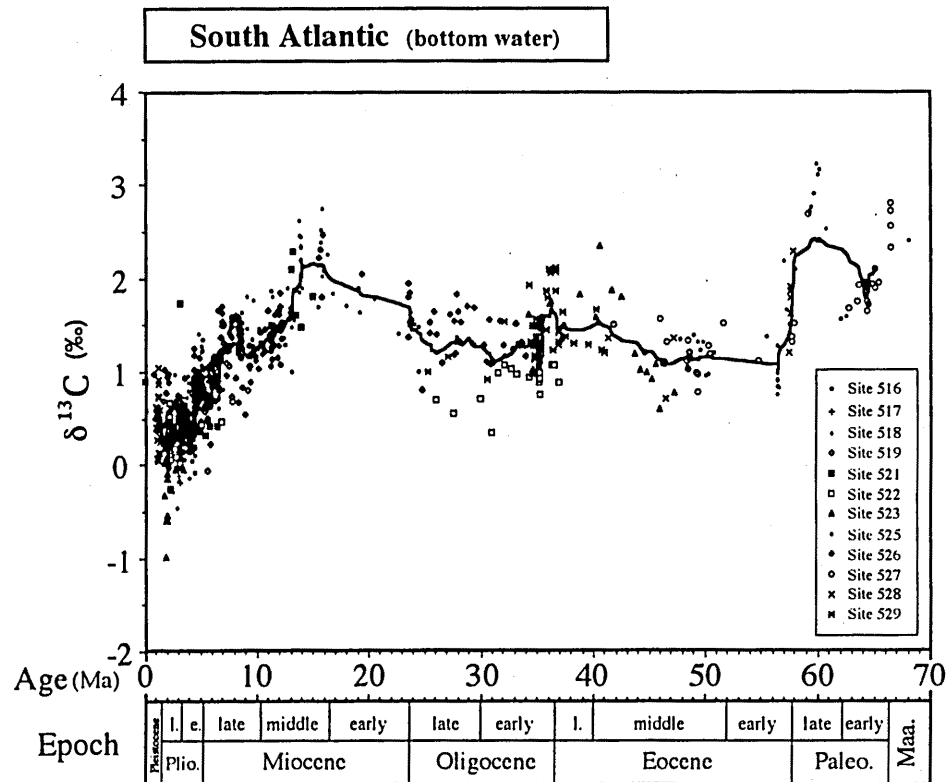


Fig. 47. Composite carbon isotope record of bottom water in the South Atlantic Ocean region. All data corrected to DIC of bottom water (See III-A-1 section). The smoothed curve is obtained by 15-point running average.

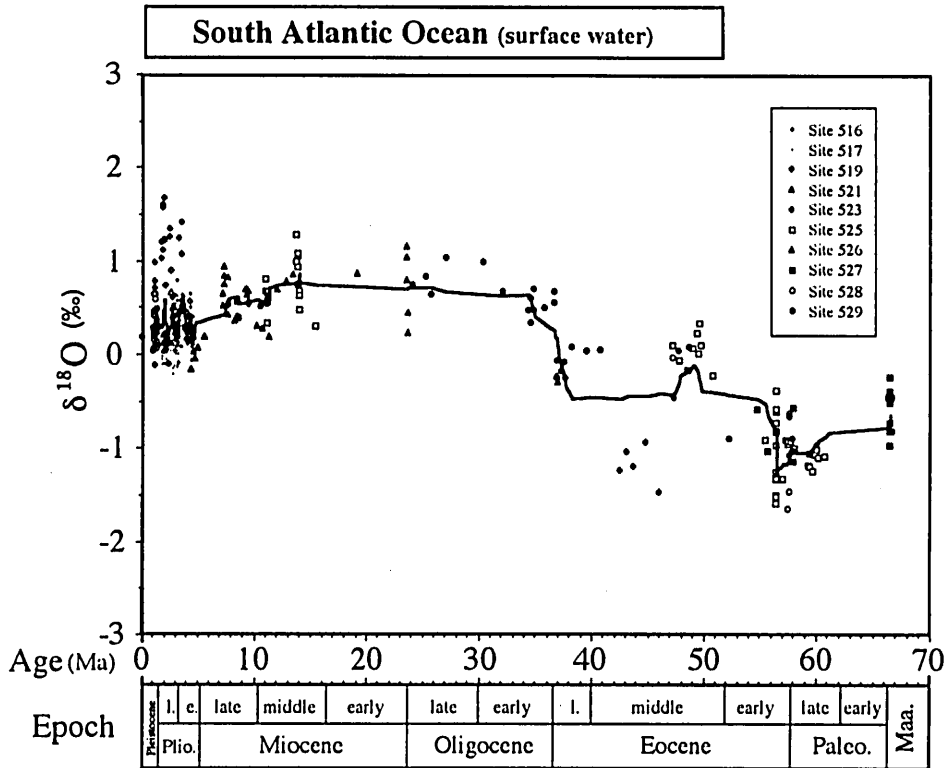


Fig. 48. Composite oxygen isotope record of surface water in the South Atlantic Ocean region. All data corrected to DIC of surface water (about 50 m below sea; See III-A-1 section). The smoothed curve is obtained by 15-point running average.

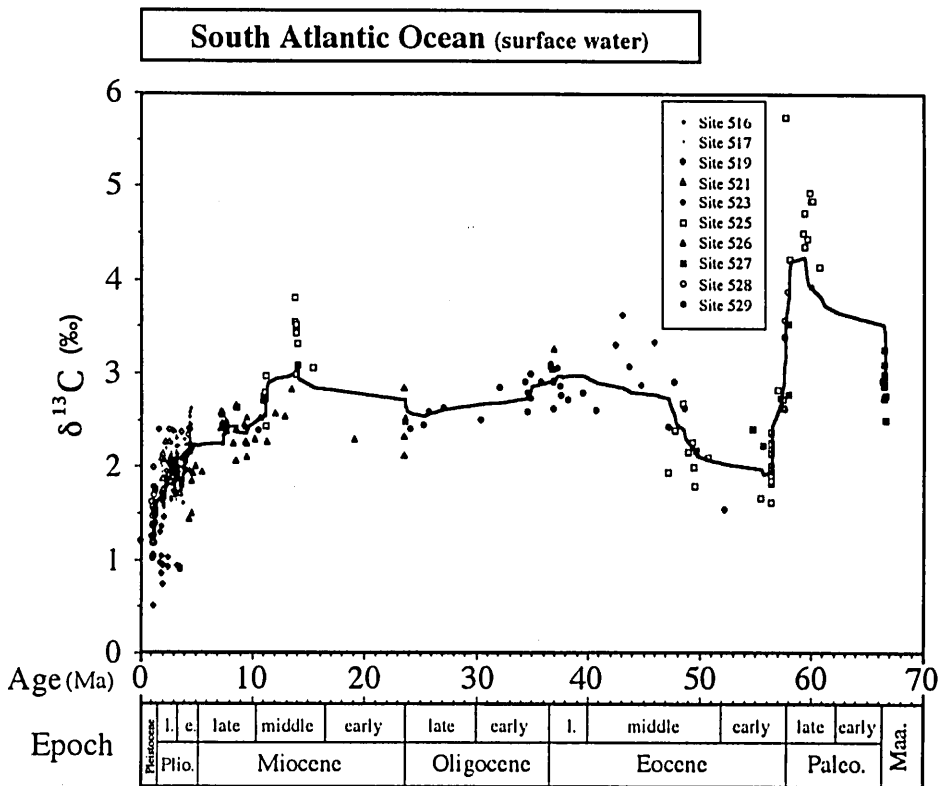


Fig. 49. Composite carbon isotope record of surface water in the South Atlantic Ocean region. All data corrected to DIC of surface water (about 50 m below sea; See III-A-1 section). The smoothed curve is obtained by 15-point running average.

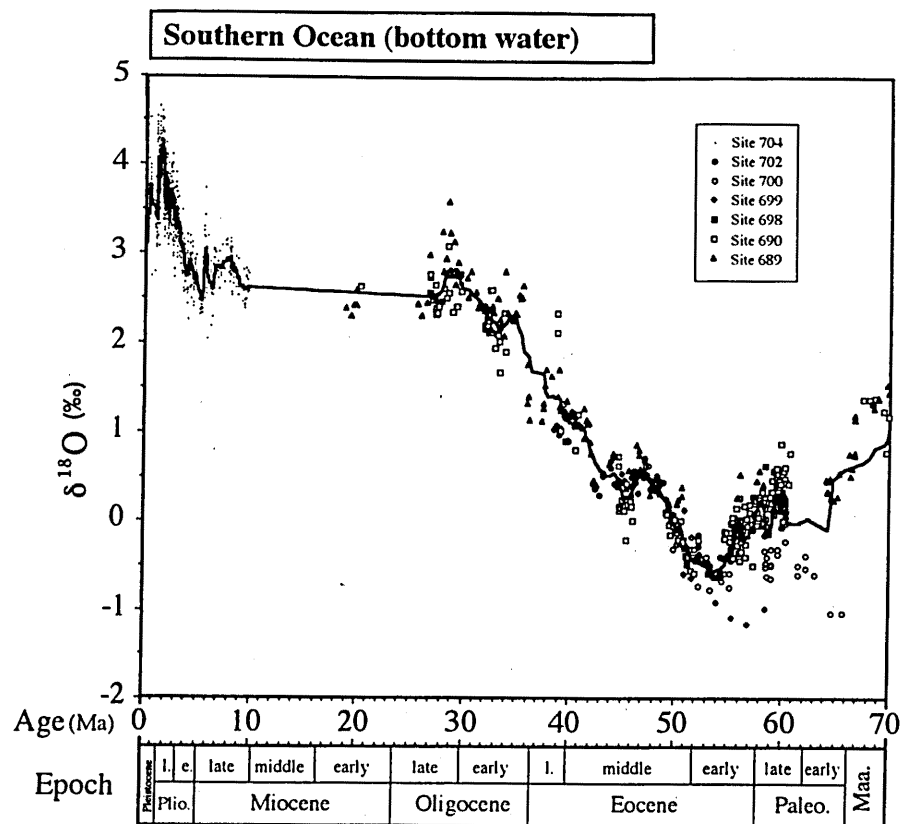


Fig. 50. Composite oxygen isotope record of bottom water in the Southern Ocean (Atlantic sector) region. All data corrected to DIC of bottom water (See III-A-1 section). The smoothed curve is obtained by 15-point running average.

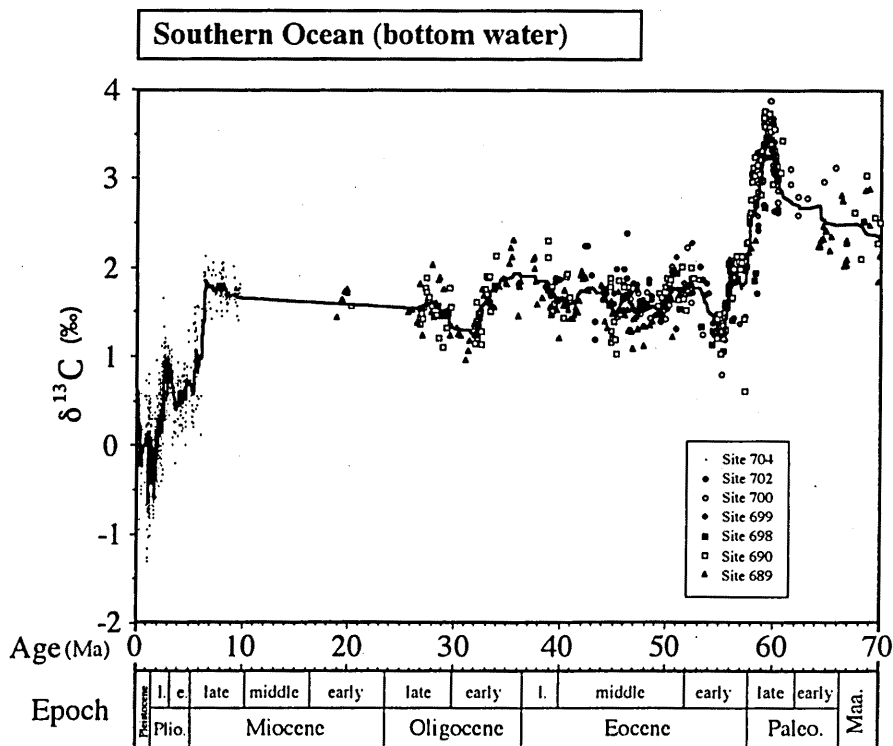


Fig. 51. Composite carbon isotope record of bottom water in the Southern Ocean (Atlantic sector) region. All data corrected to DIC of bottom water (See III-A-1 section). The smoothed curve is obtained by 15-point running average.

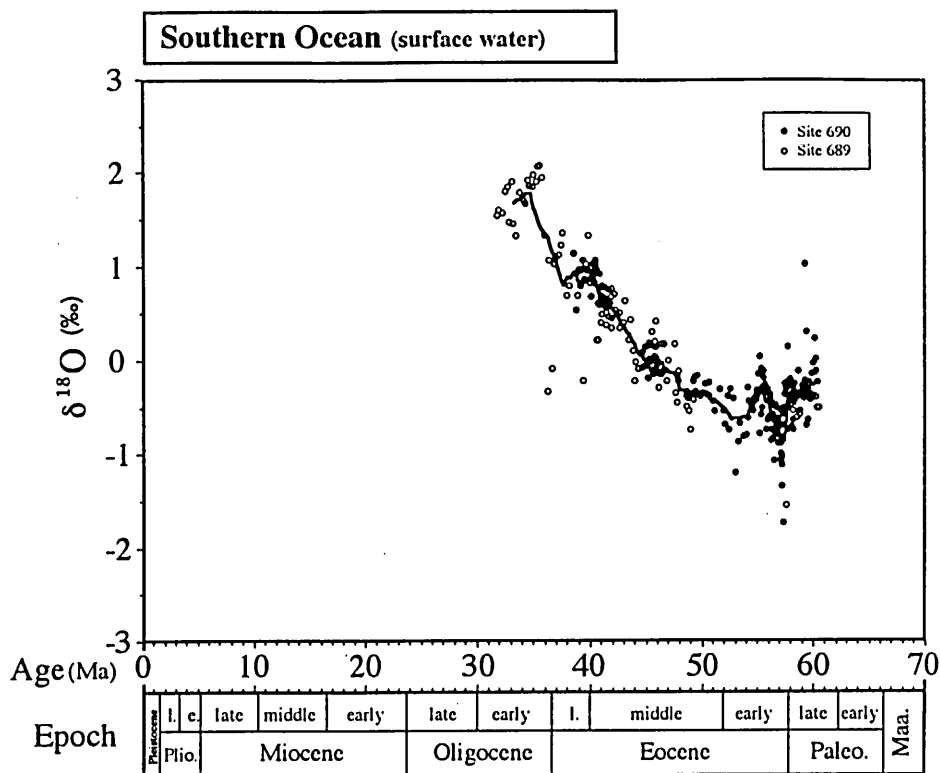


Fig. 52. Composite oxygen isotope record of surface water in the Southern Ocean (Atlantic sector) region. All data corrected to DIC of surface water (about 50 m below sea; See III-A-1 section). The smoothed curve is obtained by 15-point running average.

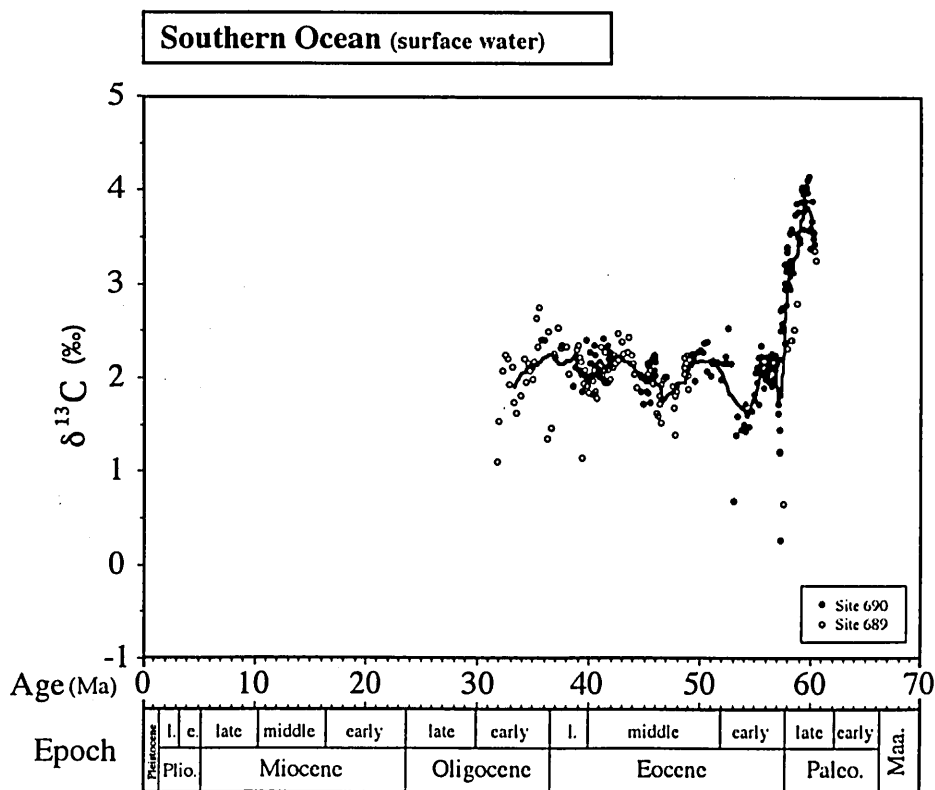


Fig. 53. Composite carbon isotope record of surface water in the Southern Ocean (Atlantic sector) region. All data corrected to DIC of surface water (about 50 m below sea; See III-A-1 section). The smoothed curve is obtained by 15-point running average.

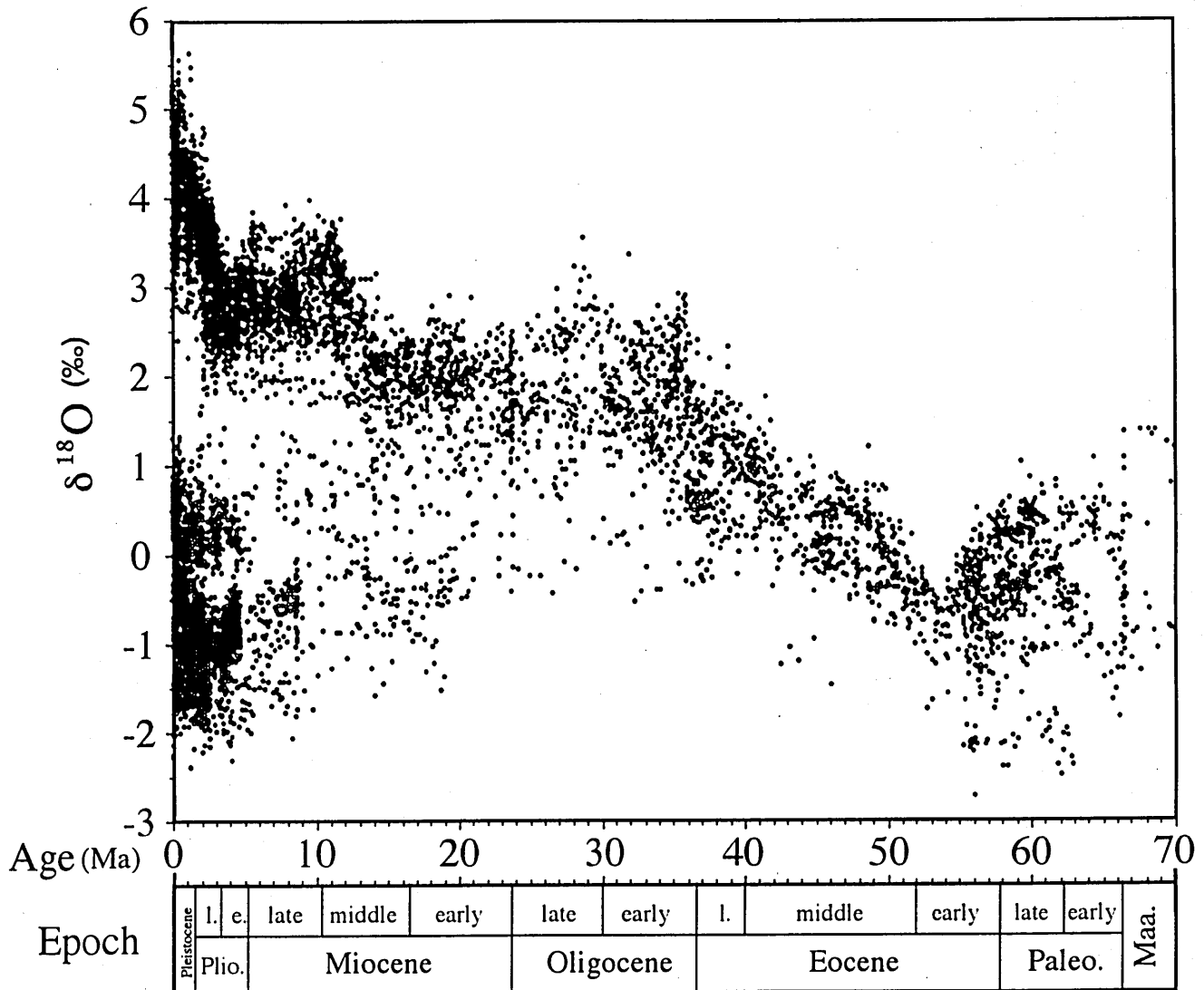


Fig. 54. All oxygen isotope data considered in this study. Open circle shows a DIC values of surface water (about 50 m below sea), and dot shows a DIC value of bottom water.

Indian, South Atlantic, and Southern Oceans. The net magnitudes are largest ($\sim 0.8\text{‰}$) in the Southern Ocean, and tend to decrease southwards as found in the shift of the bottom water record. The following $\delta^{18}\text{O}$ values differ among the oceans during the Oligocene. $\delta^{18}\text{O}$ values in the northern Indian Ocean are lowest ($\sim 0.2\text{‰}$); the Southern Ocean, southern Indian, and South Atlantic Oceans, respectively, record the values of 2.0, 1.4, and 1.0‰ , which are higher than those of the northern Indian Ocean. From the Miocene, although isotopes can only be recorded in the South Atlantic Ocean, the general trends are similar to those of the northern Indian Ocean. The degrees of decrease in those oceans are small from 12 Ma on consequently the $\delta^{18}\text{O}$ differences of those oceans expands to 1.4‰ .

Carbon isotopic record: The carbon isotopic records of surface and bottom water through the Cenozoic exhibit the

same pattern in the Indian and South Atlantic Oceans, except for a pronounced shift toward low values (by $\sim 1.0\text{‰}$) observed in the Southern Ocean (Atlantic sector) in the late Pliocene (2.5 Ma). The South Atlantic and northern Indian Oceans located in the same latitude record similar $\delta^{13}\text{C}$ values of bottom water. $\delta^{13}\text{C}$ values in the southern Indian Ocean are also similar with the exception of a lower value from the Eocene to late Oligocene. In the Central Atlantic Ocean, $\delta^{13}\text{C}$ values are slightly lower than those of the northern Indian Ocean. The difference of $\delta^{13}\text{C}$ values between the northern Indian and Southern Ocean varies according to age. $\delta^{13}\text{C}$ values in the Southern Ocean are close to those of the northern Indian Ocean in the late and middle Eocene, latest Pliocene, and Pleistocene, but they are higher through the Paleocene to early Eocene (0.5‰), and are higher through the Oligocene to early Pliocene (0.3‰).

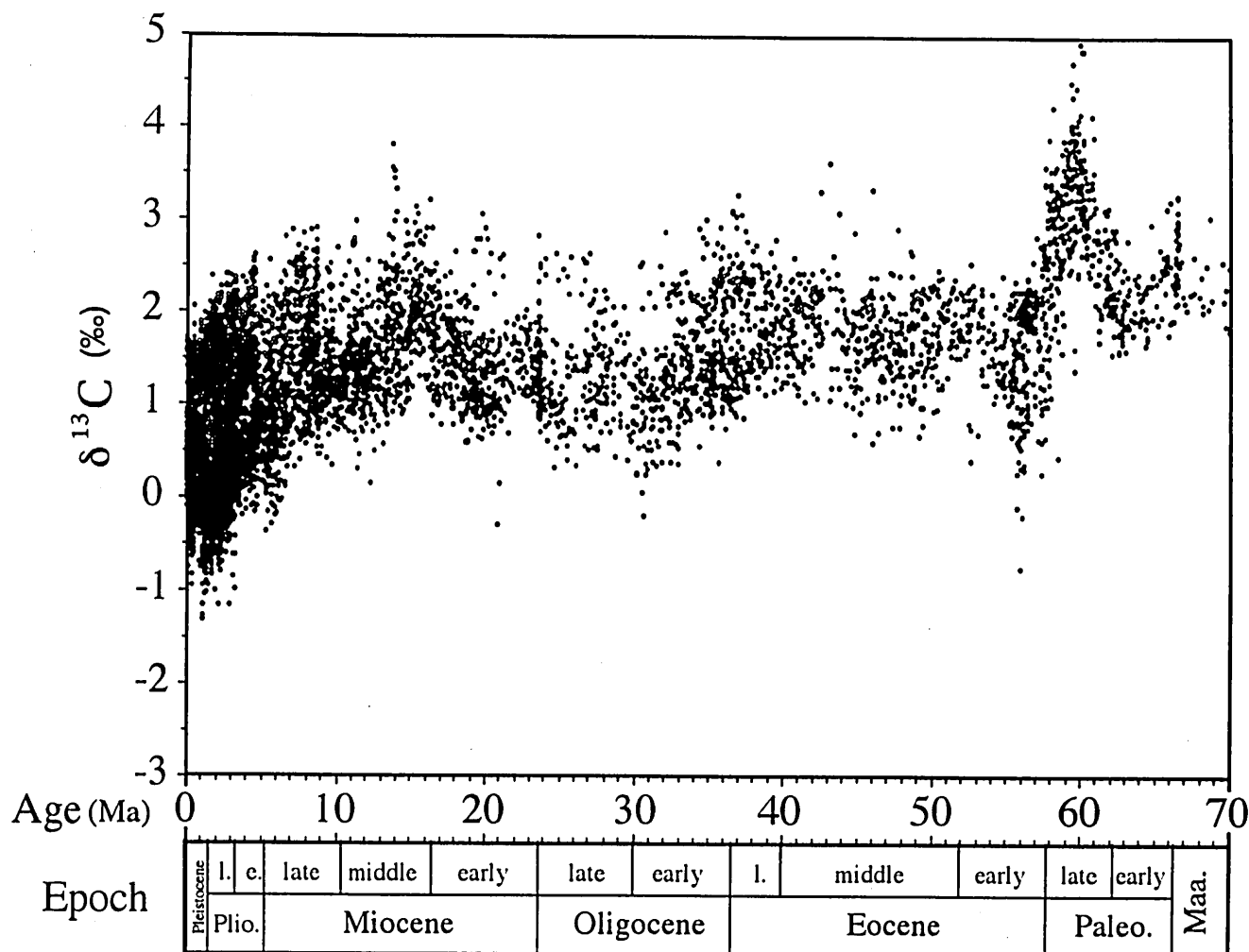


Fig. 55. All carbon isotope data considered in this study. Open circle shows a DIC values of surface water (about 50 m below sea), and dot shows a DIC value of bottom water.

For $\delta^{13}\text{C}$ values of surface water from the Paleocene to Oligocene, the southern Indian Ocean is generally similar to the northern Indian Ocean, however the former records 0.5‰ , which is lower than that of the latter in the middle Eocene. During the Pleistocene, $\delta^{13}\text{C}$ values of the Central Atlantic Ocean are lower than those of the northern Indian Ocean (1.0‰). $\delta^{13}\text{C}$ values in the South Atlantic Ocean are higher throughout the Cenozoic 0.5 to 1.0‰ . In the Southern Ocean (Atlantic sector), $\delta^{13}\text{C}$ values are 0.6‰ higher than those of the northern Indian Ocean during the Paleocene, which are similar to the early Eocene $\delta^{13}\text{C}$ values. They are slightly lower than those of the northern Indian Ocean in the middle and late Eocene.

The difference of $\delta^{13}\text{C}$ between surface and bottom water is smallest in the Southern Ocean (Atlantic sector), and largest in the South Atlantic Ocean. This difference tends to increase northwards. The Atlantic Ocean side exhibits larger difference than those of the Indian Ocean, although the two oceans are located at the same latitude.

D. General trends in the isotopic record

All oxygen and carbon isotope data used in this study are shown in Figs. 54 and 55. The general trends for oxygen and carbon isotopic records are similar in both the Indian and South Atlantic Oceans. In general, the oxygen and carbon isotopic values increase southwards.

The oxygen isotopic changes show a particular pattern from the Oligocene / Eocene boundary, that is, a gentle change followed by a sharp change in terms of a positive shift. Before and after the shift, a decrease in oxygen isotope values is recognized. The magnitudes of the shift of the Eocene / Oligocene boundary in surface and bottom water and the shift of the middle Miocene in bottom water tend to increase southwards. The shift of the late Pliocene tends to decrease southwards on the Atlantic ocean side, and decreases northwards on the Indian Ocean side. The difference in the magnitude of the shift may be caused by the difference in water mass structure.

As a general feature of carbon isotopes, the degree of variance of $\delta^{13}\text{C}$ tends to increase with a decrease in average

$\delta^{13}\text{C}$. This suggests that the carbon isotopic changes are related to ocean circulation velocity and marine productivity. The fluctuations in carbon isotopes from the latest Paleocene to early Eocene are remarkable in the Cenozoic record.

D. Distribution of oxygen and carbon isotopic values

The geographical distribution of oxygen and carbon isotopes are examined using reconstructed paleocoordinates and paleodepths. For the correlation among sites, the isotopic data are averaged in one million year intervals. The averaged values and standard deviations are shown in Table 7.

1. Oxygen isotopes

Maastrichtian: In the Maastrichtian, oxygen and carbon isotopic data are measured from Site 752 (around 50°S) in the northeastern Indian Ocean and Sites 689 and 690 (around 70°S) in the Atlantic sector of the Southern Ocean (Barrera and Huber, 1990). The oxygen isotopic values at Site 752 are constant at $\sim 0.9\text{‰}$ with a slight fluctuation, whereas those at Sites 689 and 690 tend to increase slightly by $\sim 1.0\text{‰}$. The difference between sites results from geographic difference, rather than variation in water depth, because the estimated water depth at Site 752 is close to that of Site 689.

Paleocene: From the early Paleocene to the time of the ^{18}O maximum at 61 Ma, the latitudinal gradient of oxygen isotopes in bottom water in the South Atlantic Ocean (70°–30°S) shows a slightly increasing trend toward low-latitudes, and the latitudinal difference is less than 0.2‰ (Fig. 56). In the Indian Ocean, $\delta^{18}\text{O}$ values on the high-latitude side are only 0.1‰ , which are higher than those on the low-latitude side in this interval. However, the oxygen isotopic values at 50°S are lower than those on the high- and low-latitude sides ($0.6\text{--}1.3\text{‰}$). The oxygen isotopic values in the Indian Ocean side are lower than those of the Atlantic Ocean side at similar latitude and water depth ($0.3\text{--}0.4\text{‰}$). This difference tends to reduce in this interval. $\delta^{18}\text{O}$ values on the low-latitude side in surface water are lower than those on the high-latitude side, and the difference on both sides is about 0.5‰ . In contrast, the $\delta^{18}\text{O}$ difference between surface and bottom water on the high-latitude side is lower than that on low-latitude sides. The oxygen isotopic values at about latitude 50°S in the Indian Ocean are lower than those of the high- and low-latitude sides, and similar to bottom water record ($1.0\text{--}1.5\text{‰}$). The surface to bottom difference of $\delta^{18}\text{O}$ values in this area is the largest among in the Indian and Atlantic Oceans reaching 1.6‰ at 62 Ma.

From 61 Ma in the Paleocene (Fig. 56), the oxygen isotopic values on the low-latitude side (around 40°S) in the Atlantic Ocean are close to those of the high-latitude side ($\sim 70^\circ\text{S}$). $\delta^{18}\text{O}$ values in mid-latitudes ($\sim 55^\circ\text{S}$) are $0.2\text{--}0.4\text{‰}$, lower than those of both sides. In this area, $\delta^{18}\text{O}$ values in intermediate water depth ($\sim 1000\text{m}$) are slightly higher than in deep water ($\sim 2000\text{m}$). In surface water in the South Atlantic Ocean, the oxygen isotopic values on the low-latitude side are $0.6\text{--}0.8\text{‰}$, lower than those on the high-latitude side, and the surface to bottom $\delta^{18}\text{O}$ difference on the low-latitude side is larger than that of the high-latitude sides. The latitudinal gradient of bottom water $\delta^{18}\text{O}$ in the Indian Ocean is similar to that of the Atlantic Ocean. The oxygen isotopic values on

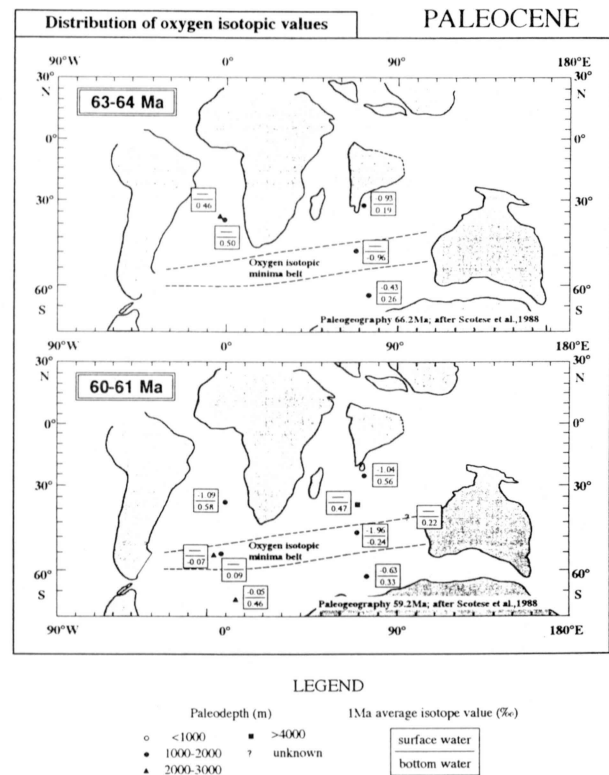


Fig. 56. Geographical distribution of oxygen isotopes during the Paleocene. The isotopic data are averaged in one million year intervals (Table 7).

the low-latitude side ($\sim 25^\circ\text{S}$) in bottom water are close to those on the high-latitude side ($\sim 65^\circ\text{S}$), and at mid-latitude ($\sim 50^\circ\text{S}$) records values lower by 0.6‰ than those of both sides. At high-latitudes, the oxygen isotopic values of bottom water in the Indian Ocean are close to those of the Atlantic Ocean. In surface water, the oxygen isotopic values on the low-latitude side in the Indian Ocean are $0.5\text{--}0.8\text{‰}$, lower than those on the high-latitude side; furthermore those at mid-latitudes are $0.6\text{--}0.9\text{‰}$, lower than those of the low-latitude side. The difference of $\delta^{18}\text{O}$ values between surface and bottom water is relatively large, being $\sim 1.6\text{‰}$ on the low-latitude side and $\sim 0.8\text{‰}$ on the high-latitude side. The difference at mid-latitudes is only 0.2‰ , larger than the low-latitude side.

Remarkably low values of oxygen isotopes at mid-latitudes are observed in the Indian and South Atlantic Oceans (Fig. 56). Low values in the Indian Ocean are situated slightly northward, compared with those in the South Atlantic Ocean. Low values in both oceans may be geographically connected by a belt-form (here called the "oxygen isotopic minima belt"). The "oxygen isotopic minima belt" may also exist in the late Maastrichtian.

Eocene: The pattern of oxygen isotopes in the early Eocene is similar to that in the latest Paleocene (Fig. 57-1). The oxygen isotopic values at high-latitudes in the Atlantic Ocean are $0.3\text{--}0.5\text{‰}$, higher than those at low-latitudes, with a smaller difference in the older part of the record. In this interval, $\delta^{18}\text{O}$ values in intermediate water ($\sim 1000\text{m}$) are

Table 7-1. Average oxygen and carbon isotopic data of the 1 Ma interval used in this study.

Interval (Ma)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)	No. of Data	Interval (Ma)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)	No. of Data
Central Atlantic Ocean				Site 667			
Site 98				Benthic			
Benthic				24-25	1.668±0.208	0.862±0.281	11
48-49	1.014	1.917	1	25-26	1.794±0.228	0.690±0.158	8
52-53	0.224±0.224	1.811±0.131	5	26-27	1.554	0.342	1
57-58	-0.162	1.392	1	27-28	1.678±0.185	0.529±0.013	3
58-59	-0.097	1.506	1	28-29	1.661	0.438	1
Site 366				29-30	1.484±0.160	0.506±0.151	2
Benthic				30-31	1.572±0.172	0.186±0.114	3
8-9	2.570±0.025	1.178±0.244	4	31-32	1.666	0.511	1
9-10	2.488±0.122	1.047±0.082	2	South Atlantic Ocean			
10-11	2.293	0.857	1	Site 516			
11-12	2.394	0.865	1	Benthic			
12-13	2.411	0.827	1	2-3	2.932±0.249	0.420±0.177	12
13-14	2.401	1.245	1	3-4	2.552±0.159	0.340±0.164	38
14-15	2.001±0.207	1.392±0.073	3	4-5	2.544±0.318	0.353±0.158	33
16-17	1.875±0.363	1.396±0.224	3	Planktonic			
17-18	1.577±0.205	1.707±0.244	4	2-3	0.346±0.157	1.882±0.137	16
18-19	1.745±0.088	1.087±0.685	2	3-4	0.336±0.184	1.974±0.164	40
19-20	1.886±0.227	0.792±0.189	4	4-5	0.265±0.151	2.226±0.201	33
20-21	1.893±0.136	0.987±0.185	8	Site 517			
21-22	2.068	1.071	1	Benthic			
22-23	1.841±0.184	1.534±0.143	7	2-3	2.645±0.124	0.646±0.025	3
23-24	1.742±0.220	1.388±0.214	7	3-4	2.561±0.263	0.431±0.236	12
24-25	1.456±0.092	0.964±0.074	3	Planktonic			
26-27	1.510	0.94	1	1-2	-0.028±0.055	1.935±0.165	6
27-28	1.445±0.007	0.895±0.191	2	2-3	0.105±0.184	2.018±0.175	33
29-30	1.873±0.273	1.060±0.185	4	3-4	0.211±0.224	2.038±0.201	25
30-31	1.857±0.159	0.797±0.272	10	Site 518			
31-32	1.669±0.193	0.697±0.170	7	Benthic			
32-33	1.645±0.179	0.597±0.187	7	2-3	3.138±0.216	-0.070±0.205	6
33-34	1.555±0.201	1.020±0.164	12	3-4	2.537±0.216	0.384±0.223	23
34-35	1.519±0.126	1.062±0.159	3	4-5	2.540±0.167	0.231±0.210	5
35-36	1.606±0.153	1.333±0.179	4	Site 519			
36-37	0.861±0.042	1.189±0.168	2	Benthic			
37-38	0.653±0.032	1.170±0.426	2	1-2	3.108±0.523	0.589±0.133	2
38-39	0.603	1.423	1	2-3	3.064±0.489	0.442±0.418	3
Site 658				3-4	2.876±0.082	0.529±0.152	4
Benthic				4-5	2.438±0.224	0.442±0.221	8
0-1	4.362±0.521	0.135±0.318	231	5-6	2.691±0.306	0.604±0.260	11
1-2	4.005±0.464	0.063±0.336	82	6-7	2.699±0.135	0.926±0.143	5
2-3	3.663±0.339	-0.010±0.350	133	7-8	2.585±0.241	1.116±0.213	6
3-4	3.309±0.251	-0.061±0.317	6	8-9	2.333±0.277	0.741±0.173	4
Planktonic				9-10	2.423±0.106	0.933±0.184	2
0-1	0.241±0.474	0.327±0.340	237	Planktonic			
1-2	0.229±0.336	0.572±0.304	90	1-2	0.085±0.290	2.190±0.283	2
2-3	0.680±0.115	0.546±0.144	4	2-3	0.482±0.398	2.040±0.384	5
Site 659				3-4	0.350±0.132	2.077±0.266	3
Benthic				4-5	0.355±0.290	2.375±0.049	2
0-1	4.419±0.495	0.252±0.376	163	Site 521			
Planktonic				Benthic			
0-1	0.202±0.482	0.450±0.333	108	1-2	3.508±0.054	0.349±0.153	4
Site 662				2-3	3.094±0.440	0.410±0.200	7
Planktonic				3-4	2.874±0.377	0.594±0.450	9
0-1	-0.947±0.203		103	4-5	2.512±0.131	0.364±0.059	4
1-2	-1.020±0.293		81	5-6	2.336±0.073	0.370±0.068	2
Site 663				6-7	2.723	0.421	1
Planktonic				13-14	2.353±0.440	1.865±0.387	4
0-1	-0.628±0.441		108	15-16	1.930	1.810	1
Site 665				Planktonic			
Benthic				1-2	0.147±0.047	1.9870.224	3
2-3	2.699±0.353	0.288±0.458	134	2-3	0.290±0.234	2.056±0.173	8
3-4	2.140±0.214	0.610±0.306	12	3-4	0.460±0.042	1.855±0.191	2
Site 667				4-5	0.440	2.080	1
Benthic				Site 522			
19-20	1.678±0.091	1.158±0.191	8	Benthic			
20-21	1.890±0.143	0.976±0.021	2	1-2	3.447±0.679	0.138±0.127	2
23-24	1.613±0.219	1.247±0.414	23	2-3	3.268±0.503	0.233±0.217	16
				3-4	2.783±0.050	0.176±0.212	2

Table 7-2. (continued).

Interval (Ma)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)	No. of Data	Interval (Ma)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)	No. of Data
Site 522				Site 525			
Benthic				Benthic			
4-5	2.508	0.294	1	15-16	1.751±0.266	2.239±0.346	7
5-6	2.708	0.436	1	16-17	2.073±0.194	2.041±0.298	2
6-7	2.588	0.464	1	18-19	1.410	1.701	1
26-27	1.893	0.693	1	19-20	1.905±0.049	1.761±0.184	2
27-28	1.956	0.556	1	20-21	2.200	1.781	1
29-30	2.273	0.713	1	22-23	2.145	1.411	1
30-31	1.996	0.356	1	23-24	2.359±0.052	1.561±0.101	2
31-32	2.410	0.980	1	24-25	1.852±0.168	1.327±0.279	3
32-33	2.076±0.127	1.056±0.028	2	47-48	0.709	1.339	1
33-34	2.146	0.966	1	49-50	0.744±0.032	1.315±0.065	4
34-35	1.936	0.946	1	50-51	0.266±0.128	1.037±0.139	3
35-36	2.193±0.169	1.199±0.217	32	55-56	-0.362	1.378	1
36-37	1.659±0.196	1.006±0.104	3	56-57	0.065±0.189	0.928±0.143	10
Site 523				Site 526			
Benthic				Benthic			
1-2	3.391±0.376	-0.211±0.429	9	0-1	3.953	0.973	1
2-3	3.021±0.199	0.012±0.334	9	1-2	3.453±0.004	0.888±0.174	2
3-4	2.728±0.172	0.205±0.198	6	2-3	3.391±0.121	0.757±0.252	5
34-35	1.993±0.387	1.363±0.312	3	3-4	3.001±0.085	0.676±0.225	3
35-36	1.995±0.191	1.530±0.046	4	4-5	3.023±0.205	0.877±0.138	40
36-37	1.720±0.269	1.735±0.021	2	5-6	2.893±0.194	0.760±0.180	29
37-38	1.450	1.360	1	6-7	2.842±0.114	1.172±0.274	23
38-39	1.292	1.820	1	7-8	2.795±0.147	1.372±0.192	10
40-41	1.420±0.173	1.962±0.541	2	8-9	2.772±0.157	1.364±0.123	24
41-42	0.872	1.860	1	9-10	2.549±0.128	1.134±0.135	7
42-43	0.890	1.790	1	10-11	2.549±0.141	1.314±0.205	7
43-44	0.910	1.190	1	11-12	2.423±0.192	1.351±0.162	19
44-45	0.730±0.113	1.000±0.028	2	12-13	2.513±0.181	1.516±0.200	11
45-46	0.535±0.064	0.995±0.106	2	13-14	2.415	1.685	1
46-47	0.560	0.600	1	15-16	2.016±0.165	2.188±0.285	4
47-48	0.482	0.770	1	19-20	1.600	2.040	1
48-49	0.790	1.050	1	23-24	1.595±0.183	1.665±0.222	6
Planktonic				Planktonic			
1-2	1.182±0.358	1.005±0.248	9	24-25	1.575	0.805	1
2-3	1.259±0.059	0.907±0.127	4	25-26	1.793±0.140	1.523±0.115	3
3-4	1.241±0.175	0.916±0.025	3	26-27	1.739±0.030	1.242±0.215	2
42-43	-1.240	3.300	1	27-28	1.666±0.240	1.493±0.260	6
43-44	-1.115±0.106	3.345±0.389	2	28-29	1.722±0.140	1.503±0.196	3
44-45	-0.940	2.870	1	29-30	1.843±0.011	1.488±0.272	2
46-47	-1.470	3.330	1	30-31	1.453±0.138	1.250±0.219	4
47-48	-0.460	2.430	1	31-32	1.453±0.018	1.311±0.310	2
Site 524				Site 525			
Benthic				Benthic			
55-56	-0.583	0.339	1	0-1	3.914	0.578	1
56-57	-0.270±0.132	1.299±0.467	2	2-3	3.534±0.313	0.782±0.264	2
61-62	-0.246	2.283	1	3-4	3.305±0.184	0.738±0.204	6
64-65	-0.706±0.042	2.193±0.184	2	4-5	2.985±0.090	0.927±0.305	10
65-66	-0.172±0.327	2.319±0.372	5	5-6	3.187±0.172	0.981±0.201	14
66-67	-0.846±0.499	2.211±0.227	6	6-7	3.133±0.324	0.841±0.136	10
67-68	-0.921±0.799	1.733±0.014	2	7-8	2.959±0.184	1.393±0.120	7
Planktonic				Planktonic			
55-56	-1.350	1.960	1	8-9	2.998±0.163	1.430±0.223	3
56-57	-1.225±0.064	2.260±0.424	2	9-10	2.732±0.250	1.487±0.319	5
61-62	-1.060	2.770	1	10-11	3.020±0.309	1.385±0.087	4
62-63	-0.780	2.910	1	11-12	2.917±0.253	1.427±0.168	20
Site 525				Site 526			
Benthic				Benthic			
0-1	3.914	0.578	1	12-13	2.830±0.167	1.528±0.241	5
2-3	3.534±0.313	0.782±0.264	2	13-14	2.473±0.226	2.029±0.484	12
3-4	3.305±0.184	0.738±0.204	6	14-15	2.220	1.891	1
4-5	2.985±0.090	0.927±0.305	10	Planktonic			
5-6	3.187±0.172	0.981±0.201	14	4-5	0.019±0.128	1.743±0.256	5
6-7	3.133±0.324	0.841±0.136	10	5-6	0.189	1.949	1
7-8	2.959±0.184	1.393±0.120	7	7-8	0.689±0.178	2.486±0.078	8
8-9	2.998±0.163	1.430±0.223	3				
9-10	2.732±0.250	1.487±0.319	5				
10-11	3.020±0.309	1.385±0.087	4				
11-12	2.917±0.253	1.427±0.168	20				
12-13	2.830±0.167	1.528±0.241	5				
13-14	2.473±0.226	2.029±0.484	12				
14-15	2.220	1.891	1				

Table 7-3. (continued).

Interval (Ma)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)	No. of Data	Interval (Ma)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)	No. of Data
Site 526				Site 529			
Planktonic				Planktonic			
8-9	0.481±0.115	2.338±0.248	4	35-36	0.501	2.919	1
9-10	0.642±0.061	2.339±0.155	6	36-37	0.234±0.456	2.935±0.216	4
10-11	0.294±0.021	2.409±0.170	2	37-38	-0.164±0.088	2.900±0.149	3
11-12	0.189	2.269	1	38-39	0.081	2.726	1
12-13	0.735±0.064	2.553±0.021	2	39-40	0.040	2.797	1
13-14	0.850	2.825	1	40-41	0.060	2.604	1
19-20	0.870	2.295	1	47-48	0.040	2.912	1
23-24	0.730±0.390	2.455±0.265	5	48-49	0.091	2.636	1
36-37	-0.225	3.045	1	52-53	-0.897	1.557	1
37-38	-0.283	3.264	1				
Site 527				Southern Ocean			
Benthic				Site 689			
4-5	2.975	0.765	1	Benthic			
7-8	3.165	0.685	1	18-19	2.382	1.439	1
41-42	1.498	1.508	1	19-20	2.428±0.117	1.702±0.049	2
44-45	1.088	1.208	1	25-26	2.427	1.494	1
46-47	0.861±0.035	1.331±0.230	3	26-27	2.571±0.361	1.563±0.218	3
48-49	0.586±0.111	1.171±0.128	4	27-28	2.498±0.040	1.465±0.159	4
49-50	0.332±0.033	0.915±0.110	3	28-29	3.081±0.314	1.706±0.251	6
50-51	0.158±0.212	1.228±0.057	2	29-30	2.838±0.183	1.476±0.185	5
51-52	0.228	1.518	1	30-31	2.664±0.151	1.263±0.047	3
54-55	-0.012	1.118	1	31-32	2.422±0.078	1.156±0.157	5
57-58	0.248±0.145	1.402±0.105	3	32-33	2.320±0.108	1.524±0.206	5
59-60	0.504	2.681	1	33-34	2.360±0.280	1.624±0.097	5
62-63	0.550	1.680	1	34-35	2.294±0.043	1.812±0.033	5
63-64	0.463±0.206	1.838±0.114	2	35-36	2.537±0.067	2.164±0.112	4
64-65	0.519±0.310	1.828±0.122	4	36-37	1.402±0.259	1.721±0.184	4
65-66	0.578±0.033	1.981±0.100	3	37-38	1.382±0.228	1.920±0.221	5
66-67	0.983±0.337	2.602±0.213	4	38-39	1.624	1.646	1
Planktonic				39-40	1.294±0.282	1.472±0.166	5
54-55	-0.574	2.410	1	40-41	1.149±0.064	1.564±0.103	6
55-56	-1.029	2.229	1	41-42	1.085±0.099	1.566±0.163	7
57-58	-0.850±0.412	3.164±0.531	2	42-43	0.624±0.232	1.859±0.088	4
66-67	-0.558±0.217	2.937±0.212	12	43-44	0.634	1.796	1
Site 528				44-45	0.593±0.167	1.448±0.131	4
Benthic				45-46	0.414±0.142	1.583±0.116	3
1-2	3.763±0.312	0.477±0.249	27	46-47	0.630±0.166	1.334±0.144	7
2-3	2.740	0.641	1	47-48	0.427±0.104	1.296±0.136	4
4-5	2.835±0.078	0.840±0.085	2	48-49	0.357±0.046	1.353±0.031	3
41-42	1.349	1.359	1	49-50	0.102±0.130	1.449±0.033	2
46-47	0.728	0.718	1	50-51	0.121±0.246	1.698±0.266	6
47-48	0.421±0.186	1.718±0.365	7	55-56	0.262	2.079	1
Planktonic				56-57	0.520	2.032	1
1-2	0.294±0.171	1.422±0.232	29	57-58	0.450	2.162	1
47-48	-0.034	2.440	1	58-59	0.490±0.113	2.252±0.057	2
57-58	-1.057±0.421	3.252±0.511	6	64-65	0.341±0.098	2.324±0.080	9
Site 529				65-66	0.409±0.177	2.256±0.113	2
Benthic				66-67	0.815±0.261	2.319±0.329	7
0-1	2.985	0.905	1	68-69	1.334±0.060	2.615±0.213	3
25-26	2.290	1.006	1	69-70	1.450±0.078	2.278±0.533	3
27-28	1.955	1.153	1	70-71	1.212	1.618	1
30-31	1.910	0.915	1	71-72	1.130±0.101	2.125±0.226	6
32-33	2.070	1.530	1	72-73	1.027±0.163	2.758±0.085	2
33-34	2.000	1.281	1	73-74	0.857±0.007	2.388±0.198	2
34-35	1.967±0.177	1.348±0.228	15	76-77	0.787±0.064	2.473±0.134	2
35-36	2.320±0.044	1.698±0.222	3	77-78	0.922	2.208	1
36-37	1.985±0.512	1.715±0.393	9	Planktonic			
37-38	1.024±0.035	1.504±0.184	2	31-32	1.578±0.039	1.325±0.304	2
38-39	1.334	1.299	1	32-33	1.672±0.184	2.111±0.140	4
39-40	1.360	1.290	1	33-34	1.618±0.271	1.823±0.207	4
40-41	1.119±0.139	1.364±0.261	3	34-35	1.799±0.107	2.074±0.107	5
48-49	1.214	1.174	1	35-36	1.994±0.069	2.458±0.233	5
Planktonic				36-37	0.602±0.755	1.992±0.543	5
0-1	0.185	1.205	1	37-38	1.203±0.120	2.355±0.134	4
10-11	0.519	2.389	1	38-39	0.773±0.109	2.230±0.128	4
24-25	0.751	2.409	1	39-40	0.833±0.425	1.954±0.346	9
25-26	0.740±0.129	2.525±0.093	2	40-41	0.731±0.269	2.012±0.128	16
27-28	1.041	2.639	1	41-42	0.592±0.149	2.112±0.111	12
30-31	0.991	2.509	1	42-43	0.516±0.125	2.236±0.138	5
32-33	0.681	2.859	1	43-44	0.260±0.299	2.293±0.099	6
34-35	0.526±0.138	2.810±0.156	5	44-45	-0.045±0.049	1.970±0.099	2
				45-46	0.178±0.229	2.020±0.128	5
				46-47	-0.077±0.155	1.799±0.183	10

Table 7-4. (continued).

Interval (Ma)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)	No. of Data	Interval (Ma)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)	No. of Data
Site 689				Site 699			
Planktonic				Benthic			
47-48	-0.183±0.216	1.763±0.198	6	50-51	0.120	1.308	1
48-49	-0.464±0.140	2.100±0.127	7	51-52	-0.475±0.259	1.895±0.190	3
49-50	-0.385±0.049	2.200±0.057	2	52-53	-0.369	2.266	1
57-58	-0.750±0.694	1.787±0.976	3	54-55	-0.916	1.707	1
58-59	-0.525±0.102	2.530±0.182	4	55-56	-1.093	1.384	1
60-61	-0.463±0.064	3.387±0.158	3	56-57	-1.160	1.361	1
				58-59	-0.993	2.569	1
Site 690				Site 700			
Benthic				Benthic			
20-21	2.624	1.556	1	46-47	0.479±0.075	1.693±0.130	4
26-27	2.678±0.106	1.452±0.069	4	47-48	0.567±0.099	1.592±0.111	3
27-28	2.437±0.107	1.643±0.142	8	48-49	0.377±0.064	1.575±0.209	2
28-29	2.687±0.269	1.416±0.149	4	49-50	-0.314	1.415	1
29-30	2.565±0.190	1.441±0.251	5	50-51	-0.180±0.071	1.848±0.233	3
31-32	2.154	1.326	1	51-52	-0.308	2.221	1
32-33	2.239±0.184	1.371±0.184	13	52-53	-0.729	1.610	1
33-34	2.035±0.213	1.815±0.205	7	53-54	-0.588±0.264	1.552±0.442	2
38-39	2.214±0.156	2.201±0.134	2	54-55	-0.554±0.129	1.341±0.116	3
39-40	1.103±0.169	1.655±0.127	5	55-56	-0.592±0.147	1.087±0.247	3
40-41	1.075±0.192	1.727±0.228	5	58-59	-0.463±0.111	3.165±0.131	5
44-45	0.372±0.268	1.522±0.343	6	59-60	-0.351±0.261	3.417±0.409	7
45-46	0.167±0.167	1.480±0.261	9	60-61	-0.067±0.137	2.853±0.124	3
46-47	0.087±0.117	1.679±0.004	2	61-62	-0.554±0.066	3.020±0.126	2
49-50	-0.012±0.104	1.650±0.105	5	62-63	-0.456±0.108	2.693±0.142	2
50-51	-0.079±0.106	1.884±0.062	4	63-64	-0.595	2.780	1
51-52	-0.411±0.088	1.705±0.181	5	64-65	-1.036	2.960	1
52-53	-0.411±0.186	1.875±0.125	3	65-66	-1.028	3.127	1
54-55	-0.214±0.140	1.344±0.124	4				
55-56	-0.089±0.156	1.583±0.291	19	Site 702			
56-57	-0.119±0.196	1.933±0.132	19	Benthic			
57-58	0.089±0.215	2.258±0.687	12	38-39	1.047±0.026	1.753±0.024	3
58-59	0.137±0.144	3.051±0.199	16	39-40	1.069±0.157	1.811±0.095	4
59-60	0.415±0.193	3.484±0.202	15	40-41	1.132±0.064	1.660±0.005	2
60-61	0.457±0.218	3.160±0.144	9	42-43	0.358±0.061	2.129±0.191	3
67-68	1.380	2.612	1	43-44	0.582±0.075	1.499±0.258	4
68-69	1.368±0.028	2.550±0.476	3	44-45	0.451±0.064	1.695±0.080	4
69-70	1.072±0.255	2.444±0.152	3	45-46	0.483±0.016	1.984±0.005	2
70-71	1.426±0.350	2.103±0.520	3	46-47	0.478±0.112	1.951±0.295	4
71-72	0.885±0.163	2.197±0.345	18	47-48	0.532±0.123	1.648±0.091	7
72-73	0.700±0.113	2.591±0.413	4	48-49	0.427±0.069	1.692±0.125	6
Planktonic				49-50	0.173±0.146	1.538±0.115	6
38-39	0.835±0.431	2.015±0.134	2	50-51	-0.085	1.528	1
39-40	0.958±0.090	2.100±0.224	4	52-53	-0.278±0.121	1.802±0.083	3
40-41	0.913±0.145	2.222±0.097	5	53-54	-0.549±0.088	1.847±0.134	4
41-42	0.630±0.125	2.197±0.191	6	54-55	-0.406	1.370	1
44-45	0.025±0.106	1.900±0.140	4	55-56	-0.237±0.057	1.230±0.059	2
45-46	-0.031±0.103	2.081±0.152	13	56-57	-0.157	1.431	1
46-47	-0.130	2.030	1	57-58	0.120±0.195	1.899±0.394	3
49-50	-0.177±0.198	2.117±0.176	6	58-59	0.095±0.228	2.196±0.420	3
50-51	-0.332±0.077	2.240±0.154	6	59-60	0.277±0.121	3.400±0.199	9
51-52	-0.460±0.139	2.103±0.107	3	60-61	0.090±0.110	2.738±0.206	4
52-53	-0.500±0.192	2.250±0.165	5				
53-54	-0.866±0.202	1.330±0.365	5	Site 704			
54-55	-0.439±0.115	1.633±0.134	7	Benthic			
55-56	-0.367±0.246	2.051±0.165	16	0-1	3.475±0.435	0.114±0.509	48
56-57	-0.661±0.180	2.122±0.110	21	1-2	3.834±0.343	-0.205±0.397	262
57-58	-0.619±0.425	2.353±0.822	23	2-3	3.517±0.250	0.502±0.415	188
58-59	-0.394±0.169	3.384±0.272	17	3-4	3.050±0.253	0.675±0.273	69
59-60	-0.243±0.391	3.852±0.229	17	4-5	2.794±0.189	0.588±0.240	58
60-61	-0.045±0.159	3.604±0.164	6	5-6	2.787±0.337	0.866±0.379	47
				6-7	2.745±0.205	1.577±0.435	29
				7-8	2.915±0.157	1.762±0.158	23
				8-9	2.829±0.166	1.746±0.160	28
				9-10	2.620±0.147	1.695±0.149	10
Site 698				Northwestern Indian Ocean			
Benthic				Site 237			
51-52	-0.406±0.102	1.715±0.207	3	Benthic			
52-53	-0.448	1.709	1	5-6	2.932±0.152	0.108±0.107	5
53-54	-0.591	1.338	1	6-7	2.949±0.052	0.183±0.207	7
54-55	-0.614±0.021	1.205±0.114	2	7-8	2.798±0.175	0.650±0.275	6
55-56	-0.216±0.228	1.321±0.272	3	8-9	3.102±0.324	0.973±0.148	11
56-57	0.024	1.403	1				
57-58	0.274	1.964	1				
58-59	0.288±0.334	2.357±0.798	3				
59-60	0.411±0.229	2.956±0.396	2				

Table 7-5. (continued).

Interval (Ma)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)	No. of Data	Interval (Ma)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)	No. of Data
Site 237				Site 709			
Benthic				Benthic			
9-10	2.950	1.180	1	11-12	2.916±0.075	1.167±0.158	11
12-13	3.010	1.290	1	12-13	2.775±0.159	1.024±0.587	4
13-14	2.550±0.208	1.263±0.365	6	13-14	1.972±0.203	1.476±0.352	14
14-15	2.070±0.157	1.570±0.336	3	14-15	2.105±0.175	1.324±0.329	24
15-16	2.255±0.035	2.170±0.212	2	15-16	1.970±0.223	1.945±0.330	6
16-17	1.835±0.218	1.797±0.538	6	16-17	2.024±0.128	1.108±0.096	9
17-18	1.968±0.080	1.413±0.100	6	17-18	1.936±0.136	0.936±0.157	5
18-19	1.948±0.146	1.328±0.237	12	18-19	1.971±0.134	1.023±0.130	4
19-20	1.869±0.088	1.165±0.111	12	19-20	2.224±0.149	1.107±0.114	12
20-21	1.990±0.173	1.083±0.153	3	20-21	2.053±0.142	0.955±0.116	4
21-22	2.095±0.163	1.205±0.021	2	21-22	2.328±0.089	1.720±0.005	2
22-23	2.000±0.057	1.150±0.042	2	22-23	2.322±0.177	1.708±0.206	8
23-24	1.835±0.223	1.285±0.420	4	23-24	2.127±0.187	1.301±0.323	8
Planktonic				Planktonic			
5-6	-0.446±0.297	1.338±0.201	6	1-2	-0.856±0.220	1.320±0.195	173
6-7	-0.602±0.200	1.743±0.203	7	2-3	-0.884±0.228	1.406±0.207	140
7-8	-0.616±0.215	2.001±0.258	6	3-4	-0.957±0.239	1.372±0.175	248
8-9	-0.457±0.162	2.018±0.228	12	4-5	-0.742±0.212	1.453±0.173	144
9-10	-0.843±0.219	2.175±0.163	2	Site 714			
10-11	-1.338	2.700	1	Benthic			
12-13	-0.078	2.090	1	7-8	2.745±0.138	0.863±0.133	4
13-14	-0.138±0.227	2.361±0.252	7	8-9	2.640±0.149	0.907±0.160	30
14-15	-0.531±0.278	2.600±0.121	3	9-10	2.456±0.034	0.955±0.078	2
15-16	-0.508±0.042	2.920±0.071	2	10-11	2.579±0.123	0.937±0.026	2
16-17	-0.784±0.145	2.616±0.467	6	11-12	2.540±0.110	1.194±0.192	10
17-18	-0.630±0.343	1.878±0.271	5	12-13	2.583±0.100	1.189±0.102	5
18-19	-0.536±0.249	1.936±0.191	11	13-14	2.140	2.020	1
19-20	-0.262±0.187	1.613±0.255	7	14-15	1.813	2.148	1
20-21	-0.238±0.398	1.657±0.029	3	15-16	1.951±0.581	1.760±0.269	2
22-23	0.662	1.280	1	23-24	1.873±0.201	1.271±0.220	6
Site 238				Planktonic			
Benthic				0-1	-1.030±0.501		20
5-6	3.137±0.233	0.097±0.276	6	1-2	-1.242±0.514		19
6-7	3.152±0.061	0.448±0.438	5	2-3	-1.054±0.279		19
7-8	2.948±0.159	0.810±0.176	6	3-4	-1.433±0.351		19
8-9	3.076±0.138	1.182±0.295	9	4-5	-1.269±0.497		16
9-10	3.550	-	1	5-6	-1.428±0.051		5
10-11	3.260	0.850	1	7-8	-2.115±0.168	2.100±0.257	4
11-12	2.825±0.304	0.955±0.686	2	8-9	-1.985±0.258	1.951±0.203	27
12-13	2.885±0.035	0.945±0.318	2	9-10	-2.330±0.283	2.060±0.099	2
13-14	2.438±0.304	1.150±0.233	4	10-11	-1.610	1.540	1
14-15	2.300±0.411	1.733±0.125	3	11-12	-1.925±0.313	2.072±0.245	10
15-16	1.935±0.325	1.735±0.409	4	Site 716			
16-17	2.020±0.453	1.560±0.255	2	Planktonic			
17-18	2.035±0.082	1.333±0.128	4	0-1	-0.941±0.381		112
18-19	2.135±0.290	0.985±0.120	2	1-2	-0.861±0.318		65
19-20	2.110	0.740	1	Northeastern Indian Ocean			
23-24	2.262±0.147	1.206±0.215	11	Site 214			
Planktonic				Benthic			
5-6	-0.886±0.297	2.047±0.270	7	5-6	3.133±0.222	0.267±0.259	6
6-7	-0.864±0.109	2.240±0.501	5	6-7	2.771±0.250	0.756±0.394	7
7-8	-1.067±0.156	2.536±0.316	5	7-8	2.967±0.097	1.300±0.138	12
8-9	-0.923±0.237	2.407±0.263	13	8-9	3.069±0.220	1.173±0.103	11
9-10	-1.014	2.270	1	9-10	2.220	0.680	1
10-11	-0.844	2.120	1	10-11	2.557±0.074	0.963±0.225	3
11-12	-0.860	2.400	1	11-12	2.570±0.156	0.550±0.339	2
12-13	-0.814±0.043	2.237±0.122	3	12-13	2.705±0.134	0.825±0.078	2
13-14	-0.882±0.294	2.379±0.158	4	13-14	2.305±0.228	0.970±0.261	5
14-15	-0.699±0.268	2.915±0.106	2	14-15	1.780	1.410	1
15-16	-0.618±0.213	2.860±0.143	6	15-16	1.745±0.148	1.565±0.007	2
16-17	-0.117±0.075	1.860±0.339	2	18-19	1.945±0.064	1.395±0.488	2
17-18	-0.318	1.910	1	19-20	1.900±0.170	1.145±0.290	2
18-19	-0.288	1.720	1	20-21	1.900	0.980	1
23-24	-0.264±0.190	2.020±0.113	2	21-22	1.970±0.057	1.130	2
Site 709				22-23	2.030	1.350	1
Benthic				23-24	1.970±0.028	1.050±0.028	2
4-5	3.120	0.270	1	Planktonic			
5-6	3.125±0.100	0.716±0.382	11	5-6	-0.821±0.197	1.984±0.210	8
7-8	3.138	1.060	1	6-7	-0.514±0.189	1.978±0.257	8
8-9	3.170	0.780	1	7-8	-0.542±0.120	2.407±0.188	18
9-10	2.995±0.021	0.885±0.078	2				
10-11	2.914±0.072	0.989±0.184	5				

Table 7-6. (continued).

Interval (Ma)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)	No. of Data	Interval (Ma)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)	No. of Data
Site 214				Site 253			
Planktonic				Benthic			
8-9	-0.354±0.158	2.133±0.284	15	32-33	1.554±0.049	1.290±0.184	2
9-10	-0.127±0.400	1.622±0.124	3	33-34	1.599	1.360	1
10-11	-0.381±0.222	1.978±0.046	2	35-36	1.384±0.191	1.455±0.064	2
11-12	-0.016±0.314	2.165±0.233	2	36-37	0.964±0.247	1.495±0.035	2
12-13	-0.008±0.255	2.155±0.134	2	Planktonic			
13-14	-0.228±0.152	2.303±0.135	5	4-5	0.288±0.297	1.920±0.041	4
14-15	-0.528	2.220	1	5-6	0.743±0.399	1.746±0.175	8
15-16	-0.631±0.141	2.168±0.301	4	6-7	1.020	2.010	1
16-17	-0.428±0.200	2.140±0.379	4	7-8	1.185±0.247	1.985±0.205	2
18-19	-0.069	2.110	1	8-9	1.240±0.076	2.173±0.156	4
Site 215				9-10	1.120±0.000	2.000±0.007	2
Benthic				10-11	1.025±0.007	2.130±0.007	2
55-56	-0.530±0.225	1.185±0.350	3	11-12	1.067±0.177	2.115±0.080	3
56-57	-0.080±0.000	1.647±0.064	2	12-13	1.075±0.247	1.840±0.403	2
57-58	-0.118±0.102	1.970±0.165	4	13-14	0.780	2.315	1
58-59	-0.035±0.108	2.343±0.791	4	14-15	0.573±0.311	2.415±0.178	3
59-60	0.270±0.221	3.074±0.218	5	15-16	0.645±0.021	2.800±0.049	2
60-61	0.470±0.269	2.829±0.169	3	16-17		2.000±0.021	2
Site 216				17-18	0.515±0.071	2.142±0.106	2
Benthic				18-19	0.475	2.157	1
8-9	3.041±0.119	0.850±0.121	7	19-20	0.455	2.657	1
10-11	2.850±0.141	0.875±0.078	2	21-22	0.110	2.607	1
11-12	2.773±0.090	0.757±0.175	3	24-25	-0.185±0.057	2.457±0.170	2
12-13	2.810±0.245	1.065±0.181	4	25-26	-0.235	2.197	1
13-14	2.453±0.194	1.477±0.110	3	26-27	-0.425	2.187	1
14-15	2.030±0.087	1.760±0.201	3	27-28	-0.075	2.137	1
15-16	1.910±0.424	1.735±0.290	2	28-29	-0.145	2.147	1
16-17	1.807±0.084	1.727±0.212	3	29-30	0.185	2.057	1
17-18	1.835±0.021	1.565±0.502	2	30-31	0.175±0.057	2.127±0.113	2
18-19	1.895±0.235	1.130±0.292	4	31-32	0.175±0.085	2.282±0.304	2
19-20	2.062±0.151	0.940±0.133	9	32-33	-0.355±0.200	2.194±0.032	3
20-21	2.195±0.325	0.735±0.521	8	33-34	-0.385	2.207	1
21-22	2.077±0.121	1.034±0.229	7	Site 752			
22-23	2.193±0.335	1.367±0.387	3	Benthic			
Planktonic				0-1	2.413	1.071	1
8-9	-0.333±0.158	1.787±0.270	7	1-2	2.313±0.156	1.421±0.028	2
10-11	-0.188±0.240	1.745±0.445	2	2-3	2.668±0.191	0.876±0.064	2
11-12	0.059±0.276	1.627±0.081	3	3-4	2.110±0.100	0.884±0.115	3
12-13	0.017±0.175	1.865±0.254	4	4-5	1.896±0.100	0.861±0.206	4
13-14	-0.238±0.014	2.105±0.092	2	5-6	2.003±0.014	0.891±0.057	2
14-15	-0.308±0.087	2.417±0.023	3	6-7	1.883±0.104	1.171±0.308	3
15-16	0.227±0.389	2.355±0.035	2	7-8	1.950±0.015	1.441±0.085	3
16-17	0.009±0.180	2.115±0.254	4	8-9	1.840±0.135	1.271±0.161	3
17-18	-0.373±0.021	2.405±0.021	2	9-10	1.703±0.014	1.566±0.177	2
18-19	-0.243±0.318	2.005±0.926	2	10-11	1.780±0.021	1.528±0.120	3
19-20	0.032±0.057	1.370±0.014	2	11-12	1.793±0.035	1.688±0.237	3
20-21	-0.165±0.316	1.627±0.087	3	12-13	1.798±0.021	1.771±0.014	2
22-23	-0.135±0.148	1.900±0.170	2	13-14	1.443±0.071	2.086±0.191	2
Site 253				14-15	1.063±0.240	1.851±0.255	2
Benthic				15-16	1.248±0.064	2.196±0.120	2
4-5	3.020±0.083	1.015±0.093	4	16-17	0.848±0.106	1.566±0.318	2
5-6	3.038±0.100	0.960±0.130	8	17-18	0.708±0.247	1.616±0.219	2
6-7	3.050±0.170	0.635±0.332	2	18-19	1.058±0.177	1.401±0.368	2
7-8	3.210±0.283	1.240±0.367	4	19-20	2.363±0.778	1.036±0.007	2
8-9	3.343±0.201	1.633±0.131	4	20-21	1.218±0.049	1.406±0.092	2
9-10	2.935±0.106	1.230±0.057	2	21-22	0.963±0.156	1.406±0.049	2
10-11	2.970±0.085	1.430±0.099	2	22-23	0.728±0.714	1.346±0.219	2
11-12	2.870±0.070	1.683±0.163	3	23-24	1.393	1.881	1
12-13	2.765±0.276	1.285±0.361	2	24-25	1.083	1.801	1
13-14	2.740	1.150	1	28-29	0.913	1.411	1
14-15	2.057±0.315	1.793±0.156	3	30-31	0.413	1.451	1
15-16	1.575±0.600	1.745±0.615	2	31-32	0.253	1.491	1
16-17	2.289±0.071	1.120±0.099	2	32-33	0.313	1.391	1
17-18	2.259	1.080	1	34-35	-0.377	1.721	1
18-19	2.369	1.170	1	35-36	-0.192±0.148	1.866±0.064	2
19-20	2.269	1.490	1	55-56	-0.789±0.199	0.655±0.282	13
21-22	2.349	1.440	1	56-57	-0.721±0.301	0.741±0.540	7
24-25	1.589	0.760	1	57-58	-0.588±0.086	1.325±0.293	3
28-29	2.009	1.010	1	58-59	-0.323±0.080	2.565±0.243	6
30-31	2.024±0.049	0.905±0.219	2	59-60	-0.308±0.140	2.578±0.274	3
31-32	1.939±0.297	1.190±0.127	2	60-61	-0.239±0.055	2.320±0.161	4
				61-62	-0.222±0.090	2.083±0.193	12
				62-63	-0.524±0.161	1.917±0.178	16

Table 7-7. (continued).

Interval (Ma)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)	No. of Data	Interval (Ma)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)	No. of Data
Site 752				Site 756			
Benthic				Benthic			
63-64	-0.955±0.179	2.066±0.307	5	23-24	1.355±0.081	1.167±0.047	3
64-65	-0.940±0.125	2.262±0.179	4	24-25	1.471±0.161	1.095±0.078	3
65-66	-1.386±0.179	2.175±0.274	6	25-26	1.567±0.174	0.911±0.255	3
66-67	-1.059±0.386	2.181±0.249	9	26-27	1.567±0.056	0.906±0.048	3
67-68	-0.875±0.345	2.168±0.087	4	27-28	1.592±0.097	1.010±0.085	3
68-69	-0.851±0.224	2.079±0.069	3	28-29	1.430±0.044	1.077±0.085	4
69-70	-0.821±0.019	1.941±0.172	3	29-30	1.397±0.032	0.841±0.099	2
70-71	-0.578±0.006	1.432±0.018	2	30-31	1.303±0.233	1.423±0.127	5
71-72	-1.105±0.371	1.433±0.150	2	31-32	1.400±0.217	1.108±0.228	9
Planktonic				Planktonic			
55-56	-2.114±0.094	1.534±0.673	7	32-33	1.608±0.335	1.020±0.224	4
56-57	-2.232±0.270	1.567±0.527	5	33-34	1.222±0.105	1.213±0.186	7
57-58	-2.264±0.173	1.914±0.243	2	34-35	1.200±0.260	1.561±0.298	5
58-59	-2.177±0.149	3.175±0.416	4	35-36	0.684±0.089	1.430±0.149	4
59-60	-1.973±0.146	3.581±0.264	2	36-37	0.571±0.146	1.516±0.146	6
60-61	-1.957±0.128	3.291±0.119	3	37-38	0.475±0.099	1.294±0.085	2
61-62	-1.976±0.228	2.843±0.218	6	Site 757			
62-63	-2.192±0.199	2.647±0.216	7	Benthic			
Site 754				Benthic			
Benthic				Benthic			
0-1	2.847±0.113	0.936±0.230	3	0-1	3.088±0.201	0.108±0.277	5
1-2	2.741±0.073	1.015±0.190	3	1-2	3.145±0.213	0.195±0.231	7
2-3	2.696±0.088	0.997±0.133	6	2-3	2.800±0.196	0.342±0.106	5
3-4	2.240±0.196	0.885±0.110	6	3-4	2.632±0.186	0.309±0.210	9
4-5	2.249±0.227	0.808±0.125	4	4-5	2.505±0.127	0.429±0.166	8
5-6	2.189±0.071	0.922±0.150	4	5-6	2.506±0.118	0.363±0.132	6
6-7	1.957±0.038	1.120±0.168	4	6-7	2.625±0.290	0.648±0.272	7
7-8	2.092±0.059	1.469±0.067	2	7-8	2.356±0.227	1.158±0.282	5
8-9	2.078±0.163	1.331±0.170	2	8-9	2.825	1.355	1
9-10	1.933±0.036	1.492±0.076	3	9-10	2.415±0.213	1.089±0.124	2
10-11	1.933	1.531	1	10-11	2.464±0.014	0.957±0.291	2
11-12	2.169±0.112	1.182±0.254	3	11-12	2.309±0.155	1.141±0.118	3
12-13	1.873±0.132	1.523±0.234	16	12-13	2.484	0.491	1
13-14	1.681±0.226	1.798±0.284	11	13-14	2.118±0.150	1.255±0.266	3
14-15	1.058±0.096	2.222±0.131	3	14-15	1.354±0.127	1.464±0.076	4
15-16	1.232±0.210	2.233±0.250	3	16-17	1.528±0.069	1.050±0.220	2
16-17	0.953±0.028	1.670±0.012	2	17-18	1.719±0.084	1.187±0.114	3
17-18	1.223±0.226	1.301±0.410	2	18-19	1.784	1.117	1
18-19	1.313	1.511	1	19-20	1.443	1.293	1
19-20	1.273±0.141	1.493±0.088	2	23-24	1.687±0.053	1.218±0.001	2
20-21	1.285±0.136	1.482±0.113	4	25-26	1.584	1.429	1
21-22	1.238±0.164	1.530±0.083	4	27-28	1.666	1.256	1
22-23	1.287±0.160	1.764±0.179	3	28-29	1.634	1.379	1
23-24	1.228±0.271	1.580±0.235	4	30-31	1.723±0.077	1.284±0.038	4
24-25	1.100±0.222	1.513±0.243	2	31-32	1.619	1.206	1
25-26	1.263	1.311	1	33-34	1.620±0.021	1.575±0.102	2
26-27	1.221±0.017	0.952±0.128	2	34-35	1.631±0.132	1.790±0.210	4
27-28	1.213	1.301	1	35-36	1.219±0.163	1.568±0.182	2
28-29	1.227±0.020	1.212±0.114	5	36-37	0.922±0.011	1.350±0.182	2
Site 756				Benthic			
Benthic				Benthic			
1-2	3.003±0.187	0.683±0.086	2	37-38	1.069±0.051	1.856±0.131	4
2-3	3.026±0.573	0.398±0.052	3	38-39	0.941±0.066	1.748±0.001	2
3-4	2.623±0.333	0.681±0.312	5	40-41	0.878±0.099	1.353±0.121	3
4-5	2.699±0.084	0.757±0.148	4	41-42	0.916±0.068	1.562±0.177	2
5-6	2.652±0.136	0.574±0.247	6	42-43	0.974	1.677	1
6-7	2.498±0.167	0.676±0.093	5	43-44	0.618±0.122	1.812±0.021	2
7-8	2.446±0.133	1.168±0.160	4	44-45	0.684	1.635	1
8-9	2.523±0.084	1.221±0.140	3	45-46	0.488	1.664	1
9-10	2.535	1.434	1	47-48	0.451±0.010	1.638±0.057	2
10-11	2.376±0.030	1.184±0.141	2	48-49	0.294	1.575	1
11-12	2.498±0.131	1.581±0.547	2	49-50	0.110±0.222	1.575±0.133	2
12-13	2.143±0.167	1.312±0.042	3	51-52	-0.187±0.141	1.793±0.019	2
13-14	1.572	1.359	1	53-54	-0.681±0.192	1.563±0.282	3
14-15	1.645	1.603	1	54-55	-0.799	1.044	1
15-16	1.394±0.245	2.174±0.270	3	55-56	-0.903	0.849	1
17-18	1.214±0.000	1.534±0.159	2	56-57	-1.141	0.972	1
18-19	1.479±0.233	1.444±0.245	2	Planktonic			
19-20	1.509±0.245	1.567±0.067	2	35-36	0.402	2.367	1
20-21	1.616±0.193	1.493±0.143	4	37-38	0.430	2.343	1
21-22	1.629±0.115	1.824±0.089	3	38-39	0.197	2.445	1
22-23	1.446±0.083	1.442±0.306	3	40-41	0.317	2.421	1
				41-42	0.182	2.431	1
				43-44	0.006	2.609	1
				45-46	-0.159	2.390	1
				47-48	-0.231	2.212	1
				48-49	-0.170	2.413	1
				49-50	-0.583±0.275	2.182±0.083	2

Table 7-8. (continued).

Interval (Ma)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)	No. of Data	Interval (Ma)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)	No. of Data
Site 757				Site 762			
Planktonic				Benthic			
51-52	-0.872±0.194	2.259±0.033	2	61-62	-0.042±0.172	2.337±0.273	9
52-53	-1.115	2.302	1	Southern Indian Ocean			
53-54	-1.433±0.292	2.094±0.255	2	Site 738			
54-55	-1.552	1.458	1	Benthic			
55-56	-1.636	1.395	1	35-36	2.034±0.394	1.440±0.223	5
56-57	-1.933	1.523	1	36-37	1.475±0.167	1.106±0.099	9
Site 758				37-38	1.543±0.220	1.447±0.141	9
Benthic				38-39	1.547±0.188	1.505±0.190	5
0-1	3.751±0.314	0.022±0.227	13	39-40	1.272±0.049	1.800±0.109	2
1-2	3.640±0.407	-0.021±0.225	12	40-41	1.012	1.763	1
2-3	3.242±0.533	-0.079±0.216	9	42-43	0.484±0.187	1.706±0.062	3
3-4	2.699±0.209	0.188±0.220	9	44-45	0.727	1.583	1
4-5	2.703±0.209	0.061±0.205	7	45-46	0.612±0.035	1.582±0.085	2
5-6	2.645±0.369	0.030±0.292	8	46-47	0.610±0.205	1.279±0.236	3
6-7	2.528±0.226	0.666±0.301	7	47-48	0.469±0.220	1.264±0.174	5
7-8	2.536±0.336	0.904±0.256	8	48-49	0.430±0.114	1.212±0.135	3
8-9	2.535±0.361	0.825±0.312	6	49-50	0.187	0.982	1
10-11	2.676	1.151	1	50-51	0.023±0.257	1.753±0.151	5
11-12	2.269±0.325	0.722±0.033	2	51-52	-0.290±0.040	1.869±0.042	3
12-13	1.679	1.008	1	52-53	-0.343±0.000	1.962±0.198	2
14-15	1.796±0.120	1.468±0.214	3	53-54	-0.553	1.972	1
15-16	1.669	1.569	1	54-55	-0.511±0.095	1.581±0.294	5
17-18	1.848	1.014	1	55-56	-0.383±0.099	1.452±0.353	2
18-19	1.793±0.062	0.985±0.556	2	56-57	-0.292±0.204	0.939±0.314	7
19-20	2.045±0.015	1.097±0.189	3	57-58	0.182	1.628	1
20-21	1.549±0.616	1.060±0.252	4	58-59	0.262	2.013	1
21-22	1.510±0.235	1.400±0.442	3	60-61	0.332±0.073	3.192±0.241	4
22-23	1.805±0.208	1.627±0.182	3	61-62	0.462±0.414	2.314±0.383	5
23-24	1.347±0.264	1.287±0.169	7	62-63	0.376±0.130	1.964±0.134	9
25-26	1.758±0.134	0.983±0.317	3	63-64	0.257±0.162	1.950±0.182	3
27-28	1.786±0.121	0.975±0.192	2	65-66	-0.033±0.164	2.710±0.293	5
28-29	1.715	0.712	1	66-67	-0.163	2.582	1
29-30	1.500	1.330	1	Planktonic			
30-31	1.684±0.259	0.728±0.514	8	35-36	0.757±0.107	2.597±0.120	3
31-32	1.790±0.183	0.574±0.264	2	36-37	0.497±0.087	2.353±0.182	7
34-35	1.675±0.303	0.938±0.541	2	37-38	0.503±0.192	2.358±0.160	9
41-42	1.775	1.033	1	38-39	0.615±0.323	1.793±0.209	5
58-59	0.291	2.836	1	39-40	0.627±0.292	2.287±0.075	2
59-60	0.340±0.320	3.031±0.163	2	40-41	0.293	2.365	1
60-61	0.563±0.158	2.731±0.118	2	41-42	0.009±0.491	2.241±0.257	2
61-62	0.321±0.075	2.221±0.150	2	42-43	-0.018±0.053	2.139±0.164	3
63-64	0.185	2.023	1	44-45	0.062	1.939	1
Planktonic				45-46	0.062±0.071	1.765	2
0-1	-1.250±0.409	1.411±0.185	142	46-47	0.009±0.160	2.127±0.074	3
1-2	-1.478±0.232	1.621±0.222	121	47-48	-0.157±0.218	1.814±0.246	6
2-3	-1.599±0.251	1.654±0.173	81	48-49	-0.305±0.111	1.911±0.081	3
3-4	-1.903±0.135	1.542±0.180	12	49-50	-0.564±0.191	1.940±0.118	3
4-5	-1.605±0.231	1.741±0.245	10	50-51	-0.680±0.072	2.110±0.386	5
5-6	-1.701±0.259	1.740±0.320	10	51-52	-0.676±0.208	2.120±0.351	3
6-7	-1.358±0.193	2.074±0.223	8	52-53	-1.355±0.531	2.060±0.000	2
7-8	-1.568±0.116	2.181±0.226	9	53-54	-0.879±0.134	2.127±0.110	2
8-9	-1.710±0.248	2.382±0.395	5	54-55	-1.078±0.099	1.980±0.179	5
9-10	-1.624±0.148	1.655±0.346	2	55-56	-0.855±0.502	1.640±0.099	2
11-12	-1.069±0.276	2.108±0.134	2	56-57	-1.173±0.215	1.708±0.240	4
12-13	-1.154	2.178	1	57-58	-0.850±0.156	2.270±0.226	2
14-15	-1.507±0.095	2.610±0.145	2	58-59	-0.560	2.500	1
15-16	-1.184	2.562	1	60-61	-0.625±0.092	3.870±0.042	2
17-18	-1.299	2.062	1	61-62	-0.465±0.163	2.730±0.099	2
18-19	-1.369±0.148	1.969±0.102	3	62-63	-0.563±0.144	2.243±0.165	4
58-59	-1.383	3.239	1	63-64	-0.426±0.047	2.288±0.134	5
59-60	-1.340±0.301	3.372±0.010	2	Site 744			
60-61	-1.037±0.025	3.132±0.044	2	Benthic			
61-62	-0.922±0.140	2.609±0.271	2	9-10	3.131±0.021	1.090±0.085	2
63-64	-0.927	2.388	1	10-11	3.023±0.212	1.338±0.235	5
Site 762				11-12	3.051±0.381	1.217±0.356	6
Benthic				12-13	2.657±0.041	1.305±0.035	2
55-56	-0.559±0.168	0.869±0.359	2	13-14	2.287±0.236	1.788±0.203	5
56-57	-0.547±0.116	0.973±0.197	3	14-15	1.937±0.237	1.810±0.141	12
57-58	-0.460±0.128	0.987±0.168	8	15-16	2.033±0.334	1.994±0.187	7
58-59	-0.074	1.977	1	16-17	1.989±0.292	1.977±0.372	8
59-60	-0.183±0.040	2.817±0.297	3	17-18	2.068±0.316	1.774±0.142	19
60-61	0.216±0.187	2.821±0.107	5	18-19	2.292±0.238	1.592±0.138	5

Table 7-9. (continued).

Interval (Ma)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)	No. of Data	Interval (Ma)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)	No. of Data
Site 744				Site 748			
Benthic				Planktonic			
19-20	2.491±0.060	1.433±0.130	4	39-40	0.344	1.761	1
20-21	2.220±0.121	1.260±0.200	5	40-41	0.201±0.009	1.801±0.226	2
21-22	2.176	1.280	1	42-43	0.012±0.229	2.021±0.268	2
24-25	2.576	1.650	1	43-44	0.034	2.041	1
25-26	2.666	1.560	1	44-45	-0.217±0.159	1.757±0.306	3
26-27	2.226±0.208	1.385±0.151	4	45-46	0.274±0.269	1.661±0.099	2
27-28	2.360±0.128	1.020±0.197	5	47-48	0.534	0.771	1
28-29	2.651±0.148	1.130±0.014	2	49-50	0.104	0.671	1
30-31	2.556	1.050	1	52-53	-0.496±0.085	0.776±0.049	2
31-32	3.356	0.990	1	55-56	-0.546	0.331	1
32-33	2.502±0.125	1.128±0.090	5	58-59	-0.426±0.115	1.904±0.330	3
33-34	2.476±0.125	1.253±0.240	4	59-60	-0.699±0.187	2.964±0.395	4
34-35	2.559±0.083	1.298±0.077	4				
35-36	2.412±0.348	1.532±0.154	9	Site 750			
36-37	1.696±0.196	1.158±0.286	5	Benthic			
37-38	1.690±0.339	1.274±0.161	5	65-66	0.195±0.021	2.555±0.073	4
38-39	1.312±0.141	1.440±0.168	8	66-67	0.173±0.217	2.341±0.131	12
39-40	1.656±0.044	1.527±0.112	3				
Planktonic				Site 751			
19-20	1.122±0.181	2.839±0.151	5	Benthic			
20-21	0.760±0.259	2.519±0.227	5	4-5	3.410±0.133	0.955±0.105	4
21-22	0.956	2.389	1	5-6	3.570±0.137	1.297±0.160	18
24-25	1.146	2.629	1	6-7	3.578±0.101	1.24±0.168	9
25-26	1.206	2.389	1	7-8	3.593±0.209	1.196±0.148	5
26-27	0.766±0.075	2.494±0.098	4	8-9	3.489±0.145	1.135±0.093	12
27-28	0.799±0.256	2.054±0.117	6	9-10	3.313±0.252	1.202±0.215	41
28-29	1.131±0.106	2.149±0.127	2	10-11	3.354±0.172	1.372±0.262	54
30-31	1.186	2.539	1	11-12	3.319±0.226	1.407±0.273	55
31-32	1.316	2.299	1	12-13	3.005±0.128	1.658±0.109	8
32-33	1.114±0.261	2.257±0.045	15	13-14	3.099±0.003	2.073±0.211	3
33-34	1.091±0.083	2.337±0.123	4	14-15	2.727±0.393	1.779±0.124	4
34-35	1.159±0.092	2.107±0.049	4	15-16	2.325±0.179	1.878±0.172	11
35-36	1.217±0.259	2.242±0.142	8	16-17	2.164±0.219	1.835±0.162	17
36-37	0.560±0.111	2.257±0.107	5	17-18	2.224±0.334	1.672±0.326	8
37-38	0.528±0.172	2.285±0.096	5	18-19	2.481±0.194	1.316±0.361	12
38-39	0.344±0.168	2.402±0.181	6	19-20	2.360±0.024	1.106±0.192	2
39-40	0.313±0.087	2.489±0.171	3				
Site 748							
Benthic							
22-23	2.292	1.149	1				
23-24	2.174±0.114	1.118±0.048	3				
24-25	2.321±0.182	0.957±0.081	2				
25-26	2.407±0.095	1.180±0.159	3				
26-27	2.409±0.125	0.955±0.573	2				
27-28	2.394±0.179	0.885±0.007	2				
28-29	2.310±0.141	0.830±0.028	2				
29-30	2.515±0.195	0.765±0.069	4				
30-31	2.521±0.126	0.677±0.117	2				
31-32	2.211±0.098	0.477±0.117	2				
32-33	2.437±0.196	0.650±0.104	3				
33-34	2.310±0.036	1.027±0.132	3				
34-35	2.270±0.198	0.985±0.049	2				
35-36	2.393±0.530	1.178±0.323	12				
36-37	1.620±0.110	1.184±0.187	5				
37-38	1.350±0.057	1.425±0.168	5				
38-39	1.662±0.071	1.484±0.097	4				
39-40	1.431±0.055	1.420±0.000	2				
40-41	1.260±0.145	1.233±0.106	3				
41-42	1.164	1.129	1				
42-43	0.970	1.070	1				
43-44	0.903±0.143	1.355±0.280	3				
44-45	0.880±0.141	1.040±0.491	2				
52-53	-0.560	0.410	1				
55-56	-0.220	-0.750	1				
58-59	-0.090	0.450	1				
59-60	0.092±0.359	2.251±0.695	5				
Planktonic							
33-34	1.377±0.212	1.839±0.310	4				
34-35	0.997±0.170	1.831±0.101	3				
35-36	1.287±0.387	2.190±0.211	10				
36-37	0.745±0.104	2.097±0.213	7				
37-38	0.455±0.259	2.379±0.098	5				
38-39	0.334±0.121	1.941±0.378	3				

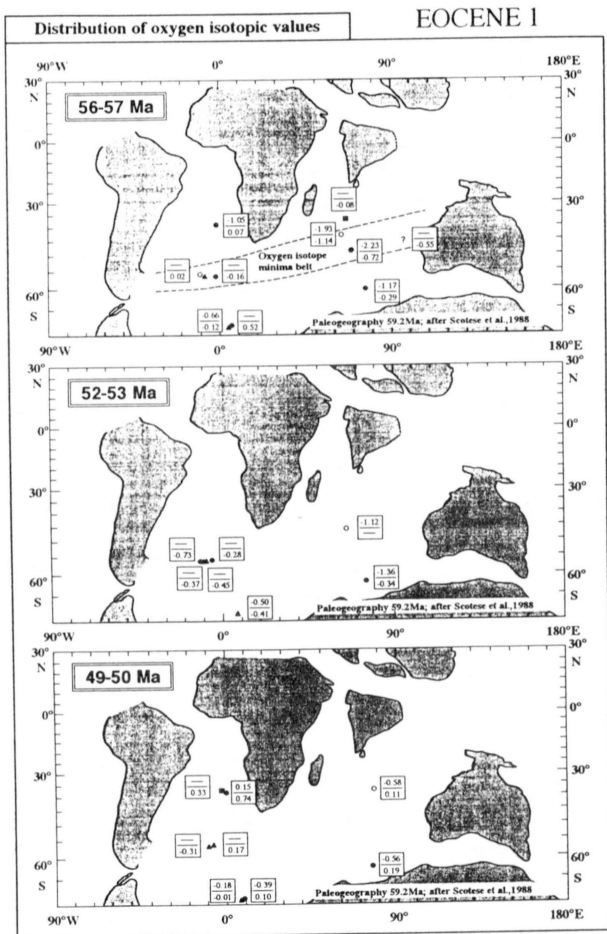


Fig. 57-1. Geographical distribution of oxygen isotopes during the Eocene. The isotopic data are averaged in one million year intervals (Table 7). See legend of Fig. 56.

distinctly high. Latitudinal difference of surface water $\delta^{18}\text{O}$ at this time is smaller than that in the latest Paleocene; for example the difference between 40°S and 70°S ranges from 0.2-0.4‰. Distinctly lower values of $\delta^{18}\text{O}$ of surface and bottom water in the Indian and South Atlantic Oceans became close to $\delta^{18}\text{O}$ values at high-latitudes in this interval, and those values became close to values recognized at the middle/early Eocene boundary. As a result, "oxygen isotopic minima belt" in the Indian and South Atlantic Oceans disappear at this time. $\delta^{18}\text{O}$ values in the Indian Ocean are slightly lower than those in the Atlantic Ocean.

During the middle Eocene, the oxygen isotopic values of bottom water in the Atlantic Ocean tend to increase northward from high-latitudes, and the oxygen isotopic difference between the high- and low-latitude sides (around 70° - 40°S) is 0.2-0.6‰ at the same depth. From 49 to 51 Ma, $\delta^{18}\text{O}$ values of intermediate water (~2000m) at high-latitudes are 0.2-0.3‰ lower than those of deep water. In contrast, $\delta^{18}\text{O}$ values of intermediate water (1000-2000m) at high-latitudes are 0.3-0.5‰ higher than those from 47 to 43 Ma. The benthic $\delta^{18}\text{O}$ values at 40 Ma are close to values at depths shallower than 2500 m. In surface water of the Atlantic

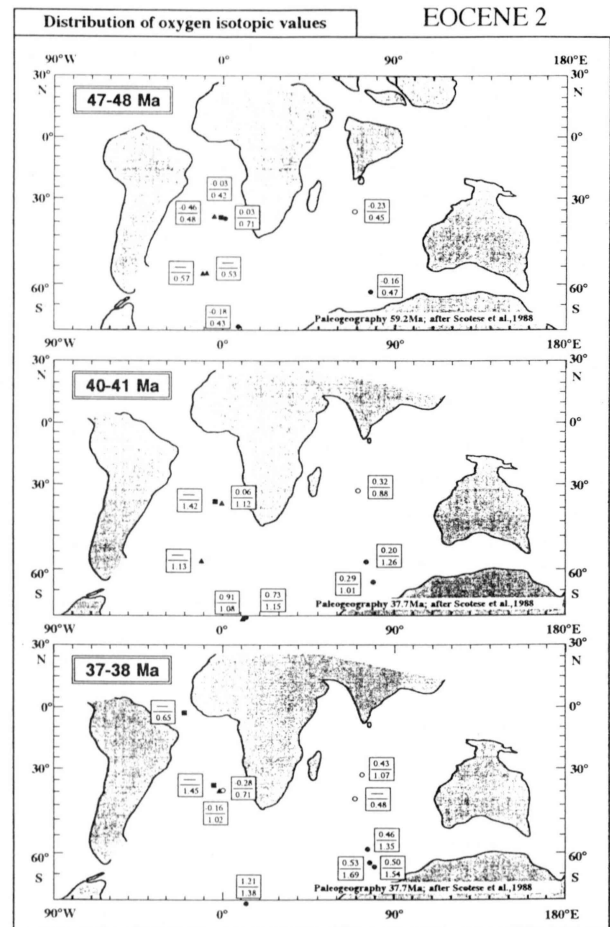


Fig. 57-2. (continued).

Ocean, $\delta^{18}\text{O}$ values at low-latitude are about 0.3‰ higher than those at high-latitudes during the period 51-48 Ma (Fig. 57). In the Indian Ocean, the oxygen isotopic values of bottom water tend to increase southward from low-latitudes in the middle Eocene, and the latitudinal difference of $\delta^{18}\text{O}$ value between 35°S and 65°S is 0.1-0.4‰ at a different water depth (Fig. 57-2). This difference tends to be expanded. The oxygen isotopes of intermediate water (1000-1500m) at high-latitudes in the Indian Ocean show relatively high values in comparison with deep water, which is a similar oxygen isotopic record to the Atlantic Ocean. No latitudinal gradient of oxygen isotopes is observed in surface water of the Indian Ocean.

In the late Eocene, the oxygen isotopic values of bottom water in the Atlantic Ocean tend to decrease toward low-latitudes the opposite pattern to the middle Eocene, and then the latitudinal difference between 70°S and 0° is about 0.6‰ (Fig. 57-2). The pattern of oxygen isotopes in surface water is similarly to that in bottom water. The latitudinal difference between 70° and 40°S in surface water is about 0.6-1.4‰. Although the latitudinal gradient of oxygen isotopes in the Indian Ocean in the late Eocene is close to that of the

middle Eocene, the latitudinal difference of $\delta^{18}\text{O}$ values between 65°S and 35°S is expanded by 0.3-0.6‰.

Oligocene: In the Atlantic Ocean, the oxygen isotopic values of bottom water in the Oligocene exhibit a decreasing trend toward low-latitudes, and the gradient at the late Oligocene is steeper than that of the early Oligocene (Fig. 58). The latitudinal difference from 70°S to 0° at ~2000m water depth ranges from 0.8-1.0‰ in the earliest Oligocene (36-34 Ma). This difference is rapidly reduced in the 34-33 Ma interval, and becomes to 0.4-0.6‰ in the late early Oligocene (34-30 Ma). This difference is again expanded by 0.6-1.2‰ in the late Oligocene. The oxygen isotopic values in intermediate water (~2000 m paleodepth) at high-latitudes are relatively higher than those in deep water. The water column gradient in other regions around the Atlantic Ocean shows a decreasing trend toward shallower depth. $\delta^{18}\text{O}$ values in surface water around the Atlantic Ocean also tend to decrease northward from the high-latitude side during the Oligocene. The latitudinal difference of $\delta^{18}\text{O}$ value in surface water between 70°S and 40°S is 1.0-1.4‰, and show a reducing trend.

The oxygen isotopic values of bottom water in the

Indian Ocean tend to decrease northward from high-latitudes similar to the oxygen isotopic record in the Atlantic Ocean. No latitudinal difference is recognized to the north of 40°S. At 2000 m paleodepth, the latitudinal difference between 65°S and 10°S is constant throughout the Oligocene (approximately 0.9‰). The latitudinal difference (65°S-35°S) in surface water is 1.4‰ from 35 to 32 Ma, reducing to ~1.0‰ at 32-31 Ma, and tends expand gently from 32 to 24 Ma. At 1000m paleodepth, the latitudinal difference between 60°S and 30°S is ~0.7‰ with a small fluctuation throughout the Oligocene. This fluctuation may reflect an isotopic change in surface water. The water column gradient in the Indian Ocean exhibits a decreasing trend toward shallower depths.

Miocene: During the Miocene, the oxygen isotopic values of bottom water in the Atlantic Ocean show a decrease north-ward from the high-latitude side, which is similar to the Oligocene isotopic record (Fig. 59-1). At 2500m water depth, the $\delta^{18}\text{O}$ difference between the high- and low-latitude sides (around 35°S-5°N) is relatively large from 25-20 Ma and 13-7 Ma, respectively (0.3-0.8‰). In these intervals, the largest differences are recognized at 24-23 Ma and 11-10 Ma, respectively. The water column gradient tends to decrease

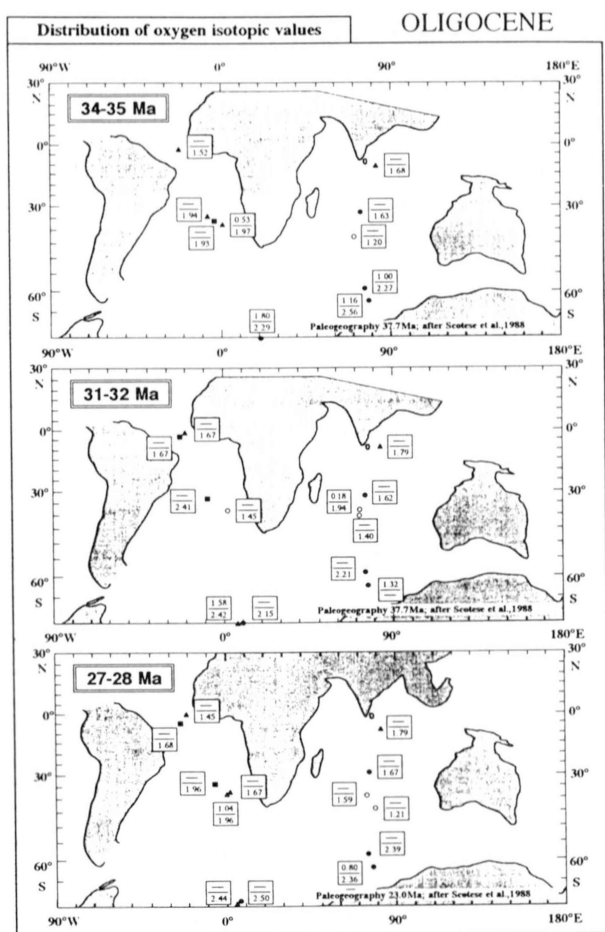


Fig. 58. Geographical distribution of oxygen isotopes during the Oligocene. The isotopic data are averaged in one million year intervals (Table 7). See legend of Fig. 56.

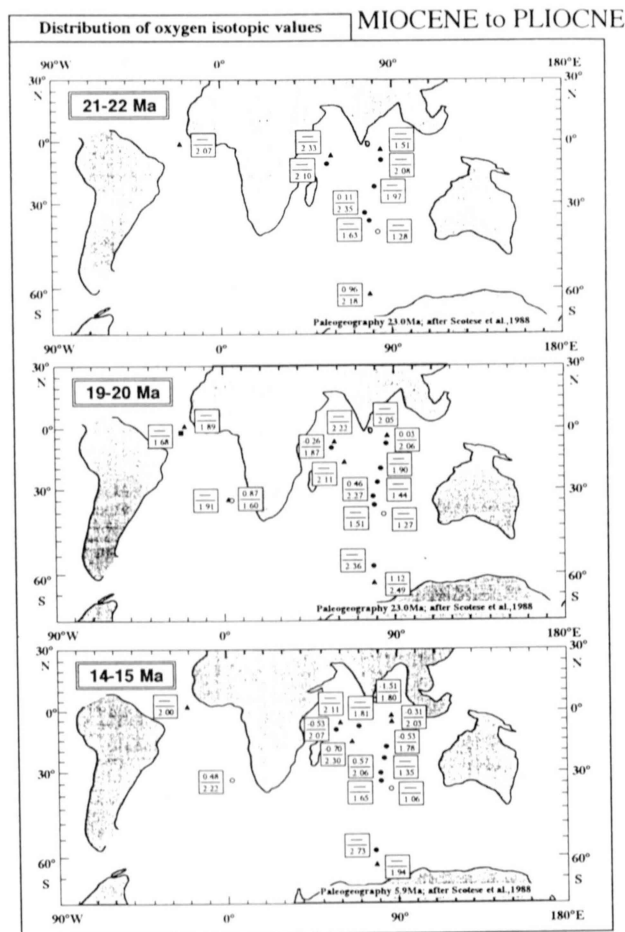


Fig. 59. Geographical distribution of oxygen isotopes from Miocene to Pliocene. The isotopic data are averaged in one million year intervals (Table 7). See legend of Fig. 56.

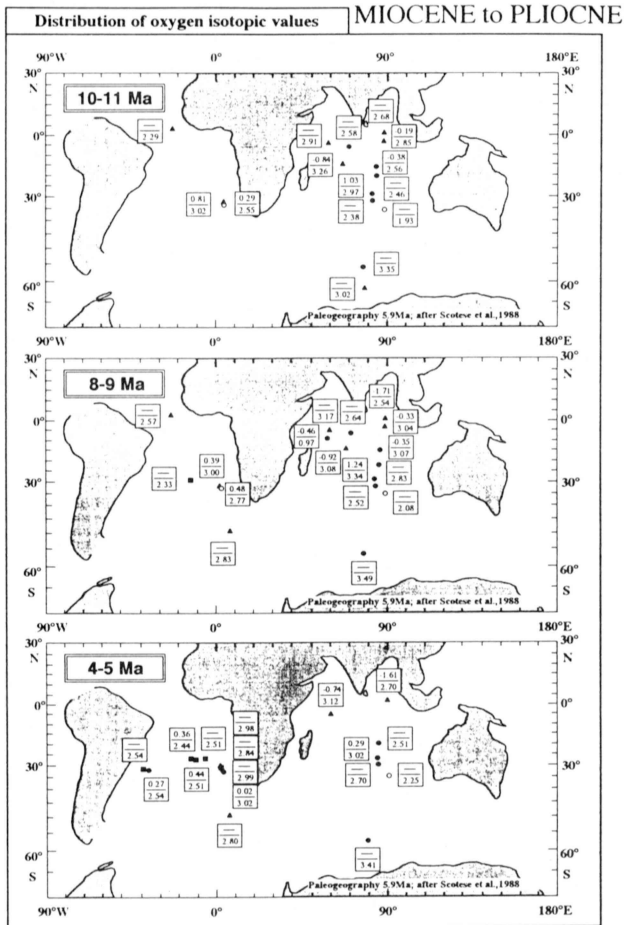


Fig. 59-2. (continued).

with depth in these intervals. In contrast, the latitudinal difference of $\delta^{18}\text{O}$ (around 35°S-5°N) is small at 20-13 Ma (<0.25‰) and in terms of the $\delta^{18}\text{O}$ water column is constant between 1000-2500m depths.

The oxygen isotopic values of bottom water in the Indian Ocean exhibit a general decrease northward from the high-latitude side and westward from the east side during the Miocene. The latitudinal difference of $\delta^{18}\text{O}$ values between 60°S and 35°S is from 0.5 to 1.5‰ at a water depth of 1500 m, and this difference is small at 26-23 Ma, 18-15 Ma, 13-12 Ma, and 9-8 Ma intervals, whereas it is large from 15-13 Ma (Fig. 59). The 18-15 Ma interval corresponds to the oxygen isotopic minima zone in the Miocene, and the 15-13 Ma interval corresponds to a shift in the middle Miocene. $\delta^{18}\text{O}$ values at 1500 m paleodepth between 55°E to 85°E show a range of 0.2-0.8‰ at the 15-9 Ma interval. This difference is not observed at other intervals in the Miocene. At a water depth of 2000m, the latitudinal difference between 60°S and 10°S in terms of $\delta^{18}\text{O}$ value is less than 0.5‰, and is relatively small at 22-20 Ma and 18-15 Ma. The latitudinal difference between 60°S and 10°S shows an increase north-

ward from high-latitudes from 15 to 13 Ma. Relations are reverse to the general gradients. Therefore, the latitudinal gradient of deep water is in contrast to that of shallow water (1500m). The longitudinal difference of $\delta^{18}\text{O}$ at 2000m paleodepth between 55°E and 85°E is consistently 0.1-0.2‰ throughout the Miocene. The latitudinal difference of $\delta^{18}\text{O}$ in surface water is large in the Indian Ocean during the Miocene: for example this difference at 35°S-5°S is 1.9-3.0‰. This difference is largest at 9-8 Ma, which corresponds to the ^{18}O maximum at 8 Ma. The longitudinal difference of $\delta^{18}\text{O}$ values in surface water during the Miocene gradually expands to the ^{18}O maximum at 8 Ma, and reduces afterward.

Pliocene-Pleistocene: The oxygen isotopic latitudinal gradient from the Pliocene to Pleistocene in the Atlantic Ocean decreases south-ward at depths shallower than 4000 m, opposite to the pattern in the Miocene (Fig. 59-2). The latitudinal differences of $\delta^{18}\text{O}$ between 50°S-30°S and 50°S-20°N at a water depth of 2500 m are less than 0.2‰ in the shift of the late Pliocene (5-3 Ma), whereas they expand to 0.5 and 1.2‰ respectively after the shift of the late Pliocene (3-0 Ma). The latitudinal gradient of $\delta^{18}\text{O}$ below the 4000m paleodepth in the Atlantic Ocean decreases to the north. The latitudinal difference in oxygen isotopes between 30°S and 0° is ~0.6‰ at 4500m depth. The oxygen isotopic ratios in surface water tend to increase toward high-latitude, and the latitudinal difference between 0° and 30°S is ~1.0‰. In the Indian Ocean, $\delta^{18}\text{O}$ values of bottom water at depths <3000m tend to decrease north-ward from the Pliocene on, which is the opposite relation to the gradient in the Atlantic Ocean. The oxygen isotopic values in the Indian Ocean are lower than those of the Atlantic Ocean.

2. Carbon isotopes

Maastrichtian: In the Maastrichtian, the carbon isotopic values at Site 752 in the Indian Ocean are ~-0.4‰, lower than those of Site 689 in the Atlantic sector of the Southern Ocean. Although Sites 689 and 690 are situated at different water depths, $\delta^{13}\text{C}$ values at Site 689 are close to those of Site 690. Therefore, the carbon isotopic column in the Atlantic sector of Southern Ocean may be uniform.

Paleocene: In general, carbon isotopic values at the low-latitudes in the Atlantic Ocean are higher than those at the high-latitudes (Fig. 60). The ^{13}C latitudinal gradient is relatively large from 66-61 Ma: for example the latitudinal difference between 55°S and 40°S is ~-1.0‰ at a water depth of 2000m. In this interval, a water column gradient is not observed at low-latitudes. In contrast, the latitudinal difference in carbon isotopes between 70°S and 40°S is small at 61-58 Ma, <0.3‰ at 1000-2500m water depth. At the 61-60 Ma section, $\delta^{13}\text{C}$ values at mid-latitudes (60°S-40°S) are uniform at 1000-2500 m paleodepth, however the $\delta^{13}\text{C}$ values at high-latitudes are relatively low. The ^{13}C difference between those regions and Site 690 at 70°S are relatively large. At the 60-59 Ma interval, which show the highest values of $\delta^{13}\text{C}$, they are uniform in the water mass between 70° and 55°S in the 1500-2500 m paleodepth, and between 55°S and ~40°S at 1000 m. The ^{13}C difference of these water masses, however, is relatively large. The surface carbon isotopic values at 61-58 Ma interval are relatively high at low-latitudes.

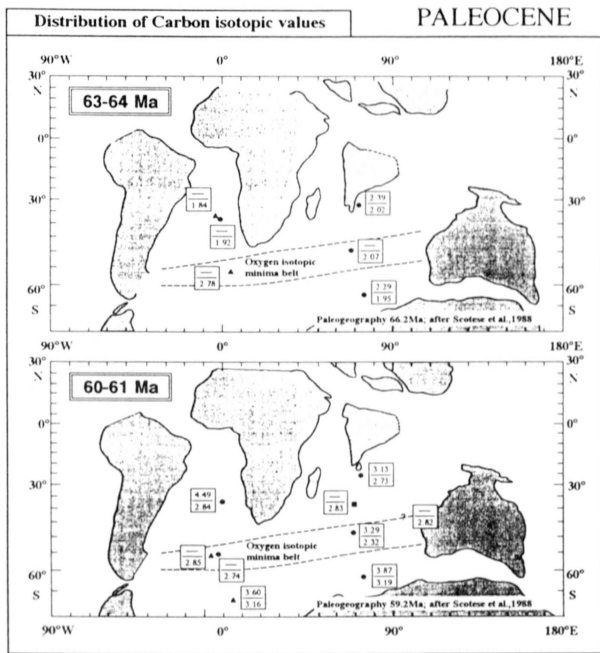


Fig. 60. Geographical distribution of carbon isotopes during the Paleocene. The isotopic data are averaged in one million year intervals (Table 7). See legend of Fig. 56.

In the Indian Ocean, the carbon isotopic latitudinal difference at ~1000 m between 65°S 50°S is ~0.5‰ from 67 to 64 Ma, with higher values at high-latitudes. At 64-61 Ma, $\delta^{13}\text{C}$ values of bottom water between 65°S and 30°S are uniform. In this area, $\delta^{13}\text{C}$ values are the lowest at 50°S. In this interval, $\delta^{13}\text{C}$ values of surface water at mid-latitudes (~50°S) are 0.1-0.4‰, higher than those on the low- and high-latitude sides (around 30°-60°S). Consequently, the surface to bottom difference is the largest at mid-latitudes. In the 61-58 Ma interval for highest values of $\delta^{13}\text{C}$, the latitudinal difference of $\delta^{13}\text{C}$ at ~1000m is relatively large: $\delta^{13}\text{C}$ values on the high-latitude side (~60°S) are ~0.5‰ higher than those on the low-latitude side (~30°S), and ~0.9‰ higher than those of mid-latitudes (~50°S), which are the lowest values in this area. The carbon isotopic values in surface water decrease toward low-latitudes, and the ^{13}C latitudinal difference between 60° and 50°S is relatively large (approximately 0.75‰). In contrast, $\delta^{13}\text{C}$ values between 50°-30°S exhibit a relatively small difference (<0.2‰). In general, the carbon isotopic values on the Atlantic Ocean side are slightly higher than those of the Indian Ocean side at high-latitudes.

Eocene: The carbon isotopic values of bottom water on the Atlantic Ocean side decrease toward low-latitudes through the Eocene (Fig. 61). In the early Eocene, the latitudinal difference of $\delta^{13}\text{C}$ in bottom water tends to expand from 58 to 56 Ma, then reduce to 52 Ma (Fig. 61-1). The largest latitudinal difference of $\delta^{13}\text{C}$ at 57-56 Ma corresponds to a ^{13}C minimum value, and the latitudinal difference between 70 and 40°S at 1000m paleodepth reaches 1.1 ‰. The carbon isotopic column gradient of bottom water in the 58-56 Ma

interval shows a decrease with depth, whereas the gradient increases during the 56-52 Ma interval. This gradient is remarkable at mid-latitudes (~55°S). No surface latitudinal difference in $\delta^{13}\text{C}$ value is observed in the 57 to 55 Ma interval. As for other intervals in the early Eocene, $\delta^{13}\text{C}$ values in surface water show slight latitudinal differences, and are randomly distributed. The benthic latitudinal gradient of $\delta^{13}\text{C}$ on the Indian Ocean side shows a decrease toward low-latitudes similar to that on the Atlantic Ocean side. The latitudinal difference in $\delta^{13}\text{C}$ tends to expand from 58 to 56 Ma. Carbon isotopic values decrease from the Atlantic to Indian Ocean sides at high-latitudes. A regional difference in $\delta^{13}\text{C}$ is recognized in the Indian and South Atlantic Oceans during the ^{13}C minima zone (57-55 Ma). The $\delta^{13}\text{C}$ variance increases between the southern and northern parts of the South Atlantic Ocean. The same feature can be observed in $\delta^{13}\text{C}$ between the South Atlantic and Indian Oceans.

In the 52-49 Ma interval of the middle Eocene, the carbon isotopic values of intermediate water (~2000 m) in the South Atlantic Ocean (70-55°S latitude) are relatively high. A ^{13}C latitudinal difference is not observed in this area. In the 49-44 Ma interval on the Atlantic Ocean side, the carbon

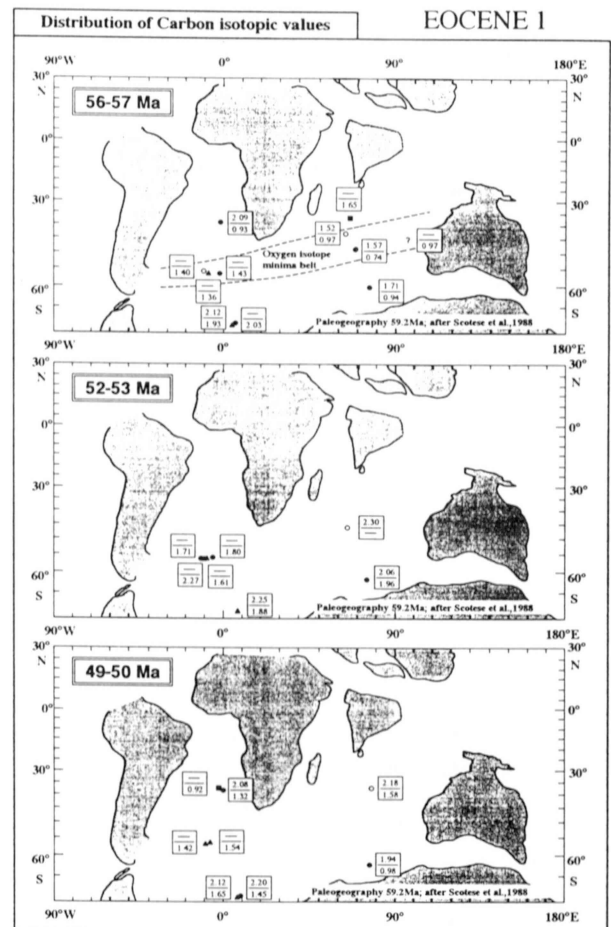


Fig. 61-1. Geographical distribution of carbon isotopes during the Eocene. The isotopic data are averaged in one million year intervals (Table 7). See legend of Fig. 56.

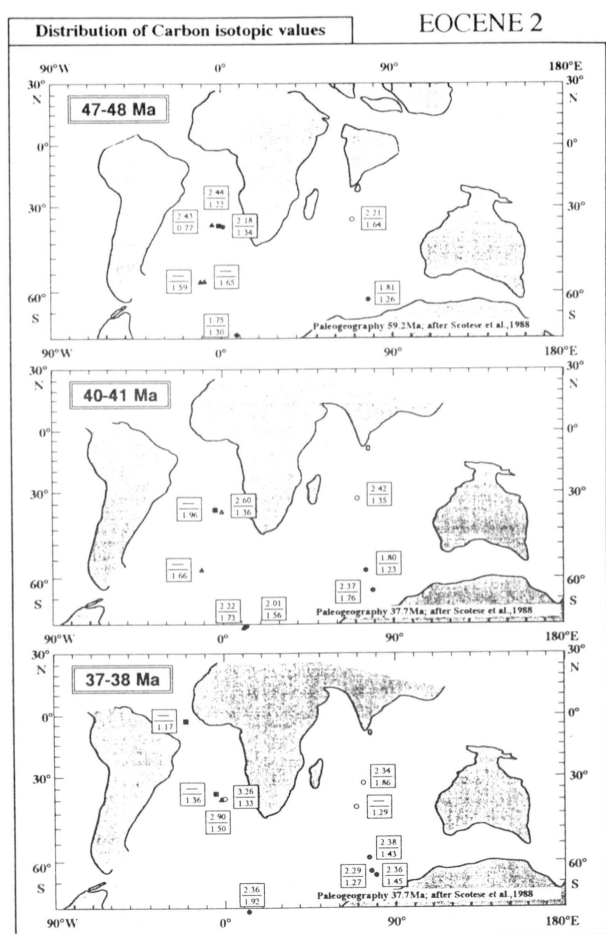


Fig. 61-2. (continued).

isotopic values are lowest at low-latitudes ($\sim 40^{\circ}\text{S}$) in deep water (2500-3000 m paleodepth). This trend is noted in the 48-46 Ma interval (Fig. 61-2). The ^{13}C latitudinal difference is relatively small except for the area showing the lowest value, and is less than 0.6‰ in this interval. In the 44-40 Ma interval, $\delta^{13}\text{C}$ values differ by $<0.6\text{‰}$ according to latitude. In surface water, a latitudinal difference in $\delta^{13}\text{C}$ between 70°S and 40°S in the middle Eocene is not recognized in the 52-49 Ma interval, whereas isotopic differences are $0.55\text{-}0.75\text{‰}$ in the 49-40 Ma interval, with higher values at low-latitudes. In the Indian Ocean, carbon isotopic values in low-latitudes tend to be higher than those at high-latitudes at different water depths through the middle Eocene. The ^{13}C latitudinal difference in the 43-40 Ma interval is $0.4\text{-}0.6\text{‰}$ in intermediate water ($\sim 1000\text{ m}$ paleodepth). In this interval, the carbon isotopic values of deep water at high-latitude ($>1500\text{ m}$) are the highest among the Indian Ocean sites, while the lowest values are found at intermediate water ($\sim 1000\text{ m}$). At high-latitudes, the carbon isotopic values of deep water ($>1500\text{ m}$ paleodepth) in the Indian Ocean are close to those of the Atlantic Ocean during the middle Eocene. Between 60°

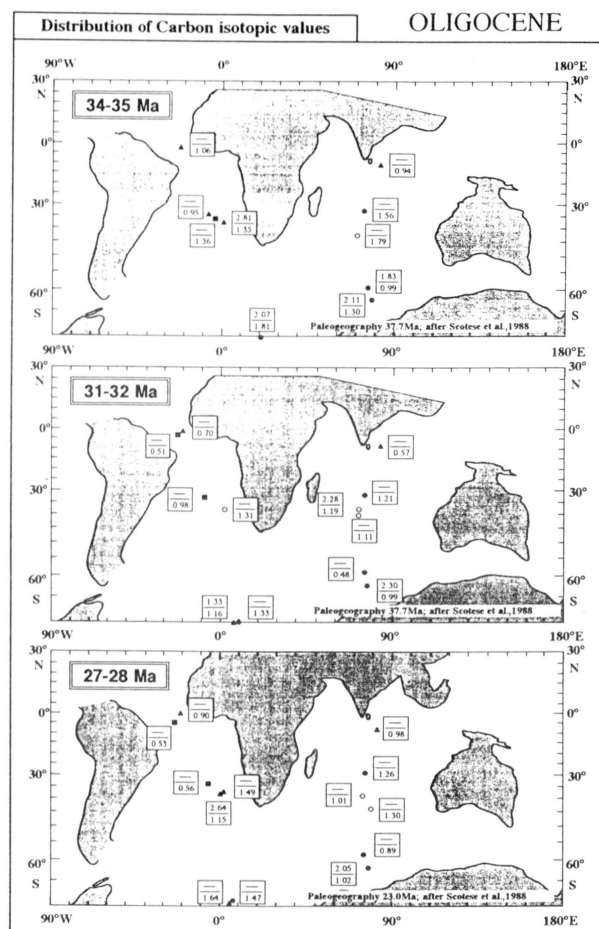


Fig. 62. Geographical distribution of carbon isotopes during the Oligocene. The isotopic data are averaged in one million year intervals (Table 7). See legend of Fig. 56.

and 40°S , $\delta^{13}\text{C}$ values of surface water at low-latitudes tend to be slightly higher than those at high-latitudes.

In the late Eocene, the ^{13}C latitudinal difference ($70\text{-}0^{\circ}\text{S}$) in the Atlantic Ocean is $<0.8\text{‰}$ in bottom water. $\delta^{13}\text{C}$ values in intermediate water ($\sim 1500\text{ m}$) are lower than those of deep water ($\sim 2000\text{ m}$) at high-latitudes ($\sim 70^{\circ}\text{S}$). The carbon isotopes in the water column are close to constant at mid-latitudes ($\sim 40^{\circ}\text{S}$). $\delta^{13}\text{C}$ values in surface water at mid-latitudes ($\sim 40^{\circ}\text{S}$) are $0.7\text{-}1.0\text{‰}$, higher than those at high-latitudes ($\sim 70^{\circ}\text{S}$). In the Indian Ocean, the carbon isotopic at low-latitude are higher than those at high-latitudes at shallower than 2000 m paleodepth throughout the late Eocene. The latitudinal difference of $\delta^{13}\text{C}$ at that time is $0.1\text{-}0.4\text{‰}$ between 60°S and 35°S , which is smaller than that in the late middle Eocene. Carbon isotopic values in the Indian Ocean are lower than those in the Atlantic Ocean, and the difference of $\delta^{13}\text{C}$ in those oceans tends to expand throughout the middle Eocene.

Oligocene: Carbon isotopes on the Atlantic Ocean side shows decrease gradually toward low-latitudes (Fig. 62).

The ^{13}C latitudinal difference between 70°S and 0° at a water depth of $\sim 2000\text{m}$ is $0.6\text{-}1.0\text{‰}$ up to $\sim 32\text{ Ma}$, and $0.4\text{-}0.6\text{‰}$ from 32 Ma to 26 Ma . This difference tends to expand from 30 Ma . The latitudinal difference of $\delta^{13}\text{C}$ at $2500\text{-}3500\text{ m}$ between 70°S and 0° is $0.8\text{-}1.1\text{‰}$ in the $32\text{-}26\text{ Ma}$ interval with an increase in upper stratigraphic levels. Surface carbon isotopes exhibit a high values at low-latitudes. Carbon isotopic values of intermediate water ($<1500\text{ m}$ paleodepth) on the Indian Ocean side increase toward low-latitudes. The ^{13}C latitudinal difference reduces upward, and is $<0.2\text{‰}$ in the $27\text{-}24\text{ Ma}$ interval. Conversely, carbon isotopic values in deep water ($>1500\text{ m}$ paleodepth) decrease toward low-latitudes. At high-latitudes ($\sim 60^\circ\text{S}$), $\delta^{13}\text{C}$ values in intermediate water ($\sim 1000\text{ m}$) are lower than those in deep water during the Oligocene. Although this trend cannot be found in low-latitudes by 31 Ma , $\delta^{13}\text{C}$ values are low at $\sim 1000\text{m}$ after 31 Ma , with high values at $\sim 2000\text{m}$ paleodepth.

Miocene: During the Miocene, carbon isotopes in bottom water on the Atlantic Ocean side decrease toward low-latitudes (Fig. 63). The latitudinal difference of $\delta^{13}\text{C}$ at 2500m between 35°S and 5°N is less than 0.4‰ in the $24\text{-}22\text{ Ma}$ interval, corresponding to a ^{13}C maximum in the earliest

Miocene (23 Ma). A value of $0.5\text{-}0.7\text{‰}$ at $19\text{-}14\text{ Ma}$ corresponds to a ^{13}C maximum in the middle Miocene (15 Ma), whereas there are large variation between 0.8 to 1.0‰ from $21\text{-}19\text{ Ma}$. From 8 to 14 Ma , the ^{13}C latitudinal difference between 35°S and 5°N at 2500m paleodepth tends to reduce from 0.8 to 0.2‰ . The water column gradient in the $1000\text{-}3000\text{ m}$ depth interval is small ($<0.3\text{‰}$). $\delta^{13}\text{C}$ values at $>3000\text{m}$ differ by $\sim 0.5\text{‰}$ from this depth section, particularly during the late Miocene.

The latitudinal gradient of carbon isotopic values in the Indian Ocean decrease toward low-latitudes at the same water depth. Carbon isotopic values in the Indian Ocean transect show a longitudinal decrease eastward from 24 Ma to 9 Ma , whereas they increase eastward from 9 Ma to 5 Ma . In the water column of the northern Indian Ocean, lower $\delta^{13}\text{C}$ values are observed at $1500\text{-}2000\text{m}$, and $\delta^{13}\text{C}$ values are relatively high immediately above and below this interval. The ^{13}C latitudinal gradient in surface water tends to decrease northward from high-latitudes. The ^{13}C latitudinal difference is large up to 19 Ma , and then is greater than 1.0‰ between 60° and 5°S . After 19 Ma , the ^{13}C latitudinal difference is relatively small during the Miocene.

Pliocene-Pleistocene: During the Pliocene, carbon

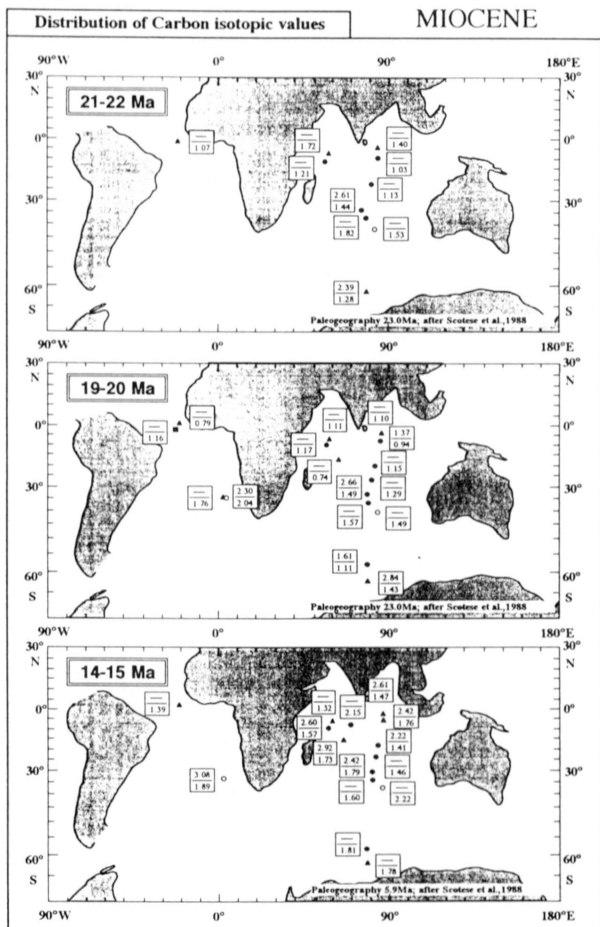


Fig. 63-1. Geographical distribution of carbon isotopes from Miocene to Pliocene. The isotopic data are averaged in one million year intervals (Table 7). See legend of Fig. 56.

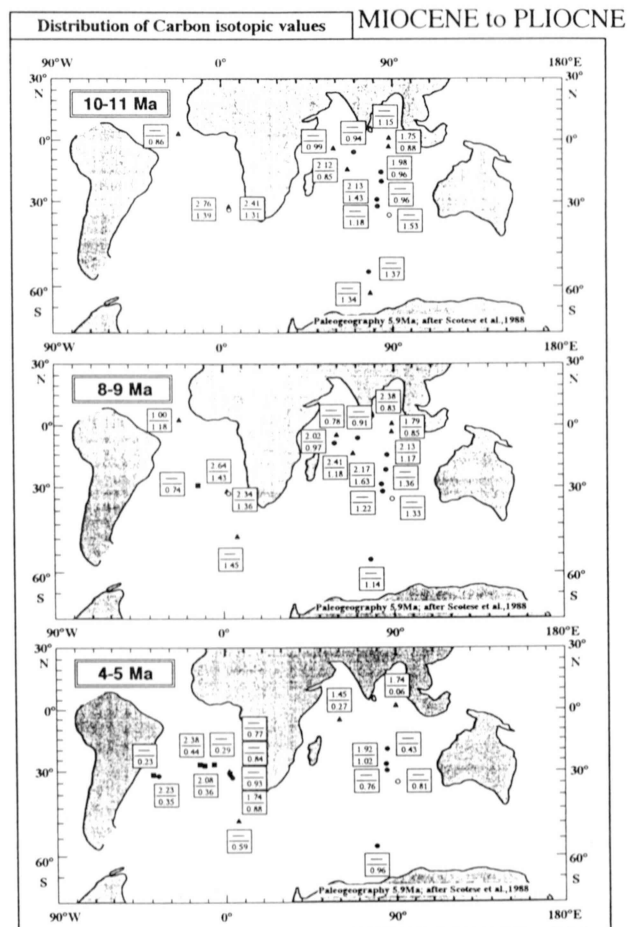


Fig. 63-2. (continued).

isotopic values on the Atlantic Ocean side tend to decrease northward. At 2500m water depth, $\delta^{13}\text{C}$ values around 30°S are 0.1-0.2‰, higher than those at 50°S. The latitudinal difference at 2000m between 30°S and 20°N is $\sim 0.8\text{‰}$. This difference decreases throughout the Pleistocene. The water column gradient of $\delta^{13}\text{C}$ is relatively large, and $\delta^{13}\text{C}$ values at $>4500\text{m}$ paleodepth are particularly low. From 30°S-20°N, the surface carbon isotopic values on the south side are 0.8-1.5‰, higher than the north side, with low values in the young part. During the Pleistocene and Pliocene, the gradient of carbon isotopic values in the Indian Ocean decreased a northward at high-latitudes. In the 5 to 4 Ma interval, $\delta^{13}\text{C}$ values between 60°S and 25°S are uniform ($<0.25\text{‰}$). In general, the carbon isotopic values at depth $>2500\text{m}$ are low during the Pleistocene and Pliocene.

IV. Paleoceanographic reconstruction of the Indian and South Atlantic Oceans

A. Paleo-ocean circulation viewed from carbon and oxygen isotopic ratios

Carbon isotopic ratios in sea water are mainly influenced by the biological production of organic carbon and the decomposition of organic matter. $\delta^{13}\text{C}$ values in surface water are generally 1‰ higher than those of bottom water (Kroopnick, 1974; 1980; Berger and Vincent, 1986). This is caused by selective absorption of ^{12}C -enriched CO_2 by phytoplankton in surface water. On the other hand, Kroopnick (1974, 1980) investigated the $\delta^{13}\text{C}$ values of DIC (dissolved inorganic carbon) of the modern ocean, and noted that the $\delta^{13}\text{C}$ values depend on the circulation velocity of deep water. Modern $\delta^{13}\text{C}$ distribution of DIC in the Atlantic Ocean show a southward decrease from the Northern hemisphere (Kroopnick, 1974; 1980; Leonard et al., 1983; Duplessy et al., 1984; Woodruff and Savin, 1989). This gradient is explained by a supply of ^{12}C by oxidative respiration of ^{12}C -enriched organic carbon through flow of deep water. $\delta^{13}\text{C}$ values in aged deep water are low because of high ^{12}C . Therefore, a ^{13}C gradient indicates the deep water flow (Woodruff and Savin, 1989).

$\delta^{13}\text{C}$ values of DIC have been recorded in the carbonate of foraminiferal tests (e.g., Shackleton et al., 1984; Woodruff and Savin, 1989). Woodruff and Savin (1989) showed that the distribution pattern of foraminiferal $\delta^{13}\text{C}$ are similar to the modern distribution of water DIC $\delta^{13}\text{C}$, and they concluded that the benthic $\delta^{13}\text{C}$ gradient can be used to tracer ancient ocean circulation. Based on this evidence, they proposed the existence of Tethyan Indian Saline Water (TISW), flowing from the Tethys into the northern Indian Ocean.

In this study, the paleo-directions of deep flow are deduced from the foraminiferal $\delta^{13}\text{C}$ gradient. In the case of a single direction of deep-water flow, the relationship between $\delta^{13}\text{C}$ values and migration distance are characterized by a linear function (Fig. 64-A). If this relationship is not recognized, the geographical distribution of $\delta^{13}\text{C}$ may represent the some water mass flow (Fig. 64-B).

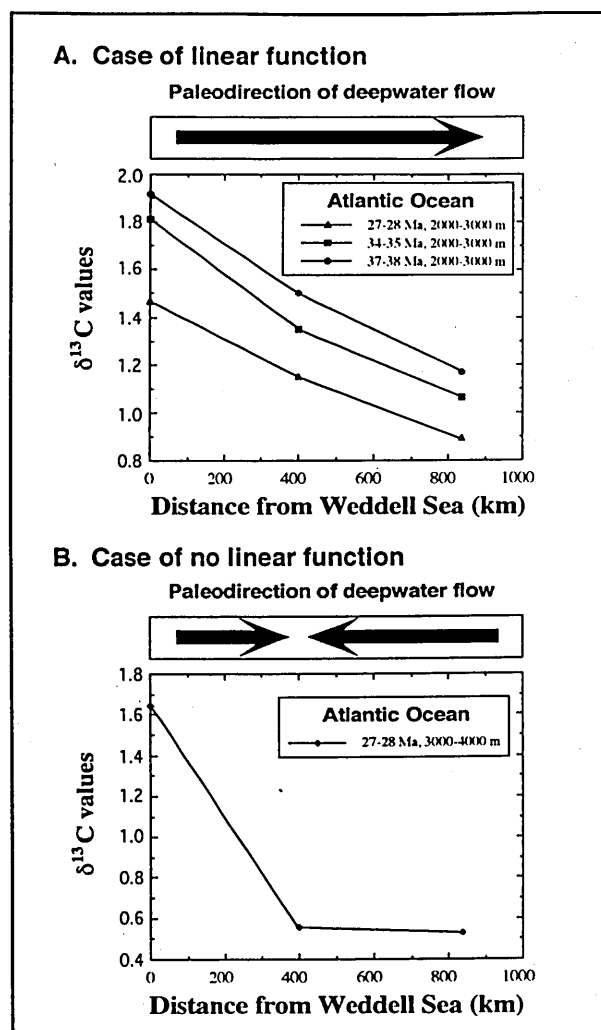


Fig. 64. Paleodirection of deep water flow and the ^{13}C gradient. A. In the case of a linear function, the paleodirection indicates a single direction flow. B. For no linear function, the paleodirection indicates some water-mass flow.

B. Isotopic water of column structure

The water column structure of oxygen and carbon isotopes from the late Maastrichtian to Pleistocene has been reconstructed, shown in Figs. 65 and 66. The relationship between the estimated water depth and isotopic data at representative ages are shown in Figs. 67 and 68. Each reconstruction covers a relatively large area (i.e., for the northeastern Indian Ocean region; the latitudinal range is about 35°), therefore even small regional differences can be recognized. The reconstruction of the Neogene in the southern Indian and Southern Oceans cannot be illustrated because of few data.

1. Oxygen isotopes

During the Paleocene, oxygen isotopic ratios of the northeastern Indian Ocean increased in a linear fashion with depth down to 1500m, and become constant below this level. The differences between the $\delta^{18}\text{O}$ values of surface and

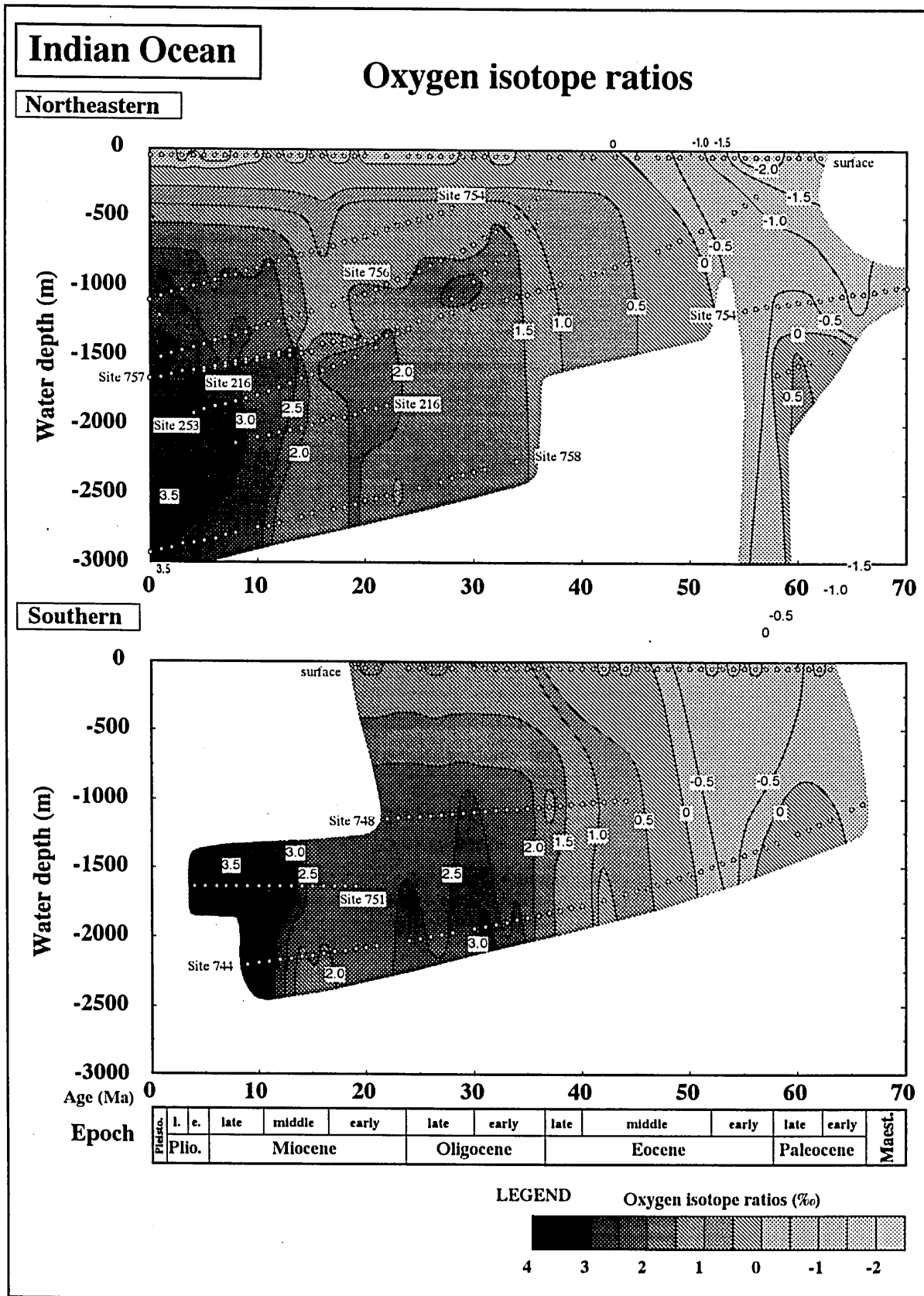


Fig. 65-1 DIC $\delta^{18}\text{O}$ value of water column at the Cenozoic in the four ocean region. The isotopic data are averaged in one million year intervals (Table 7). Paleodepth is mainly reconstructed using published data calculated from a subsidence curve of the "backtrack method" (See III-A-3 section).

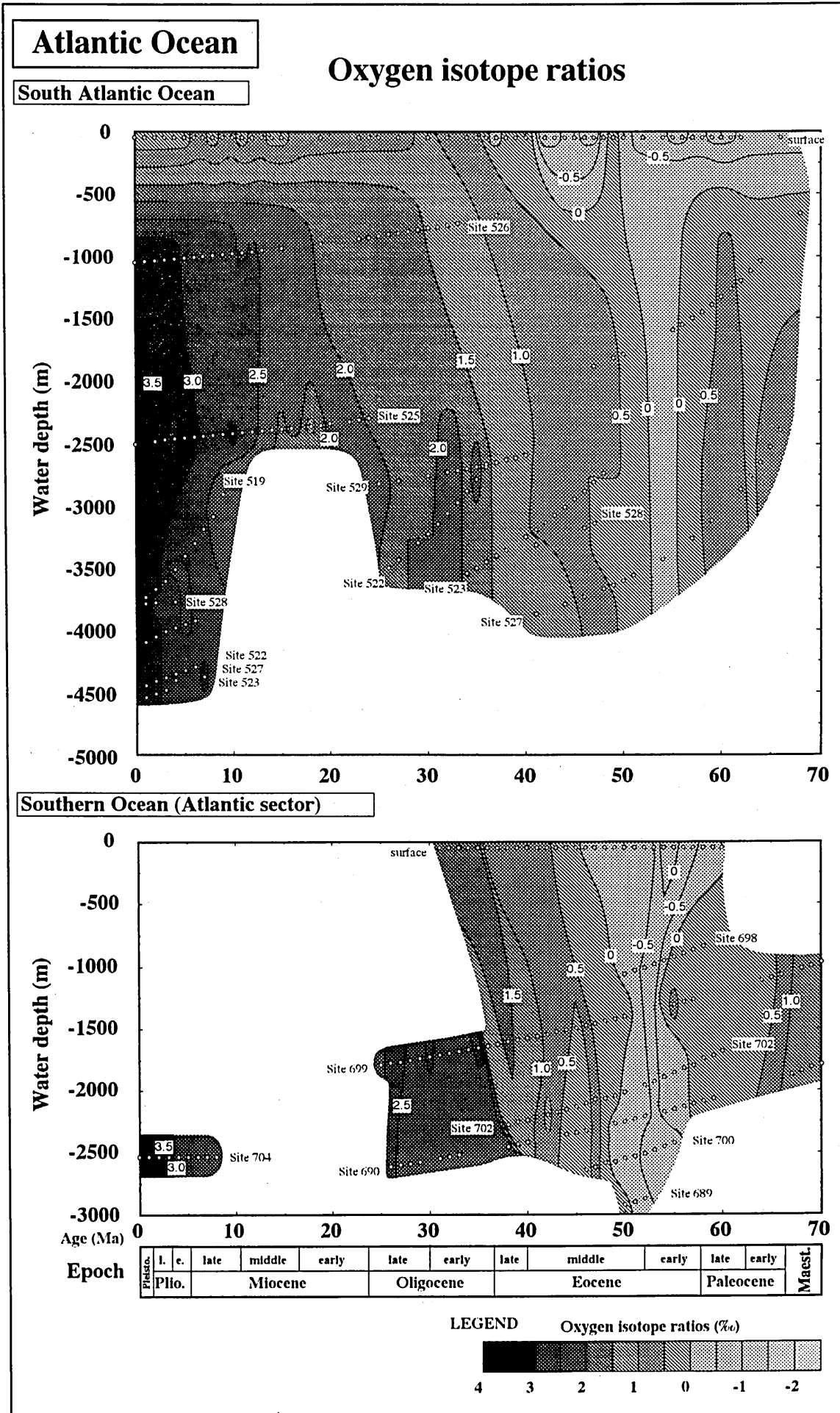


Fig. 65-2. (continued).

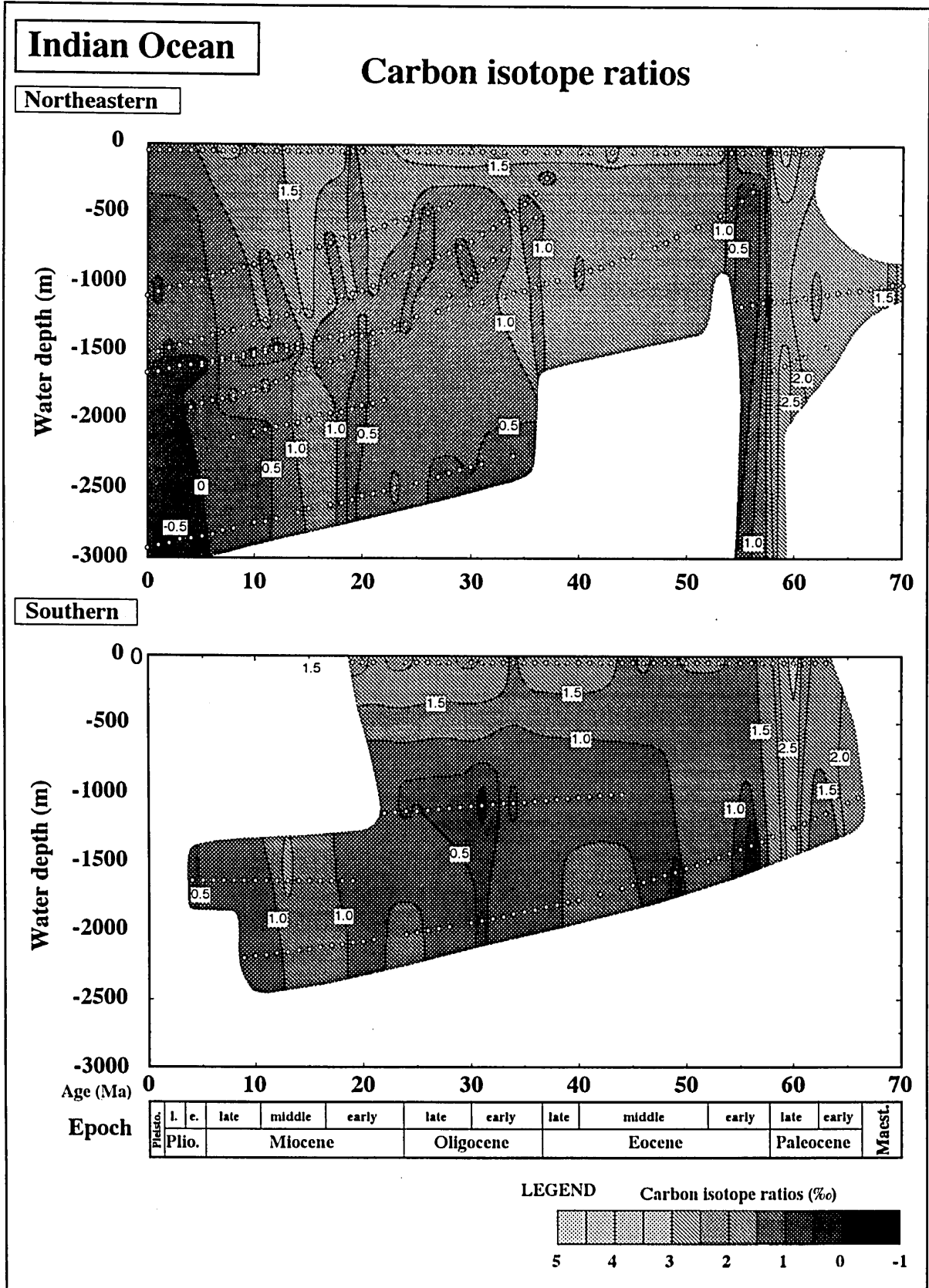


Fig. 66-1 DIC $\delta^{13}\text{C}$ value of water column in the Cenozoic in the four ocean region. The isotopic data are averaged in one million year intervals (Table 7). Paleodepth is mainly reconstructed using published data calculated from a subsidence curve of the "backtrack method" (See III-A-3 section).

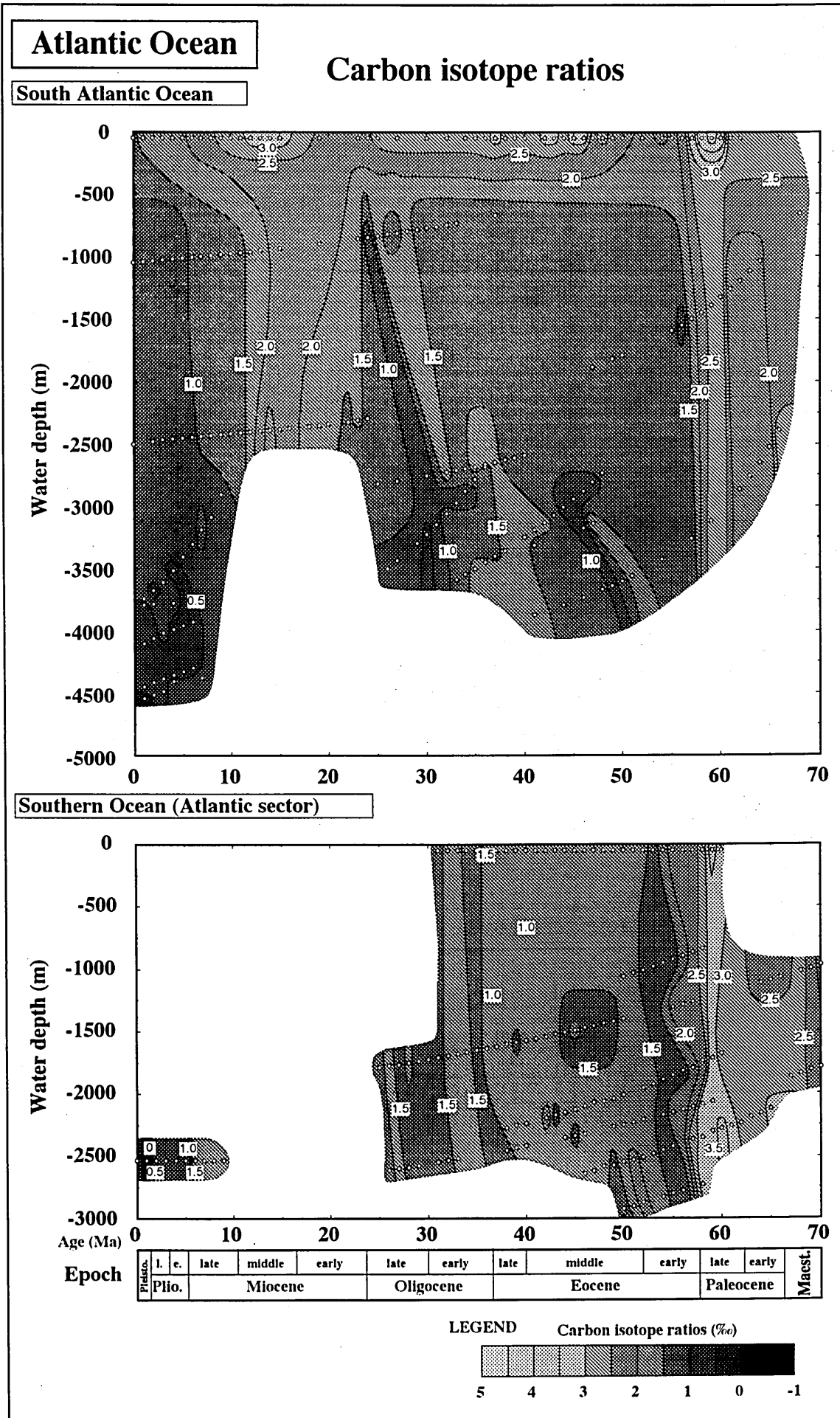


Fig. 66-2. (continued).

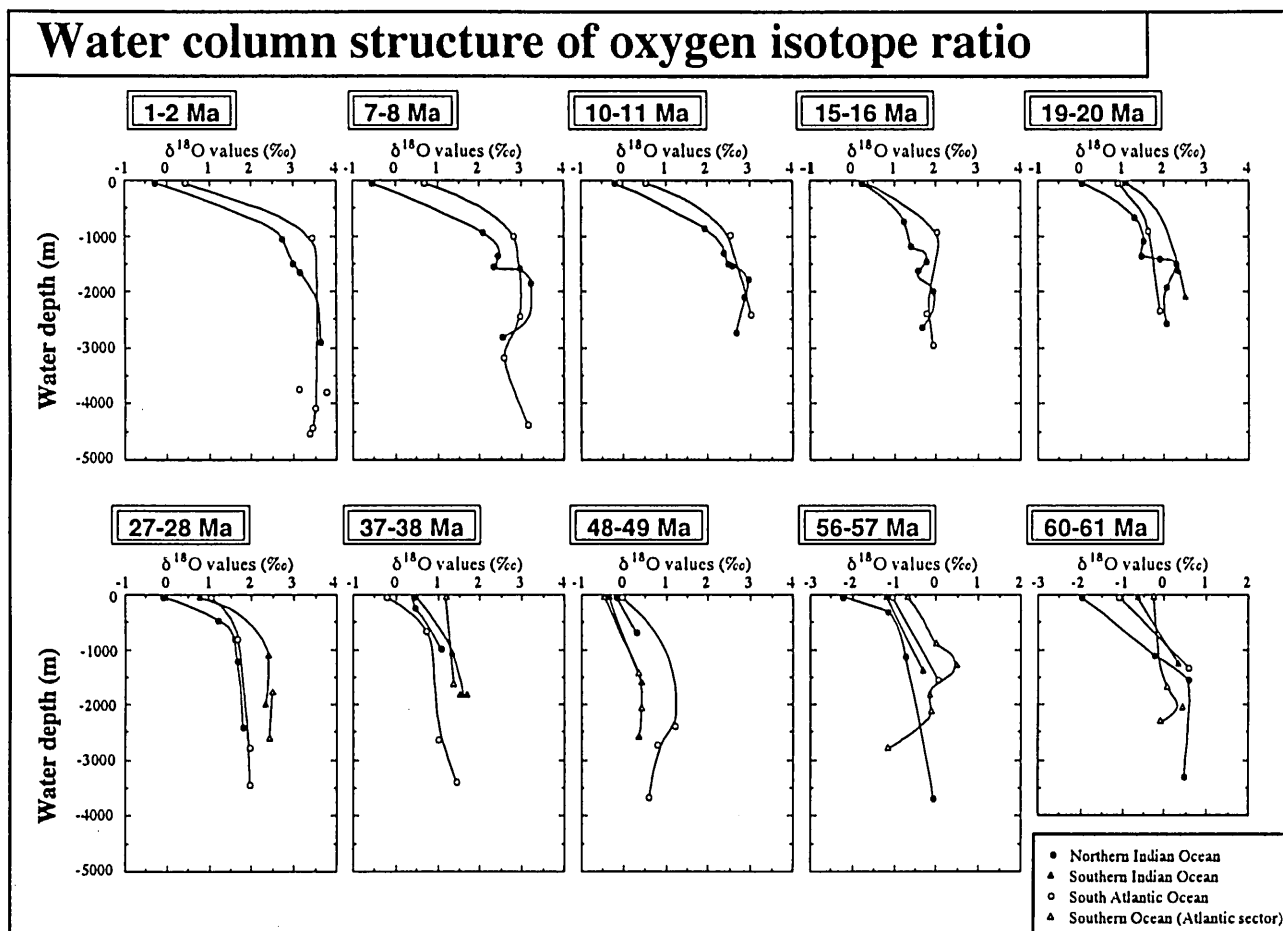


Fig. 67 Vertical change of DIC $\delta^{18}\text{O}$ value at representative age in the four ocean region. The isotopic data are averaged in one million year intervals (Table 7). Paleodepth is mainly reconstructed by published data calculated from a subsidence curve of the "backtrack method" (See IV-A-3 section).

bottom water are relatively large (up to $\sim 2.5\%$). The water columns in other oceans exhibit the same gradient to that of the northeastern Indian Ocean except for the Southern Ocean (Atlantic sector), which shows homogeneous values through the water column. A ^{18}O maximum at 61 Ma is observed in all the oceans except for the Southern Ocean (Atlantic sector), and is particularly high at the northeastern Indian Ocean region. The surface to bottom difference in oxygen isotopes of the northeastern Indian Ocean during this time is the largest in the Paleocene section.

In the ^{18}O minimum zone within the early Eocene, $\delta^{18}\text{O}$ values are low through the entire water column, and the surface to bottom difference is relatively small. In the northeastern Indian Ocean, bottom water oxygen isotopes increase with increasing depth. During the increase of the middle and late Eocene, the surface to bottom difference in oxygen isotopes is relatively small, and the isotopic difference according to depth in bottom water is also small, especially in the Southern Ocean where the entire water column records relatively homogeneous values.

The shift at the Oligocene / Eocene boundary is remarkable in the Southern Ocean. During the Oligocene, $\delta^{18}\text{O}$ values rapidly increased from the surface to depths of about 1000m, and gradually increased from 1000m.

In the early Miocene, the water column in the northeastern Indian Ocean was divided into intermediate and deep water by a conspicuous discontinuity at about 1500m. $\delta^{18}\text{O}$ values of intermediate water are $\sim 0.6\%$, higher than those of deep water. The difference between $\delta^{18}\text{O}$ values of surface and intermediate water during this time is similar to that of Oligocene sections. In the ^{18}O minima zone of the Miocene, the intermediate/deep water boundary is indistinct, and the surface to bottom difference reduces. After the shift of the middle Miocene, this conspicuous discontinuity is again observed around 1500m, and the surface to bottom difference increases. All sections of bottom water show high values at the 8 Ma ^{18}O maximum. From the latest Miocene to early Pliocene, $\delta^{18}\text{O}$ values are distinctively high in the water column from 1600-2500m paleodepth. These discontinuities are only recognized in the northern Indian Ocean. During the Miocene, the water column gradient in the South Atlantic

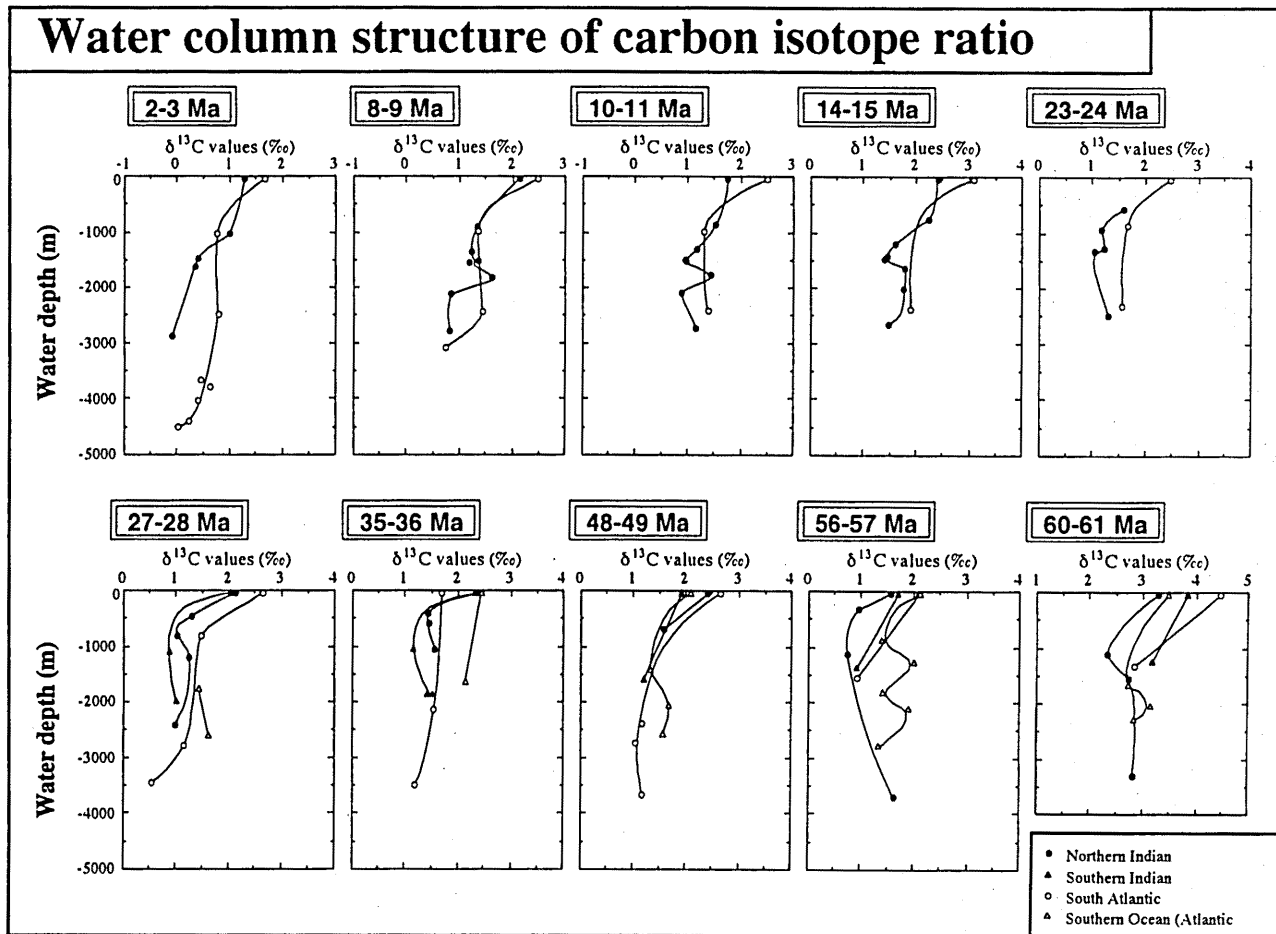


Fig. 68 Vertical change of DIC $\delta^{13}\text{C}$ value at representative age in the four ocean region. The isotopic data are averaged in one million year intervals (Table 7). Paleodepth is mainly reconstructed by published data calculated from a subsidence curve of the "backtrack method" (See IV-A-3 section).

Ocean displays a trend of gradual increase with depth. From the late middle Miocene to early Pliocene, $\delta^{18}\text{O}$ values are relatively high in intermediate water (1000-2800 m).

After the shift of the late Pliocene, the surface to bottom differences in oxygen isotope values expanded in the Indian and South Atlantic Oceans. In bottom water, $\delta^{18}\text{O}$ values tended to increase with depth in the northern Indian Ocean, and were uniform in the South Atlantic Ocean.

2. Carbon isotopes

During the Paleocene, carbon isotopes in the water column of the northeastern Indian Ocean region showed an increase by 0.5‰ from 400 to 1200m, and were uniform below 1200m. In the Southern Ocean (Atlantic sector), the surface to bottom water difference of carbon isotope was relatively small. The surface to bottom water differences, however, were large except for the Southern Ocean (Atlantic sector). The shift at the Paleocene / Eocene boundary is remarkable as observed in other deep sea sections.

The minimum value zone within the early Eocene exhibits a similar pattern in all oceans, as in the Paleocene.

The surface to bottom difference at this time is slightly smaller than those in the Paleocene. Although the intermediate water (500-1500 m) in the southern Indian Ocean shows relatively low values, the carbon isotopic ratios of bottom water are uniform from the minimum value zone in the early Eocene to the shift in the earliest Oligocene in all oceans. The surface to bottom difference is the largest (1.5-2.0‰) in the South Atlantic Ocean. In contrast to this, $\delta^{13}\text{C}$ values of bottom water are close to that of surface water of the Southern Ocean. The surface to bottom difference in the Indian Ocean is 0.5-1.0‰. Although the water column exhibits a general decreasing trend with increasing paleodepth during the late Oligocene, $\delta^{13}\text{C}$ values from 400-1000m in the northeastern Indian Ocean and of 400-1500m in the southern Indian Ocean are relatively low. The surface to bottom difference in this interval is larger than that of the early Oligocene in the northeastern Indian Ocean. At the ^{13}C maximum at 23 Ma (near the Miocene / Oligocene boundary), the entire water column records are roughly homogeneous.

From the Miocene onward, the carbon isotopic values through the water column in the northeastern Indian Ocean show a gradual decrease with depth. At approximately

1500m, carbon isotopic values are remarkably low (0.2-0.5‰) during the Miocene. From 11 to 7 Ma, a discontinuity is observed at a depth of 2000m, and $\delta^{13}\text{C}$ values below 2000m deep reduce by $\sim 0.5\text{‰}$. $\delta^{13}\text{C}$ values of surface water increase from 6 to 10 Ma, consequently the surface-bottom difference increases. From the Pliocene to Pleistocene, a discontinuity with a value of 0.6‰ is recognized at 1300m paleodepth. No discontinuity is recognized in other oceans after the Miocene.

C. Source of deep water

Ocean regions with high $\delta^{13}\text{C}$ values is close to the source of bottom water (Miller et al., 1987) because the ^{13}C -enriched surface water produced from the source region flows into the bottom water. Therefore, the oxygen and carbon isotopic ratio of bottom water in these regions is also close to that of surface water. The oxygen isotopic ratio of bottom water in the source region exhibits the highest value, because this water is characterized by either colder or highly saline waters. Based on water density, the source water may have flowed to a warmer and less saline deep area.

In the South Atlantic and Indian Oceans, the averaged value of $\delta^{13}\text{C}$ in one million year intervals are plotted as a function of paleo-latitude (Fig. 69). Throughout the Cenozoic, carbon isotopic values are high on the high latitude side of the Southern Ocean, and tend to decrease toward middle to low latitudes. In the former region, oxygen isotopic ratios are higher, and the carbon isotopic values of bottom water are also close to those of surface water. This suggests that the source water formed at a high latitude in the Southern Ocean throughout the Cenozoic. Bottom water influenced by this source water may correspond to AABW (Antarctic Bottom Water) or its prototype.

During the Paleocene, $\delta^{18}\text{O}$ and $\delta^{13}\text{C}$ values in the northern part of the South Atlantic and Indian Oceans are also distinctly high. The ^{13}C difference between surface and bottom water is small in the northern part of Indian Ocean (Site 758), which is similar to that of the Southern Ocean (Site 690). This shows that the source water exists in the northern part of the South Atlantic and the Indian Oceans. This water may be warm and highly saline, because this source water probably formed at low latitude. Brass et al. (1982) called such deep water "Warm Saline Deep Water (WSDW)", and considered that this water was formed by the concentration of saline water through evaporation in the Tethyan Sea or lower latitudes. The gradient of carbon isotopic ratios in the Indian Ocean supports the supposition that WSDW may have formed in the Tethyan Sea.

High $\delta^{13}\text{C}$ values are recognized at mid latitudes ($\sim 30^\circ\text{S}$) from the Miocene. This is probably caused by the influence of AAIW (Antarctic Intermediate Water).

During Miocene, the carbon isotopic values are high at 10°S in the northwestern Indian Ocean. This may be influenced by TISW (Tethyan Indian Saline Water) as proposed by Woodruff and Savin (1989).

D. Reconstruction of paleo-ocean circulation

During the Paleocene, two water masses (proto-AABW and WSDW) probably existed in the South Atlantic and Indian Oceans. The gradient of oxygen and carbon isotope ratios shows that Proto-AABW flowed northward

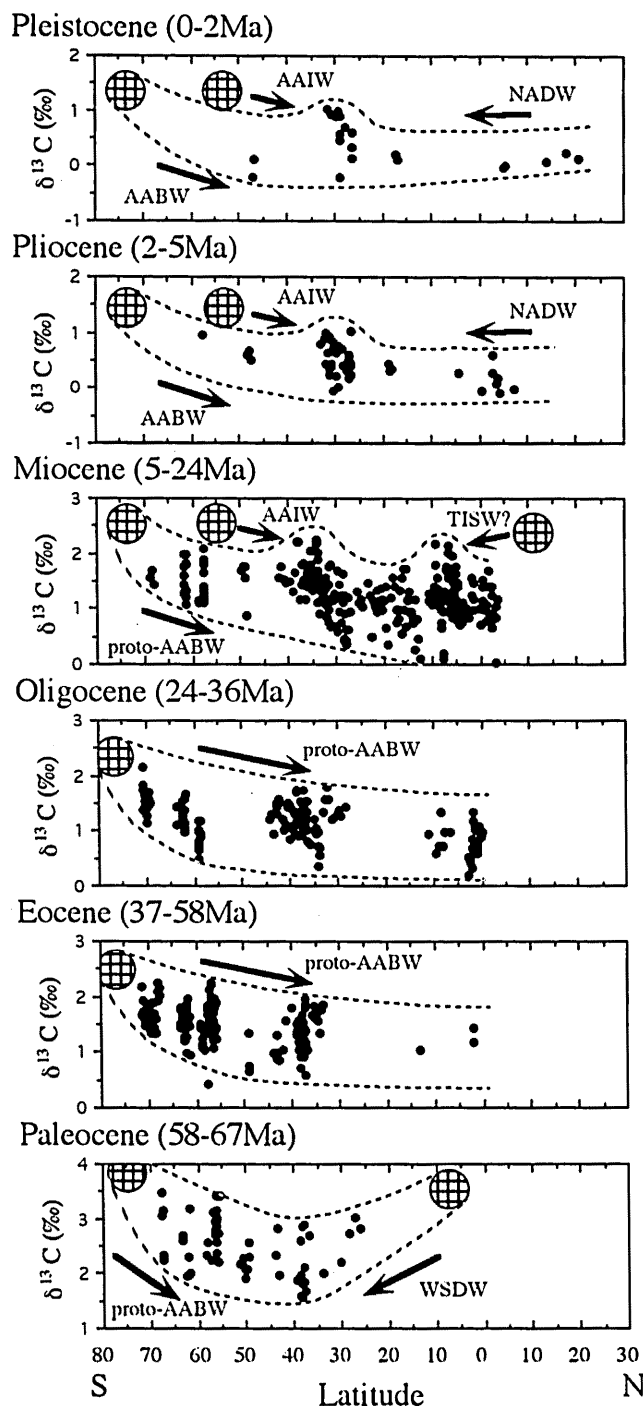


Fig. 69 ^{13}C latitudinal change of one million average values during the each period. Carbon isotopic ratios of source water show a high value. In general, carbon isotopic values decrease from high to low latitudes.

from the Southern Ocean, and that WSDW flowed probably southward from the Tethyan Sea. In the South Atlantic and Indian Oceans during the Paleocene, an "oxygen isotopic minima belt" existed at mid-latitudes (50-60°S). $\delta^{18}\text{O}$ values in this belt are 0.6-1.6‰, lower than that of high- and low latitudes. $\delta^{18}\text{O}$ values at Site 752 are lowest among the "oxygen isotopic minima belt". The "oxygen isotopic minima belt" may have been caused by the confluence of Proto-AABW and WSDW, as pointed out by Nomura et al. (1992), because the carbon isotopic ratio is lowest among the ^{13}C gradient. Across the Paleocene / Eocene boundary, a sudden decrease in carbon isotopic ratios has been recognized from 59 to 56 Ma (Kennett and Stott, 1990; Barrera and Huber, 1991; Zachos et al., 1992a). However, the distribution of oxygen and carbon isotopes is similar to that of the Paleocene. Therefore, the ocean circulation system seems to be similar to that of the Paleocene.

After 56 Ma, the ^{13}C latitudinal gradient shows that Proto-AABW flowed into the South Atlantic and Indian Oceans from high latitudes. Until about 50 Ma, $\delta^{18}\text{O}$ values at Site 757 in the northern Indian Ocean are higher than that at Site 689 in the Southern Ocean. This suggests that WSDW remained an influence in the Indian Ocean until about 50 Ma, and that WSDW flowing from the north part of Indian Ocean disappeared after that. This water mass may have been reduced with the closing of the Tethyan Sea. WSDW on the Atlantic Ocean side, however, may have flowed until about 40 Ma judging from the ^{13}C geographical gradient of the South Atlantic Ocean. The carbon and oxygen isotopic ratios at Site 738, where the paleodepth is slightly deeper than at Site 748, show similar values to those at Sites 689 and 690. This indicates that the water mass flowed into deep water in the southern Indian Ocean from the Southern Ocean. This water mass can be referred to as the proto-type of CPDW (Circumpolar Deep Water).

The sharp change in oxygen isotopic ratios and low carbon isotopic ratios are recognized at ~1000 m depth in the northern Indian Ocean at about 30 Ma. The water mass in the deep part may be equivalent to the proto-type of AABW (Proto-AABW), and the water mass in the shallow part is equivalent to the present AAIW. AAIW may be formed by melting of sea ice, as in the present ocean, based on the lower $\delta^{18}\text{O}$ values.

Carbon isotopic ratios in the northwestern Indian Ocean are higher during the Miocene. This supports the scenario of Woodruff and Savin (1989), where a warm saline plume (Tethyan / Indian Saline Water: TISW) flowed into the northern Indian Ocean from the Tethyan Sea. However, a northward decrease of the carbon isotope gradient is recognized in the intermediate water of the northeastern Indian Ocean. This suggests that TISW is not the main flow in the Indian Ocean. This intermediate water is probably the Antarctic Intermediate Water (AAIW), which formed by melting of sea ice, because the $\delta^{18}\text{O}$ values of intermediate water are low in comparison with that of deep water. The sharp change of oxygen isotopic ratios in the water column of the northern Indian Ocean, assuming that the AAIW / Proto-AABW boundary, are recognized at ~2000 m paleodepth. The boundary is deeper than that recognized in the late Oligocene. During the late Miocene, the AAIW / Proto-AABW boundary deepens to the north, as indicated by the

distribution of low $\delta^{13}\text{C}$ values. This may be caused by the northward movement of Proto-AABW. As a result, AAIW flows into the deeper part of the northern Indian Ocean.

After the late Pliocene, ocean circulation of the deep sea is the same as the present circulation. The oxygen isotopic ratios of Central Atlantic Ocean are higher than those of the South Atlantic Ocean. The carbon isotopic gradient shows a southward decrease in the South Atlantic Ocean. These features indicate a water mass corresponding to NADW in the modern ocean. Before the early Pliocene, positive evidence of NADW is not recognized in the South Atlantic Ocean. This indicates that NADW developed rapidly near the early/late Pliocene boundary (about 3 Ma), and affected to AABW and CPDW.

VI. Concluding Remarks

(1) Oxygen and carbon isotopic ratios were studied in Cenozoic sediments at six sites (Sites 752, 754, 756, 757, 758, and 762) of ODP Legs 121 and 122 in the northeastern Indian Ocean. They recorded such global events as the sharp increase of $\delta^{18}\text{O}$ values near the middle Miocene and Eocene / Oligocene boundary, the increase of $\delta^{18}\text{O}$ values in the Eocene (Miller et al., 1987), the chron-6 shift and the chron-16 shift of $\delta^{13}\text{C}$ ratios (Vincent et al., 1980; 1985), and the drastic change of $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ values across the Paleocene / Eocene boundary.

(2) Benthic and planktonic foraminiferal isotopic were data converted into δ values of dissolved inorganic carbon of marine water on the basis of the adjustment values calculated from foraminiferal interspecific differences in isotopic ratios. Those data are compiled from the Indian Ocean and the South Atlantic Oceans. The general trends of oxygen and carbon isotopic records are similar in all studied oceans. However, the oxygen and carbon isotopic values show an increase southwards.

(3) Averaged isotopic values in one million year intervals are calculated at each ODP and DSDP site, and the time and spatial distributions of the oxygen and carbon isotopic values are examined based on the estimated paleodepth and paleocoordinates. In the Paleocene water column, vertical change in isotopic ratios is not observed. In the Miocene, a notable oxygen isotopic discontinuity and low carbon isotopic values are recognized at ~1500m paleodepth in the northeastern Indian Ocean. This suggests that two water masses may have existed in the Miocene water column of the Indian Ocean. The modern ocean shows a complicated water circulation in the Indian Ocean (Warren, 1984). Based on those results, it is concluded that the water column structure became more complicated from Paleocene to the Recent.

(4) The source of deep water is discussed using the ^{13}C geographical gradient. In the South Atlantic and Indian Oceans, the source water may be formed on the high latitude side of the Southern Ocean through the Cenozoic because the carbon isotopic values are consistently high on the high latitude side of the Southern Ocean. The bottom water formed in the source region should correspond to AABW (Antarctic Bottom Water) or its prototype. In the Paleocene, the $\delta^{18}\text{O}$ and $\delta^{13}\text{C}$ values in the northern part of the South Atlantic and Indian Oceans also are remarkably high. This water may be

"Warm Saline Deep Water (WSDW)". High $\delta^{13}\text{C}$ values are recognized at mid latitudes ($\sim 30^\circ\text{S}$) from the Miocene onwards. This is probably caused by the influence of AAIW (Antarctic Intermediate Water). During the Miocene, carbon isotopes showed high values at 10°S in the northwestern Indian Ocean. This may be due to the influence of TISW (Tethyan Indian Saline Water; Woodruff and Savin, 1989).

(5) The circulation patterns of deep water within the Cenozoic have been reconstructed from isotopic evidence. The gradient of oxygen and carbon isotopic values reveals that Proto-AABW flowed northward from the Southern Ocean, and that WSDW flowed southward probably from the Tethyan Sea. WSDW is characterized by high temperature and high salinity. WSDW may have formed in the shallow Tethyan Sea associated with high evaporation. This water mass rapidly reduced with the closing of the Tethyan Sea across the Paleocene/Eocene boundary, and disappeared at about 50 Ma in the Indian Ocean, but it may have developed until 40 Ma in the South Atlantic Ocean. After 56 Ma, the ^{13}C latitudinal gradient shows that Proto-AABW flowed into the South Atlantic and Indian Oceans from high latitude, including the present Weddell Sea region. During the Paleocene and early Eocene, an "Oxygen isotopic minima belt" is recognized at mid-latitudes in the South Atlantic and Indian Oceans. A sharp change in oxygen isotopic ratios and a low value zone of carbon isotopic ratios formed at 1000-2000 m depths in the water column of the northern Indian Ocean after 30 Ma. This sharp change may indicate the formation of the AAIW/Proto-AABW boundary. AAIW may have flowed northward from high latitude judging from the northward decrease of the carbon isotope gradient. NADW (North Atlantic Deep Water) rapidly developed during the early/late Pliocene boundary time (~ 3 Ma), and affected AABW and CPDW.

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Appendix A-10. (continued).

Core & Interval						Specimen size		$\delta^{18}\text{O}$ $\delta^{13}\text{C}$ $\delta^{18}\text{O}$ $\delta^{13}\text{C}$				Core & Interval						Specimen size		$\delta^{18}\text{O}$ $\delta^{13}\text{C}$ $\delta^{18}\text{O}$ $\delta^{13}\text{C}$																																																																																																																																																																																																																																																																																																																																																																																																																																																																					
Section	Depth (mbsf)	Age (Ma)	D (μm)	T (μm)							Section	Depth (mbsf)	Age (Ma)	D (μm)	T (μm)							Section	Depth (mbsf)	Age (Ma)	D (μm)	T (μm)																																																																																																																																																																																																																																																																																																																																																																																																																																																															
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<i>Stensioina beccariiiformis</i> (Cont.)										<i>Nuttallides truempyi</i> (5 specimens) (Cont.)																																																																																																																																																																																																																																																																																																																																																																																																																																																																															
31X-1	0.75-0.80	286.65	61.65	455	228	0.108	1.116	0.108	1.116	27X-3	0.68-0.73	406.18	57.71	228-278	152-202	-0.701	-0.088	27X-4	0.70-0.75	407.70	57.87	253-329	152-202	-0.592	0.189	28X-1	0.71-0.76	412.71	58.35	228-342	152-215	-0.418	0.826	29X-1	0.76-0.80	422.26	59.27	278-304	152-177	-0.464	1.312	29X-2	0.69-0.74	423.69	59.41	304-329	164-202	-0.435	1.872	29X-3	0.69-0.73	425.19	59.55	304-329	164-202	-0.524	1.554	30X-1	0.69-0.74	431.69	60.18	266-304	152-177	-0.065	1.812	31X-1	0.70-0.74	441.20	60.74	329-354	190-253	-0.248	1.490	31X-3	0.67-0.71	444.17	60.88	278-342	164-177	-0.058	1.749	31X-5	0.68-0.72	447.18	61.02	278-329	164-190	-0.299	1.713	32X-1	0.65-0.69	450.65	61.18	278-329	164-190	-0.138	1.286	33X-1	0.71-0.75	460.21	61.61	266-291	152-164	-0.286	1.222	33X-3	0.69-0.73	463.19	61.67	278-342	164-215	-0.383	1.259	33X-5	0.71-0.75	466.21	61.73	278-329	164-190	-0.260	1.059	34X-1	0.61-0.65	469.61	61.80	278-316	164-190	-0.395	1.312	34X-3	0.65-0.69	472.65	61.86	304-316	164-190	-0.258	0.873	34X-5	0.65-0.70	475.65	61.92	291-329	164-190	-0.405	1.109	35X-1	0.68-0.73	479.18	61.98	253-329	164-202	-0.972	0.856																																																																																																																																																																																																																																																																																																																																
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26X-1	0.70-0.73	398.70	56.91	443	278	-0.720	0.475	-0.720	0.475	24X-1	0.77-0.82	379.77	55.74	481	278	-0.628	-0.547	29X-1	0.76-0.80	422.26	59.27	734	253	-0.236	1.605	-0.525	1.368	29X-1	0.76-0.80	422.26	59.27	392	202	-0.764	1.240	29X-1	0.76-0.80	422.26	59.27	417	228	-0.576	1.258	29X-2	0.69-0.74	423.69	59.41	493	266	-0.396	1.922	-0.397	1.946	29X-2	0.69-0.74	423.69	59.41	468	253	-0.281	1.996	29X-2	0.69-0.74	423.69	59.41	430	266	-0.513	1.920	29X-3	0.69-0.73	425.19	59.55	367	215	-0.340	1.688	29X-3	0.69-0.73	425.19	59.55	531	329	-0.602	1.583	-0.600	1.718	30X-1	0.69-0.74	431.69	60.18	392	215	-0.084	1.724	29X-3	0.69-0.73	425.19	59.55	455	215	-0.754	1.653	31X-1	0.70-0.74	441.20	60.74	430	240	-0.410	1.568	-0.318	1.502	29X-3	0.69-0.73	425.19	59.55	531	278	-0.444	1.919	31X-1	0.70-0.74	441.20	60.74	392	215	-0.225	1.435	30X-1	0.69-0.74	431.69	60.18	481	278	0.083	1.824	-0.120	1.750	31X-3	0.67-0.71	444.17	60.88	557	278	-0.076	1.618	-0.234	1.571	30X-1	0.69-0.74	431.69	60.18	493	354	-0.266	1.671	31X-3	0.67-0.71	444.17	60.88	405	228	-0.149	1.592	30X-1	0.69-0.74	431.69	60.18	468	240	-0.178	1.756	31X-3	0.67-0.71	444.17	60.88	430	240	-0.476	1.502	30X-2	0.69-0.74	433.19	60.32	557	278	-0.001	1.794	0.038	1.820	31X-5	0.68-0.72	447.18	61.02	417	228	-0.460	1.722	-0.167	1.748	30X-2	0.69-0.74	433.19	60.32	380	177	0.128	1.880	30X-2	0.69-0.74	433.19	60.32	380	202	-0.012	1.786	30X-3	0.69-0.74	434.69	60.43	607	304	0.210	1.679	0.165	1.649	31X-5	0.68-0.72	447.18	61.02	443	253	0.120	1.691	30X-3	0.69-0.74	434.69	60.43	468	253	0.120	1.618	32X-1	0.65-0.69	450.65	61.18	493	253	-0.061	1.403	-0.021	1.420	31X-1	0.65-0.69	450.65	61.18	405	202	0.072	1.596	31X-1	0.70-0.74	441.20	60.74	380	202	-0.351	1.533	-0.351	1.533	32X-1	0.65-0.69	450.65	61.18	380	215	-0.073	1.260	31X-3	0.67-0.71	444.17	60.88	430	202	-0.068	1.820	-0.143	1.654	33X-1	0.71-0.75	460.21	61.61	455	266	-0.075	1.344	-0.106	1.295	31X-3	0.67-0.71	444.17	60.88	405	228	-0.194	1.564	31X-3	0.67-0.71	444.17	60.88	417	228	-0.166	1.577	31X-5	0.68-0.72	447.18	61.02	380	202	-0.318	1.776	-0.278	1.781	33X-1	0.71-0.75	460.21	61.61	405	228	0.015	1.303	31X-5	0.68-0.72	447.18	61.02	417	228	-0.238	1.785	31X-5	0.68-0.72	447.18	61.02	417	228	-0.238	1.785	32X-1	0.65-0.69	450.65	61.18	380	240	-0.282	1.322	-0.345	1.319	33X-1	0.71-0.75	460.21	61.61	380	215	-0.065	1.377	32X-1	0.65-0.69	450.65	61.18	392	177	-0.408	1.316	33X-1	0.71-0.75	460.21	61.61	443	253	-0.555	1.290	33X-1	0.71-0.75	460.21	61.61	455	266	-0.600	1.291	33X-3	0.69-0.73	463.19	61.67	607	342	-0.135	1.111	-0.312	1.156	33X-5	0.71-0.75	466.21	61.73	430	253	-0.323	1.412	-0.176	1.241	33X-3	0.69-0.73	463.19	61.67	405	240	-0.234	1.236	33X-5	0.71-0.75	466.21	61.73	443	240	-0.163	1.265	34X-1	0.61-0.65	469.61	61.80	405	228	-0.129	1.383	-0.129	1.383	34X-3	0.65-0.69	472.65	61.86	417	228	0.040	1.182	0.040	1.182	34X-5	0.65-0.70	475.65	61.92	405	190	-0.678	1.084	-0.404	1.108	34X-5	0.65-0.70	475.65	61.92	405	177	-0.430	1.080	34X-5	0.65-0.70	475.65	61.92	405	215	-0.104	1.159
<i>Nuttallides truempyi</i> (5 specimens)										<i>Nuttallides truempyi</i> (5 specimens)																																																																																																																																																																																																																																																																																																																																																																																																																																																																															
23X-1	0.69-0.74	370.19	55.33	253-329	164-190	-0.770	0.040	24X-1	0.77-0.82	379.77	55.74	253-329	177-228	-1.055	-0.306	25X-1	0.72-0.76	389.22	56.14	253-304	152-202	-0.999	-0.324	25X-3	0.71-0.75	392.21	56.27	266-329	164-177	-0.768	0.064	26X-1	0.70-0.73	398.70	56.91	278-342	164-215	-0.855	-0.073	26X-2	0.71-0.75	400.21	57.07	278-342	177-215	-0.976	-0.489	26X-3	0.76-0.81	401.76	57.24	253-304	152-202	-0.607	0.073	26X-4	0.69-0.74	403.19	57.39	253-316	152-228	-0.839	-0.134	27X-1	0.69-0.74	403.19	57.39	278-304	164-177	-0.847	-0.155	27X-2	0.69-0.74	404.69	57.55	266-304	152-177	-0.803	-0.011	26X-5	0.70-0.75	404.70	57.55	278-329	164-177	-0.567	-0.014																																																																																																																																																																																																																																																																																																																																																																																																		

Appendix A-11. (continued).

Core & Interval Section	Depth (mbsf)	Age (Ma)	Specimen size		$\delta^{18}\text{O}$	$\delta^{13}\text{C}$	$\delta^{18}\text{O}$ ave.	$\delta^{13}\text{C}$ ave.	Core & Interval Section	Depth (mbsf)	Age (Ma)	Specimen size		$\delta^{18}\text{O}$	$\delta^{13}\text{C}$	$\delta^{18}\text{O}$ ave.	$\delta^{13}\text{C}$ ave.
			D (μm)	T (μm)								D (μm)	T (μm)				
Planktonic Foraminifera									Leg 121 Site 757 Hole B								
Leg 121 Site 752 Hole A									Subbotina spp. (6 specimens) (Cont.)								
<i>Acarinina primitiva</i> (5 specimens)									15H-1 0.70-0.75 130.50 38.55 253-329 202-253 0.807 1.595								
13X-1 0.70-0.75 113.60 55.27 329-380 278-304 -1.640 1.685									15H-5 0.70-0.75 136.50 40.98 253-329 177-228 0.927 1.571								
13X-4 0.62-0.67 118.02 55.46 329-380 278-329 -1.410 1.074									16H-1 0.70-0.75 140.20 41.89 253-329 190-278 0.792 1.581								
14X-1 0.70-0.75 123.30 55.70 342-380 291-304 -1.661 -0.101									16H-5 0.70-0.75 146.20 43.38 253-329 177-228 0.616 1.759								
14X-3 0.70-0.75 126.30 55.83 329-380 304-329 -1.632 1.209									17H-5 0.70-0.75 155.80 45.88 266-304 202-304 0.451 1.540								
15X-1 0.70-0.75 133.00 56.12 329-354 278-304 -1.510 0.592									18H-1 0.70-0.75 159.50 47.09 215-278 177-202 0.379 1.362								
15X-5 0.70-0.75 139.00 55.75 329-380 278-304 -1.595 1.634									18H-5 0.70-0.75 165.50 48.67 253-278 190-228 0.440 1.563								
16X-1 0.70-0.75 142.70 55.83 329-354 278-354 -1.547 1.371									19H-1 0.70-0.75 169.20 49.24 266-329 202-253 0.222 1.273								
16X-5 0.25-0.30 148.25 55.95 329-405 304-329 -1.798 1.845									20X-1 0.70-0.75 175.40 49.90 278-329 215-253 -0.167 1.391								
17X-1 0.70-0.75 152.40 56.04 329-354 278-304 -2.160 0.377									20X-5 0.70-0.75 181.40 51.37 278-329 202-228 -0.125 1.432								
17X-3 0.70-0.75 155.40 56.11 329-405 278-329 -1.689 1.828									21X-1 0.70-0.75 183.40 51.85 278-304 202-228 -0.399 1.386								
<i>Morozovella marginodentata</i> (5 specimens)									21X-4 0.70-0.75 187.90 52.70 253-304 177-228 -0.505 1.452								
15X-1 0.70-0.75 133.00 56.12 316-392 164-202 -1.794 2.631									22X-1 0.70-0.75 193.10 53.05 228-278 177-202 -1.029 1.063								
16X-1 0.70-0.75 142.70 55.83 304-329 228 -1.662 1.751									22X-3 0.70-0.75 196.10 53.25 253-329 202-253 -0.616 1.424								
16X-5 0.25-0.30 148.25 55.95 304-354 177-202 -1.650 3.383									23X-3 0.70-0.75 205.70 54.19 253-329 202-253 -0.942 0.608								
17X-1 0.70-0.75 152.40 56.04 329-342 190-228 -2.291 2.047									24X-1 0.58-0.61 212.28 55.27 266-329 202-253 -1.026 0.545								
17X-3 0.70-0.75 155.40 56.11 278-380 177-228 -1.540 3.729									24X-4 0.58-0.61 216.78 56.01 278-304 177-228 -1.323 0.673								
18X-1 0.71-0.75 162.11 56.74 253-380 152-228 -1.724 2.744									Leg 121 Site 758 Hole A								
18X-2 0.67-0.70 163.57 56.91 278-380 177-215 -1.810 2.868									<i>Acarinina praecursoria</i>								
19X-1 0.70-0.75 171.80 57.82 354-380 228-253 -1.677 3.225									31X-5 0.75-0.80 292.65 63.62 557 291 -1.154 2.092 -1.201 2.023								
19X-3 0.75-0.79 174.85 57.91 291-329 202-228 -1.825 2.773									31X-5 0.75-0.80 292.65 63.62 455 342 -0.777 1.614								
20X-1 0.70-0.75 181.40 58.09 304-329 202-253 -1.520 3.618									31X-5 0.75-0.80 292.65 63.62 557 405 -1.672 2.363								
<i>Subbotina spp.</i> (5 specimens)									<i>Acarinina primitiva</i>								
17X-3 0.70-0.75 155.40 56.11 278-316 202-253 -1.409 0.716									28X-2 0.75-0.80 259.15 58.18 430 329 -1.804 4.655 -1.596 4.519								
18X-1 0.71-0.75 162.11 56.74 240-291 177-215 -1.460 1.174									28X-2 0.75-0.80 259.15 58.18 481 354 -1.226 4.481								
18X-2 0.67-0.70 163.57 56.91 240-291 190-228 -1.324 1.123									28X-2 0.75-0.80 259.15 58.18 455 329 -1.757 4.420								
19X-3 0.75-0.79 174.85 57.91 253-304 177-228 -1.873 0.999									<i>Morozovella velascoensis</i>								
20X-1 0.70-0.75 181.40 58.09 278-304 215-253 -1.622 1.777									28X-2 0.75-0.80 259.15 58.18 632 278 -1.463 4.968 -1.460 4.867								
21X-1 0.70-0.75 191.10 58.36 253-304 202-253 -1.821 2.478									28X-2 0.75-0.80 259.15 58.18 607 304 -1.328 4.855								
21X-1 0.70-0.75 191.10 58.36 253 177-202 -1.736 2.477									28X-2 0.75-0.80 259.15 58.18 708 316 -1.590 4.778								
22X-1 0.70-0.75 200.80 58.63 253-329 202-253 -1.437 2.561									28X-6 0.75-0.80 265.15 59.27 620 329 -1.843 4.987 -1.853 4.829								
22X-3 0.70-0.75 203.80 58.76 253-304 152-228 -1.554 2.556									28X-6 0.75-0.80 265.15 59.27 632 354 -2.126 5.017								
23X-1 0.54-0.56 210.34 59.04 253-291 202-228 -1.466 2.544									28X-6 0.75-0.80 265.15 59.27 531 278 -1.591 4.483								
24X-1 0.70-0.73 220.20 59.72 253-304 177-228 -1.259 2.917									29X-1 0.75-0.80 267.35 59.67 734 329 -1.693 5.445 -1.631 5.084								
25X-1 0.79-0.84 229.89 60.53 278-329 202-253 -1.205 2.499									29X-1 0.75-0.80 267.35 59.67 607 304 -1.642 4.840								
25X-3 0.79-0.84 232.89 60.74 253-329 202-228 -1.453 2.520									29X-1 0.75-0.80 267.35 59.67 607 304 -1.558 4.968								
26X-1 0.97-1.00 239.77 60.96 178-329 177-253 -1.383 2.304									29X-3 0.75-0.80 270.35 60.22 683 304 -1.635 5.048 -1.492 4.799								
26X-5 0.97-1.00 245.77 61.14 253-304 202-228 -1.342 1.923									29X-3 0.75-0.80 270.35 60.22 582 278 -1.399 4.851								
27X-1 0.70-0.75 249.10 61.24 253-329 190-228 -1.274 1.743									29X-3 0.75-0.80 270.35 60.22 557 278 -1.443 4.499								
27X-3 0.70-0.75 252.10 61.33 253 190-202 -1.501 1.837									30X-1 0.75-0.80 277.05 60.86 708 329 -1.501 5.093 -1.506 4.935								
28X-1 0.70-0.75 258.80 61.53 253-304 190-228 -1.137 1.936									30X-1 0.75-0.80 277.05 60.86 708 380 -1.664 5.228								
28X-5 0.70-0.75 264.80 61.72 253-304 202-228 -1.189 2.226									30X-1 0.75-0.80 277.05 60.86 683 329 -1.353 4.484								
29X-1 0.70-0.73 268.40 61.83 253-267 177-202 -1.753 2.292									30X-3 0.75-0.80 280.05 61.11 632 304 -1.783 4.115 -1.587 4.015								
29X-5 0.70-0.73 274.40 62.01 278-304 202-228 -1.870 2.046									30X-3 0.75-0.80 280.05 61.11 493 253 -1.551 4.168								
30X-1 0.73-0.76 278.13 62.12 278-304 190-228 -1.404 1.986									30X-3 0.75-0.80 280.05 61.11 531 354 -1.428 3.762								
31X-1 0.70-0.75 280.10 62.18 253-291 177-202 -1.598 2.026									31X-1 0.75-0.80 286.65 61.65 531 304 -1.362 3.763 -1.497 3.790								
31X-5 0.70-0.75 286.10 62.36 278-316 177-228 -1.345 1.652									31X-1 0.75-0.80 286.65 61.65 582 380 -1.595 3.971								
32X-1 0.70-0.75 289.40 62.46 278 177-202 -1.419 1.713									31X-1 0.75-0.80 286.65 61.65 557 304 -1.535 3.635								
33X-1 0.68-0.71 299.08 62.76 266-278 190-202 -1.683 1.518									<i>Subbotina pseudoeocaena</i>								
33X-3 0.57-0.60 301.97 62.85 253-329 177-202 -1.752 1.640									28X-6 0.75-0.80 265.15 59.27 544 380 -1.240 2.761 -1.223 2.540								
Leg 121 Site 752 Hole B									28X-6 0.75-0.80 265.15 59.27 481 354 -1.090 2.304								
<i>Globorotalia pseudobulloides</i> (8 specimens)									28X-6 0.75-0.80 265.15 59.27 519 380 -1.340 2.555								
10R-1 1.05-1.08 346.15 64.84 253-316 152-190 -2.020 2.150									29X-3 0.75-0.80 270.35 60.22 506 392 -0.977 2.496 -0.901 2.568								
10R-2 1.00-1.02 347.60 65.06 228-291 139-202 -2.760 2.162									29X-3 0.75-0.80 270.35 60.22 506 380 -0.978 2.784								
10R-3 1.12-1.15 349.22 65.30 266-367 139-190 -2.303 2.306									29X-3 0.75-0.80 270.35 60.22 493 380 -0.747 2.424								
10R-4 0.79-0.81 350.39 65.48 228-342 127-202 -2.219 2.219									30X-1 0.75-0.80 277.05 60.86 506 367 -0.603 1.892 -0.603 1.892								
10R-7 0.41-0.43 354.51 66.10 278-329 101-152 -2.377 1.595									<i>Subbotina eocaena</i>								
<i>Rugoglobigerina pennyi</i> (9 specimens)									30X-1 0.75-0.80 277.05 60.86 683 380 -0.660 2.026 -0.873 2.092								
11R-3 1.12-1.14 358.92 66.42 266-304 177-202 -2.538 3.016									30X-1 0.75-0.80 277.05 60.86 632 380 -0.867 2.049								
12R-1 1.04-1.07 365.44 66.96 253-316 152-202 -2.718 2.280									30X-1 0.75-0.80 277.05 60.86 632 430 -1.093 2.202								
12R-3 0.10-0.13 367.50 67.10 253-304 177-228 -2.599 2.513									30X-3 0.75-0.80 280.05 61.11 721 481 -0.653 1.828 -0.799 1.696								
12R-5 0.54-0.57 370.94 67.32 253-304 164-215 -2.988 2.759									30X-3 0.75-0.80 280.05 61.11 696 405 -1.022 1.721								
Leg 121 Site 757 Hole B									30X-3 0.75-0.80 280.05 61.11 670 380 -0.721 1.540								
<i>Subbotina spp.</i> (6 specimens)									31X-1 0.75-0.80 286.65 61.65 531 405 -0.756 1.437 -0.852 1.442								
14H-1 0.70-0.75 120.80 35.34 304-342 202-253 1.012 1.517									31X-1 0.75-0.80 286.65 61.65 506 329 -0.876 1.537								
14H-5 0.70-0.75 126.80 37.28 266-354 177-240 1.040 1.493									31X-1 0.75-0.80 286.65 61.65 531 316 -0.925 1.351								

Appendix A. (continued).

Core & Section	Interval	Depth (mbsf)	Age (Ma)	Specimen size		$\delta^{18}\text{O}$	$\delta^{13}\text{C}$	$\delta^{18}\text{O}$ ave.	$\delta^{13}\text{C}$ ave.
				D (μm)	T (μm)				
Leg 121 Site 758 Hole A									
<i>Subbotina sp.1</i>									
29X-1	0.75-0.80	267.35	59.67	531	380	-0.455	2.642	-0.711	2.593
29X-1	0.75-0.80	267.35	59.67	519	329	-1.044	2.657		
29X-1	0.75-0.80	267.35	59.67	519	354	-0.635	2.480		
<i>Subbotina spp.</i> (5 specimens)									
28X-2	0.75-0.80	259.15	58.18	304-316	240-253	-0.773	2.389		
28X-6	0.75-0.80	265.15	59.27	304-354	240-354	-0.942	2.515		
29X-1	0.75-0.80	267.35	59.67	253-329	202-253	-0.517	2.529		
29X-3	0.75-0.80	270.35	60.22	316-367	228-266	-0.445	2.313		
30X-1	0.75-0.80	277.05	60.86	253-329	177-253	-0.409	2.251		
30X-3	0.75-0.80	280.05	61.11	266-329	202-228	-0.213	1.950		
31X-1	0.75-0.80	286.65	61.65	278-329	202-253	-0.411	1.567		
31X-5	0.75-0.80	292.65	63.62	278-329	190-228	-0.317	1.538		

