

Doctoral Thesis

**Analyses, Dynamics and Fates of Reactive Oxygen Species
(ROS) in Natural Waters: Emphasis on Photochemical
Reactions**

(Summary)

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学 位 論 文 の 要 旨

論文題目 Analyses, Dynamics and Fates of Reactive Oxygen Species (ROS) in Natural Waters:
Emphasis on Photochemical Reactions.

(天然水中の活性酸素種の測定、動態、運命——特に光化学反応に関して)

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A brief introduction of the reactive oxygen species (ROS) – hydrogen peroxide (H_2O_2), hydroxyl radical ($\cdot\text{OH}$), nitric oxide ($\text{NO}\cdot$) and superoxide radical ($\text{O}_2^{\cdot-}$) – measured in some natural waters in this study was reported in chapter 1. H_2O_2 and $\text{O}_2^{\cdot-}$ mediate in redox chemistry of metals (Fe, Cu, Mn) and other chemical species in natural waters, thus making them important targets of many environmental studies. Moreover, $\text{O}_2^{\cdot-}$ is a precursor of H_2O_2 . Similarly, $\cdot\text{OH}$ is a potent oxidizing agent capable of degrading recalcitrant organic pollutants in natural waters. $\text{NO}\cdot$ is photoformed in natural waters containing nitrite and it could be a potential sink of $\text{O}_2^{\cdot-}$ due to their high reaction rate. The reaction of oxygen with photo-irradiated dissolved organic matter (DOM) in natural waters produces $\text{O}_2^{\cdot-}$. However, measurement of this important ROS ($\text{O}_2^{\cdot-}$) has been a perennial challenge due to lack of suitable analytical technique. This and other environmental issues relating to the ROS in natural waters were addressed in this study.

Chapter 2 presents a report on the monthly $\text{NO}\cdot$ measurements in the Kurose River at Higashi-Hiroshima city. Results show that $\text{NO}\cdot$ photoformation rate ranged from 0.01 to $35.4 (\times 10^{-10} \text{ M s}^{-1})$ with steady-state concentrations in the range of 0.02–68.5 ($\times 10^{-11} \text{ M}$). There was a strong correlation ($r^2 = 0.95$) between $\text{NO}\cdot$ photoformation rate and the nitrite concentration in the river showing NO_2^- as a major $\text{NO}\cdot$ precursor. On the average, 98% of the photoformed $\text{NO}\cdot$ came from river nitrite. The $\text{NO}\cdot$ lifetime ranged from 0.05 to 1.3 s in the river and remained fairly stable in the upstream and downstream. The $\cdot\text{OH}$ radical, which was quantified during the study, had a photoformation rate of 0.01 to $13.4 (\times 10^{-10} \text{ M s}^{-1})$ and a steady-state concentration of 0.04 to $119 (\times 10^{-16} \text{ M})$ with a lifetime that ranged from 0.3 to 23 μs . $\cdot\text{OH}$ only accounted for $\leq 0.0011\%$ of the total $\text{NO}\cdot$ scavenged, showing that it was not a major sink for river $\text{NO}\cdot$.

In chapter 3, monthly measurements of H_2O_2 and $\cdot\text{OH}$ in rainwater and the Kurose River were carried out in 2013. H_2O_2 concentrations in the rainwater and river were highly season-dependent. H_2O_2 concentrations in the rainwater varied from the lowest 0.03 μM in winter to the highest 14.3 μM in spring. In the year, estimated wet deposition of H_2O_2 in Higashi-Hiroshima city was 7.5 $\text{mmol m}^{-2} \text{ y}^{-1}$. In the Kurose River, H_2O_2 concentrations of 0.06–0.37 μM were measured. The lowest and highest concentrations were found in the winter and summer, respectively. There was good correlation between solar intensity and H_2O_2 concentrations measured in the rainwater ($r = 0.79$, $p < 0.01$) and the river ($r = 0.81$, $p < 0.01$), which indicate photoproduction as one of the major H_2O_2 sources in the natural waters. The $\cdot\text{OH}$ was 1 order of magnitude photoformed in the river ($\sim 10^{-10} \text{ M s}^{-1}$) than in the rainwater. H_2O_2 and the unknown (which probably include photo-Fenton reaction) were predominantly $\cdot\text{OH}$ photochemical sources in the rainwater, accounting for 0.2–48%, 43–84%, respectively. In contrast, NO_2^- was the major $\cdot\text{OH}$ source (49–80%) in the Kurose River, while H_2O_2 contribution was negligible ($< 1\%$).

In chapter 4, daytime flux of $\text{NO}\cdot$ at the Seto Inland Sea-atmosphere boundary was measured in

September, 2013 and June, 2014. The average daytime $\text{NO}\cdot$ concentration measured in the sea surface was $1.9 \times 10^{-11} \text{ mol L}^{-1}$ ($\sim 9.87 \times 10^{-9} \text{ atm}$), while that measured over the sea was $5.2 \times 10^{-10} \text{ atm}$. Using an average wind speed of 3.2 m s^{-1} measured over the Seto Inland Sea during the cruise, a daytime $\text{NO}\cdot$ efflux of $0.22 \text{ pmol m}^{-2} \text{ s}^{-1}$ from the sea to the atmosphere was estimated. Assuming 8 h of solar intensity per day, about $1.54 \times 10^6 \text{ g NO y}^{-1}$ efflux of $\text{NO}\cdot$ from the 23,000 km^2 Seto inland Sea surface was determined.

In chapter 5, a method was developed for the measurement of $\text{O}_2\cdot^-$ in sunlit seawater using a fluorogenic probe – 3',6'-(diphenylphosphinyl)fluorescein (PF-1). Reaction of the photoformed $\text{O}_2\cdot^-$ with PF-1 produced fluorescein, which was separated by isocratic HPLC and measured using a fluorescence detector at 490/513 nm (excitation/emission wavelength). The reaction rate constant of the probe with $\text{O}_2\cdot^-$ was pH-dependent: $(3.2\text{--}23.5) \times 10^7 \text{ M}^{-1} \text{ s}^{-1}$ at $\text{pH}_{\text{TOT}} 7.65\text{--}8.50$. The detection limit of $\text{O}_2\cdot^-$ photoformation rate was 1.78 pM s^{-1} . The method produced results that were consistent with those obtained in literatures when applied to ten (10) seawater samples from the Seto Inland Sea. The $\text{O}_2\cdot^-$ photoformation rates were $3.1\text{--}8.5 \text{ nM s}^{-1}$, with steady-state concentrations ranging $(0.06\text{--}0.3) \times 10^{-10} \text{ M}$. The method requires no technical sample preparation and can analyze large number of samples.

In chapter 6, $\text{O}_2\cdot^-$, $\cdot\text{OH}$, $\text{NO}\cdot$, and H_2O_2 , in surface seawater obtained from the Seto Inland Sea were determined by concerted photochemical measurements. $\text{O}_2\cdot^-$ was photoformed at rates (range of $10^{-10} \text{ M s}^{-1}$) 1–2 orders of magnitude higher than $\cdot\text{OH}$ and $\text{NO}\cdot$. About 27% of the $\text{O}_2\cdot^-$ was transformed to H_2O_2 . About 3% of the $\text{O}_2\cdot^-$ photoformed was consumed by $\text{NO}\cdot$, while 15–21% of the $\text{NO}\cdot$ was consumed by the $\text{O}_2\cdot^-$. Therefore, $\text{O}_2\cdot^-$ could be a major sink for $\text{NO}\cdot$ in the sea. Estimated consumptions of $\text{O}_2\cdot^-$ by the $\cdot\text{OH}$ and due to bimolecular dismutation were negligible. In the seawater, (+)-catechin- and humic-like fluorescent dissolved organic matter contributed 3.2–4.1% and 3.6–4.8% of the $\text{O}_2\cdot^-$ photoformation rates, respectively.

Chapter 7 gives a general discussion and major conclusions of this study. The four ROS in this study are ubiquitous in sunlit natural waters. Nitrite is a major photochemical source of $\text{NO}\cdot$ and $\cdot\text{OH}$ in natural waters. The $\text{O}_2\cdot^-$ concentration ($\sim 10^{-11} \text{ M}$) in the Seto Inland Sea would be sufficient for speciation of the resident metals and other redox-active compounds.